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## Holtavirkjun

**Geological Report** 

Geological Investigations<br>2001-2006





# Landsvirkjun

**Key page** 



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Landsvirkjun's project manager's signature

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#### **Summary**

The bedrock in the area is comprised by a series of basalts and conglomerates from the extinct Stóra Laxá central volcano. The series are comparatively impermeable and reasonably competent rocks. They are overlain by series of competent interglacial lava flows of varying thickness. They are little or un-altered. On top of the interglacial lavas there are series of Holocene silt and sand. The silt is sandy and considerably cemented even if the apparent UCS strength is low but the thin layer of sand and gravel on top of the silt is poorly or un-cemented. The youngest formation is the Þjórsá lava which covers a good portion of the intake pond and forms the foundation for part of the project components. The lava is competent in most places but is very thin near the dam area

Several active fissures have been mapped in the area, some of which were mapped after the earthquake of 17. June 2000. The quake had an epicenter just south of the project are. Several active fissures have been found near most of the project components.

The main sources of permeability can be traced to the scoria of the Þjórá lava and to secondary permeability due to faults and especially active fissures.

The diversion structure and the weir will be founded on the Þjórsá lava. The lava is rather thin near the dam site and leakage paths to the gorge at Hamarinn are very short. The powerhouse and intake structure appears to be partly founded on old hydrothermally altered basalts and partly on younger interglacial basalts. Its present location appears to be near a fissure zone near the west flank of the Hamarinn. More fissure zones were found at the east flank. The mid section of Hamarinn appears to be least faulted site near the powerhouse. The tailrace canal will mostly be excavated in sand and gravel deposits near the Þjórsá river.

#### **1 General geology**

#### **Introduction**

The underlying part of the bedrock in the project area (Drawing 1) belongs to the Stóra-Laxá central volcano series. The Stóra-Laxá (SL) volcano was first identified and partly mapped by Ingvar B. Friðleifsson 1970 [20]. The centre of the volcano lies to the north and west of the project area. Orkustofnun updated the mapping 1976-1980 [11] as a part of a geothermal survey in the area. The report provided the first geological data used in the early studies of the present project area. A more detailed study of the area is found in a report published by Orkustofnun [5], see also geological map in drawing 3. The geological investigation was continued in 2000-2002 [1]. Geological observations in the area were intensified still in the present study, which started with a drilling campaign, which lasted from October 2006 to December 2006. Almenna Consulting Engineers performed the present investigation. The geologists Snorri P Snorrason, Gunnlaugur Þorbergsson, Melkorka Matthíasdóttir, Sigmundur Einarsson and geographer Áki Thoroddesen prepared this report. Ágúst Guðmundsson, from JFS Geological services, assisted and evaluated some details of the geological observation and performed magnetic and VLF surveys. Active fissures were mapped by The University of Iceland in 2000 and 2001 [15].

The present drilling campaign was mostly done in October and November 2006. The first holes were drilled on the Árnes Island but it was not fully completed as ice in the Þjórsá river made crossing untenable for the drillers crew. The diversion structure at Búði Waterfall and the canal from the intake and down to the intake lake was not investigated as planned. The investigation of the easternmost part of weir was not completed either.



*Figure 1-1. Þjórsá river (Árneskvísl) iced over 16. November 2006.*

#### **Regional geology, a brief outline**

The bedrock in the southern lowlands of Iceland was formed some 2-2.5 million years ago during a cold climatic period extending from 3.1 million years ago to the end of the Pleistocene ten thousand years ago. This era in Icelandic geology is often referred to as Plio-Pleistocene as it covers the Pleistocene and a part of the Pliocene eras of the Tertiary period. The climate in Iceland during the era is consistent with the Ice age of the Pleistocene.

The bedrock is characterized by series of basaltic lavas interbeded with relatively thin layers of conglomerate and sometimes irregular thick layers or heaps of hyaloclastite. The Stóra-Laxá central volcano dominates the volcanic pile to the north of the area. Activity in The Stóra-Laxá central volcano ceased in the Pleistocene era and after that erosion became the dominant geological process forming silt deposits in both marine and lacustrine environment [11]. The erosion process was interrupted by several Interglacial lava flows and a Holocene lava flow, all of which have their origins far outside the studied area.

#### **Plio-Pleistocene rocks**

#### Skarðsfjall tholeiite group (STG)

The oldest rocks in the project area are series of tholeiite and olivine tholeiite basalt layers and hyaloclastite, found at the base of Akbrautarholt where they are exposed and on the bank of Þjórsá River just downstream from the Búði waterfall. They are of the same age as the Stóra-Laxá central volcano and belong to its rock suite. These layers are not found at surface elsewhere in the project site, although basalt layers belonging to this group are found at the powerhouse site of Holtavirkjun. These rocks are hydrothermally altered and often intersected with dykes. They are generally highly jointed in the area.



#### **Interglacial formations**

MR

#### Interglacial silt deposits (LTS)

After the volcanic activity of Stóra-Laxá central volcano ceased a period began of alternating erosion of valleys and their filling of silt deposits and lava flows which had their origins outside the study area.

During this period the glaciers eroded a valley in SW direction from Búrfell down to the moraines at Búði waterfall. This valley is now covered by the Þjórsá-lava. Evidence of this buried valley is found to a certain degree in most of the boreholes that penetrate the Þjórsá-lava. The age of the valley is not entirely certain. The presence of interglacial lava in the area of Hvammsvirkjun overlying a sequence of silt layers shows that the valley has existed for a long time and its minimum age dates as far back as to the Eem Interglacial Period. The silt deposits observed in boreholes display a lower silt series of Pleistocene age (LTS) and it must be assumed that this older silt series is to be found in the general area near the present Þjórsá river. The possibility of encountering silt formations belonging to these series during construction phase of the project are however remote.

#### Akbraut Interglacial Basalt (HIB)

The main river course in the area has probably been very close to the present location of the Þjórsá river for a long time. The cube jointed interglacial basalt in boreholes near Skarðsfjall indicates its presence between Skarðsfjall and Núpur. The cubic joint pattern shows the presence of water (probably in the form of a river flowing forth on top of hot lava). The interglacial lava found in this area is either almost fresh or slightly altered cube jointed olivine basalt. It is very similar to the Akbraut interglacial basalt found in the area of Holtavirkjun.

The simplest relation between the interglacial lavas in the aforementioned holes and in Holt district is that they represent the same formation or lava flows that have descended through the pass between Skarðsfjall and Núpur across the silt-covered valley to south towards Holt district were it is found today. The river would have kept on flowing down on top of the lava, cooling it at first but later eroding it with time. The glaciers of the Pleistocene period have later contributed to the process. The Akbraut Interglacial lava is found in Akbrautarholt, at the powerhouse site and also below the dam site in cored hole NK-38. It is not found in NK-36, but it is almost 40 m thick in NK-40. These observation shows that the layer is discontinuous and its thickness varies greatly.

#### Hreppar Interglacial Tholeiite

These layers are found in the hills of Flagbjarnarholt just east of the intake construction. They consist mainly of fresh or little altered tholeiite lavas sometimes columnar jointed or cube jointed. The low weir at the diversion structure might rest against these rocks.

#### **Holocene formations**

#### Silt and tillite series (UTS)

The valley described in previous chapters could have reached from the sea at Eyrarbakki and possibly as far up-land as Búrfell. Kristján Sæmundsson, 2001 (pers com) has pointed out thick layers of silt under the Þjórsá-lava in the area east of Hestfjall. The moraine at Búði is mostly reasonably cemented siltstone and the un- or poorly consolidated silt in area of Hvammsvirkun indicates that the glacier has retreated from the Búði moraine at the end of the last glaciation [8], forming a laccustrine lake (similar to the present Breiðamerkurlón) between the moraine and the glacier in the process. The UTS serie is found below the Þjórsá lava at the weir site and it is found in the river bed of Þjórsá River at the main dam site. The silt layer is about 10-16 m thick in cored holes NK-36, NK-37, NK-38 and NK-40, but only some 3 m in NK-39. Silt and siltstone of same age is also found under loose gravel at the tailrace site, as can be seen in cored holes NK-43 and NK-44.



#### Tephra sand (TS)

After the glacier retreated from the area, leaving a lake mostly filled with silt, Þjórsá eroded a river course through the area most probably into the old riverbed from the last interglacial near its present location. Soon after that a thick layer of sand was deposited all over the area. The sand appears to be water-born black tephra with occasional clear plagioclase crystals. Such a thick layer (15-30 m) of tephra-rich sand suggests a big subglacial eruption, followed by a glacial outburst ("Jökulhlaup") in the upper reaches of Þjórsá/Tungná. The tephra sand has spread evenly over the area upwards of the Búði moraine possibly reaching all the way to Rangá in the east and further down the riverbed of Þjórsá. See further reading on the origins of the layer in Hvammsvirkjun Geological Report 2007. The sand was found in most boreholes that penetrate the Þjórsá lava, drilled near Hvammsvirkjun but is mixed with gravel near the former riverbed of Þjórsá river. A thin layer of the sand (2-5 m) was found directly under the Þjórsá lava at the dam site in holes NK-37, NK-38, NK-39 and NK-40 (For location of boreholes see Drawing 4).

Sand with the same description and chemical composition has been found under the lava in Sultartangi reservoir and in the area between Búðarháls and Ósöldur, near Þórisvatn in upper reaches of Tungnaá. The chemical composition of the sand found at Núpur and Búðarháls area [10] is very similar to the composition of the tephra found in Lake Saksunarvatn in Faeroe Islands [12], see Fig 1.2. This event has levelled out all minor irregularities in the area and it has left a flat plain of sand, only disturbed by local river erosion in the area best visible in the former riverbed of Þjórsá where the river has carved a passage in the loose sand east of Þjórsárholt from NK-2 down to NK-14 and down to the present passage at the Búði waterfall. This event has laid a foundation for the great extension of the Þjórsá-lava, erupted one or two thousand years later. Similar sand has been discovered elsewhere under the lava. The area of Selfoss and Eyrarbakki (Johnny Símorarson 2000, pers com) could be mentioned and similar sand has been discovered in the Urriðafoss area in boreholes at the dam site (the hole ULO-16 at 34 m depth) [2]. Chemical analysis has not been carried out but the visual appearance of the sand is similar. The conclusion is that the sand can be expected anywhere under the lava but its thickness could vary greatly. Near the old river course the sand would be mixed with gravel of various grain sizes.



*Figure 1-2. Comparison of chemical components of tephra sand at Búðarháls, Núpur and Lake Saksunarvatn.* 

#### Þjórsá lava (ÞL)

The latest event in the geological history of the area is the Þjórsá lava flow. Originating in an eruption in the Veiðivötn area some 8700 years ago, it is the largest postglacial lava known in Iceland. It flowed down from the Veiðivötn area to the coast at Eyrarbakki, a distance of 130 km. [9]. The lava covers the flat land west of the Búði moraine to Kálfá in the north and almost to Ytri Rangá in the east. A large portion of the Þjórsá lava flowed through the Búði passage along the old rivercourse of Þjórsá River. The Þjórsá lava is rather thin at the site of the weir of Holtavirkjun especially near the intake. It seems likely that the main lava flow has descended down the old riverbed and the thickness of the lava in the cored holes indicate that the old river

course has been north of the weir and also north of the cube jointed tholeiite in Þinghólar. (see Drawing 8). The lava is rather thin near the main dam at the gorge of Þjórsá River (down to 4 m in NK-40) and in NK-37 it was not found. After the eruption the river found a new path down at Hestafoss and it flowed partly over the lava and eroded away the top scoria of the lava (see drawing 8 and fig 1-3). The gorge at Hamarinn was eroded later.



*Figure 1-3. The Þjórsá lava a the dam and weir site. The scoria of the lava has been eroded away by the river.* 



*Figure 1-4. The Þjórsá river gorge at Hamarinn and the weir site in background* 



*Figure 1-5. Thickness of scoria and soil in Þjórsá Lava* 

#### **2 Tectonics**

#### *Tectonic setting*

The project area is situated in an active tectonic area in southern Iceland. The area has been known for its earthquakes and fissures for centuries. Earthquakes have occurred in several places in the area with a frequency of up to one hundred years between episodes. The magnitude of the earthquakes can reach 7 on Richter scale or possibly higher [18].

Many faults have been observed in the area and also several fault orientations. Three set of faults are most frequent: Old north-easterly-directed faults that are associated with the period of active volcanism in the area and younger active faults oriented in a northerly direction, are associated with the South Icelandic seismic zone. The third set which has a variable direction of N70°A to N105°A is probably also related to South Icelandic seismic zone. Faults of various types have been found, both strike–slip and normal [5].

#### **Active faults and fissures**

In the years 2000-2001, the Science Institute of the University of Iceland mapped active fissures related to the South Iceland seismic zone [15].

The location of the fissures is shown on drawing 9

The appearance of the fissures can be highly variable. They are easily identified in some areas for example at Minnivellir, but the fissure sets to the east from Minnivellir tend to be obscure. Active fissures in the project are reasonably mapable.

Several sets of active fissures were found in the project area and they all have the common north-south direction: The easternmost set (1) strikes at the farm Vindás just east of Búðafoss (Drawing 9). These fissures will not affect the construction site.

Fissure set (2) strikes at Búðafoss near the fish farm. The geothermal field at Laugar is almost certainly connected to the fissure set. A few fissure features near this set were mapped by AV. They show northeasterly direction.

A fissure set (3) strikes near Hestafoss and the farm Lækur. These fissures are part of the fissures related to the earthquake of June 17th 2000 just south of the project area. The fissures at Hestafoss are quite obvious and signs of movements were visible in the area just after the quake on both sides of the river, but they have not been included in previous mappings. They are quite narrow and display irregular openings of ~5 cm and sometimes up to ~10 cm. A fissure was found near the river just north of the Lækur farm and another one between the farm houses (Andrés Eyjólfsson 2007 pers.com). Fissures of this set can be traced some distance to the north from Hestafoss in the island Árnes. According to the current plan the weir will cross these fissures. Springs related to the fissures can be seen just below Hestafoss.

Fissure set (4) is found south of the farm Akbraut and is not clearly defined and it can be divided into three subsets of fissures. One is just east of Akbrautaholt and runs near Þinghólar in Árnes.

The second subset can be traced to the south from the Akbraut farm just west of the road to Akbraut. The northernmost fissures of this recently active set, were mapped some 1200 m to the south of Akbraut. It is likely that their continuation can be found at the eastern cliff edge of Akbrautarholt close to the proposed powerhouse site and possibly crossing the gate section of the main dam in the gorge of Þjórsá, just north of Akbraut farm.

The third subset is located just west of the Akbraut farm and it appears to cross the tailrace canal. Weak signs were found in a fissure survey at the Akbraut farm in fissure trench AK-1, indicating the presence of a fault there. Signs of recent movement were not found.

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The fissure sets (3) and (4) were active in the earthquake of June 17th 2000.



*Figure 2-1. Earthquake fissures at Hestafoss. Fissures have also shaped the riverbed.* 

#### **Fissure survey**

Active fissures were mapped by the University of Iceland in 2000-2001 in the project area [15]. The mapping is the main source of information regarding the fissures. Some additional investigations were carried out as a part of the geological investigation of this report. The result is shown on drawing 9. The fissures present at Hestafoss were traced in gravel deposits by the river soon after the earthquake (of June17th 2000) and in open bedrock fissures by the farmhouse of the farm Lækur (Andrés Eyjólfsson 2007 pers.com). A spring wells out of such fissure (~10-20 l/s) at Hestafoss, see drawing 12. The leak path is however very short.



*Figure 2-2. Fissure trench AK-1 at Akbraut farm.* 



*Figure 2-3. Location of fissure trenches at the farm Akbraut.*

Several fissure trenches were excavated in the western slope of Hamarinn. The overburden was mostly eolian sand and fine grained gravel from the river. Neither of those leave any indication of fissure movement in the past. The groundwater is high and it was not possible to dig the trenches to full length because of inflow of groundwater. No direct evidence of fissures was found in any of the trenches. Steep face in conglomerate was found in the fissure trench AK-1 just where it meets the slope uphill. The sign was rather steep rockface, probably a fault face. No signs of recent movement were found in any of the trenches. The overburden was not suitable to preserve signs of fissures. In AK-1 and AK-2 the soil was mostly sand and gravel and in AK-3 and AK-4 sand and stony peat with frostlifted stones of the bedrock and till.

The possibility of digging fissure trenches in the eastern slopes, south of Hamarinn was considered but the slope is too steep and the excavator would have to stand in the river with a limited reach. The survey was abandoned on the east side and decided to investigate the fissures by drilling an inclined hole at the west end of the damsite (NK-42).



*Figure 2-4. Fissure trench AK-1 just after excavation. Unsure signs of a fissure or fault are in the shadow at the light spot just above the inflowing water.*

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*Figure 2-5. Fissure trench AK 3 in process of excavation. The inflow of groundwater made the study difficult* 

## **3 Lithology**

The classification of the bedrock in the Holtavirkjun area is based of field observation and inspection of hand specimens and cores from the holes.

#### **Basalt**

#### Plio-Pleistocene basalts

Each basalt lava flow has typically a three way subdivision, top scoria (often some 10-25% of the layer), dense or semi vesicular middle part (frequently 70-80% of the lava) and a layer of bottom scoria 3-8%. The top

scoria is often well consolidated. It consists of scoriaceous and vesicular basaltic fragments. It often contains sedimentary infiltration of silt or sandstone from overlying layers. The top scoria is often hydrothermally altered and in most cases the alteration increases the degree of cementation. The crystalline middle segment of the lava consists in many cases of hard dense rock with irregular joint system derived from cooling of the lava. In addition to the cooling joints there is flow cleavage related joints and tectonic joints. Bottom scoria is usually free from sedimentary fillings and is most often fairly well consolidated.

The basalts are classified into three main types, tholeiite lava, olivine basalt (olivine tholeiite) and porphyritic basalt (G.P.L. Walker 1959). The olivine basalt is in most cases olivine tholeiite. It should be pointed out that these three types of basalts form a complete transition between them. The secondary minerals serve best as indicators in recognizing each type. The basalt layers at Holtavirkjun are mostly olivine basalt.

Olivine basalt occurs both as single flows and as a sequence of belted flows with scoriaceous or vesicular belts between the layers. Porphyritic basalts are basically tholeiite or olivine basalt with a various content of plagioclase phenocrysts and sometimes a small amount of olivine phenocrysts. Plagioclase phenochryst content higher than 5% is a common reference but if plagioclase crystals are rare in the area this number has a tendency to drop. Low contents of plagioclase content (ca 1-3%) is not significant in terms of engineering properties of the rocks but as phenocrysts content is increased, the number of primary joints often drop.

All rocks from the active period of Stóra-Laxá central volcano have been hydrothermally altered. Hydrothermal alteration as observed in the field and in the cores has most often reduced the strength of the crystalline rocks and increased the strength of the scoria parts. The powerhouse will most probably be founded in layers of this series and the tailrace canal will be excavated into these layers a short distance closest to the powerhouse.

The Olivine basalts display a medium or high strength with typical apparent UCS value in the range of 50-160 MPa. Permeability values are 8-13 LU.

The porphyritic basalts display a medium strength with typical apparent UCS value in the range of 50-80 MPa. Permeability values are 10-14 LU.

#### Interglacial lavas

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Olivine basalt of Akbraut interglacial basalt are series found below the Þjórsá-lava in drillholes NK-38, NK-40 and at the surface at Akbrautarholt and in a large area south of the project area.

This basalt is slightly or unaltered with dark fillings in joints. It is generally massive and with little scoria. Cube jointing is most prominent in the boreholes, but the rocks in outcrops at Akbrautarholt are highly jointed and the joints appear to be of tectonic origin. The thickness of this lava or series of lavas varies very much over a short distance, indicating river channels or similar landscape in the area at the time of formation. The intake channel, penstock and the powerhouse will be excavated and partly founded in this formation. The main dam will also be partly founded on it in the riverbed of Þjórsá.

Similar Interglacial lavas are found east of Árnes Island, in Flagbjarnarholt and to the south of these places. They appear to be older than the Akbraut intergalcial basalt. These layers are found in several places to the north of the project area in Þjórsárholt and Skaftholt and both Núpur and Miðhúsafjall. The lavas seem to have filled up shallow valleys in these places. Cube and columnar joints are prominent features of this series but no boreholes have so far been drilled into it. They appear to be competent rocks and have obviously withstood glacial erosion better than most layers in the area.

The Interglacial lavas display a medium or high strength with typical apparent UCS value in the range of 160-200 MPa, UCS 186 MPa. Permeability is in the range of 10-20 LU.

#### Holocene lava

The Þjórsá-lava is the only Holocene lava in the area. It has a very typical cross section of tholeiite lava described in the Plio-Pleistocene section. It is usually 20-30 m thick and has often a 5-10 m thick top-scoria. The top-scoria can be (in some cases) difficult to excavate with backhoe but a percussion drill penetrates the scoria fast or very fast. The crystalline part is of porphyritic olivine lava, with 5-10% plagioclase phenocrysts and occasional olivine crystals. It is usually 10 to 25 m thick and is normally solid and most often returns core well. But the lava is considerably thinner near Akbrautarholt were it can be as thin as 4 m in NK-40 and at the damsite it has flowed up to a silt bank and it is not found in cored hole NK-37. The scoria part is largely absent in the area next to Hamarinn (thickest 1 m in NK-38). The most plausible reason for this is that the scoria has



been eroded away when the Þjórsá river flowed over the area, see Drawing 8. The scoria has not yet been investigated in Árnes, east of hole NK-39, at the inner part of the weir or at the diversion at Búðafoss, but according to observation in the field the scoria appears to be thicker. The thickness of scoria in the Þjórsá lava is shown in fig 1-5.

The Þjórsá lava displays a rather mixed results with apparent UCS values of 50 and 60 MPa. UCS test yielded 138 MPa. Permeability is in the range of 10-50 LU for the solid part of the lava but the permeability of the bottom scoria varies. (Over 200 LU in NK-38 but only 14 in NK-40). It is not possible to measure the scoria separately. Part of the lava and the sand directly under the lava are always included the test.

#### **Dykes**

#### Basaltic dykes

Several basaltic dykes have been found in the cored boreholes. They appear to be some variation of thin and inclined cone sheets and/or inclined dykes. The inclination of the dykes varies. They tend to intrude the more soft layers as the scoria segments of the basalt layers, or conglomerates. The breaking strength of the dykes is high as the average point load varies from 60-100 MPa. The joint pattern is mostly inclined with rough and undulating joint surfaces. The most common dykes are made of grained dark olivine basalt somewhat finer grained towards the margins. The dykes are typically dense and slightly jointed or even joint free. Joints are rough and undulating. Hydrothermal alteration of the dyke is usually slight but joints are often coated with various secondary minerals. Zeolites, calcite (often Iceland spar) and pyrite are prominent. Clay minerals are less common in the dykes than in the surrounding rocks. The margins of the dykes are often fused or welded to the surrounding rock and the contact is displayed intact in the core.

The dykes display a medium or high strength with typical apparent UCS value in the range of 60-100 MPa, Permeability is not conclusive for the dykes.

#### **Sedimentary rocks**

#### Plio-Pleistocene siltstone and conglomerates

The few sedimentary layers found in the project area are generally fairly well cemented, fine grained conglomerate, competent rock. The fine-grained conglomerate is the most common type but sandstone, siltstone and coarser types of conglomerates are also known to be present in the general area. The breaking strength of the conglomerate varies from low strength to high (25-75 MPa). Permeability is rather low. Indicated permeability (tested with other formations) is 16 LU but that appears to be too high, 5 LU seems to be a more likely figure.

#### Interglacial tillite and silt

The interglacial silt layers are found in several boreholes under the Þjórsá-lava. They consist of a blend of silt and fine sandstone often layered with thin layers of fine sand. Their apparent UCS value is low (1-8 MPa in Hvammsvirkjun, measured parallel to layering) with weaker layers of sand in between. These layers are known to be present in the general area of the project but they have not been found in direct contact with the project components.

#### Holocene silt

The Holocene layer of silt and fine sand is found in several boreholes under the Þjórsá-lava (NK-36, NK-37, NK-38, NK-39 and NK-40). It is 10-15 thick in most holes and it is exposed in the Búði moraine and in banks along Þjórsá River below the Hamarinn and in cored holes NK-44 and NK-45. This layer or layers are often poorly cemented. They are generally layered with alternating layers of silt and fine sand.

The core tends to break along the sand layers. The layering and the joints are planar and smooth. In some cases the silt tends to become coarse near the base of the formation, resembling silty tillite. This coarser part is far more competent than the alternating silt and sand layers. The silt layer is found under the Þjórsá lava at

the damsite and the main dam will probably be partly founded on this layer in the Þjórsá gorge east of Hamarinn.

The Holocene silt displays a low strength with typical apparent UCS value in the range of 6-13 MPa (measured parallel to layering). Permeability is in the range of 10-15 LU, under the Þjórsá lava at the dam and weir site. Higher values, 70 LU were encountered in NK-40 (22-27,2m). These values indicate the presence of a fissure or unknown weakness near the hole.

#### Holocene tephra sand (TS)

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In several places under the Þjórsá-lava a thick layer of tephra rich sand has been found. No core has been recovered from the sand and its properties must mostly be deduced from recovered cuttings and general drilling operation during ODEX drilling.

The sand layer can be divided into three parts, each of varying thickness. One or more parts may not be present in any given section. At the bottom and top there is often a layer of loose sand and gravel displaying rounded material with high content of tephra. The middle section is of pure black tephra with grains of clear plagioclase crystals. The grain size is medium to coarse sand in most cases. The cuttings and drilling operation in general indicate a loose or poorly cemented layer. Drilling operation (ODEX) in NK-4 and 5 were particularly fast and fine-grained gravel (with high content of tephra) was observed in the top section of the sand. A thick layer of coarse tephra sand is exposed along the banks of Rangá underlying the Þjórsá-lava. Similar tephra rich deposits (poorly cemented sandstone and packed sand) are exposed along the bank of Þjórsá down from Búrfell [7], towards the project area but it disappears under the lava and alluvium from the river above the projected intake pond. The mineralogy and chemical composition of these layers has not been studied in detail but they suggest that the tephra sand below the Þjórsá lava in the present study area could be the similar or same. Similar deposits have also been observed along the banks of Rangá a considerable distance to the east.

The permeability of Holocene tephra sand and gravel was evaluated 10-20 LU (lowest test results of combined scoria and sand), as no direct test could be made. The core recovery of the sand was often very little and grain size of the recovered material varied. The evaluation is therefore uncertain.

## **4 Engineering geology**

#### **Logging of boreholes**

Detailed graphic logs of holes in the area are shown in Annex 1. The standard RQD values are given as well as the values for 30, 50 and 100 cm criteria, both for drilling intervals and rock units as detailed as possible. In addition to that location and results of point load tests are given along with results of UCS tests, permeability tests and rock mass quality Q-values. Photographs of the cored sector in boreholes are presented in Annex 2.



*Table 1 Cored holes at Holtavirkjun* 

#### **Field tests on core**

The drilling campaigns were carried out during winter, often in frost and snow. The core was transferred to a nearby house or directly to Reykjavík in a closed vehicle. Special care was made to protect the silt and other sediments from freezing. The RQD measurements were therefore sometimes carried out off site and Q value estimated simultaneously. Samples for UCS test and partly for point load test were selected on site and covered in plastic to maintain original moisture content as possible.

The results of the RQD measurements (summarized 10-, 30-, 50-, and 100 cm lengths) are given in each log in Annex 1 with Q values.

#### Point load tests

Point load tests were carried out as soon after drilling as possible. The core is 45 mm in diameter and conversion factors were used to calculate IS50 and the "apparent" UCS calculated accordingly. The results from the point load tests and calculated apparent UCS are shown in Annex 6.2. The overview of apparent UCS is summarized in the following histogram, see figure 4-1.

The conversion feature of 18 was used to calculate the apparent USC value. The conversion factor chosen is in good concordance with the individual basalt types but is maybe too high for the softer rocks, for instance the siltstone. It was however used in the case of the softer rocks as no other values were available.

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*Figure 4-1. Apparent uniaxial strength of different rock types, estimated from point load tests.* 



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*Figure 4-2. Length of tested intervals vs. permeability in Lugeon values.* 



*Figure 4-3. Distribution of permeability values.* 



*Figure 4-.: Lekage from top section of the Þjórsá lava into salmon waterway at Budi waterfall.* 

#### Laboratory tests on cores



*Table 2 Uniaxial compressive strength.* 

#### **Groundwater observation**

#### Temperature readings

After completion the temperature of the groundwater was measured in most of the holes (see Annex 5-1). Some of the holes were drilled using hot water (~60°C) as drill fluid. Some of the holes had not recovered after drilling when they were measured so relevant temperature measurements were not achieved but the temperature logs show variation related to permeability.

The inclined holes are difficult to measure because the thermometer tends to get stuck on the way down, increasing the possibility of loosing the instrument during the measurement.

The temperature readings in the holes show a low and even temperature of 4°C to 5°C in the Þjórsá lava, with little variation. One hole is different, NK-39. It shows very steep thermal gradient below ~25 m depth (below the Þjórsá lava). The gradient appears to be over 60°C/100 m. The total depth of the hole is only 32m so this thermal gradient can not be regarded certain. It nevertheless indicates a high gradient in the area. NK-39 is located near a fissure mapped as active in 2001, see drawing 9. NK-39 is also near a potential geothermal site in Árnes, see drawing 6. No temperature measurements exist for this site but a shallow hole was drilled there several years ago for investigation and lukewarm water was found there (enough to feed a small bathing pond).

The first temperature measurements at the powerhouse site were made right after drilling in NK-41 and NK-42. They show the effect of the hot water used in the drilling process. The temperature curves are irregular and show probably outflow (places of higher permeability) where the curves show reversed gradient. NK 42 was measured several days after completion. The measurement shows even and low temperature (Annex 5-1).

The temperature of NK-43 to NK-NK is normal for the site. The measurements were carried out during a cold spell and they reflect the weather directly.

High temperatures found in NK-14 (up in the Þjórsá-lava) suggest an intrusion of warm water into the groundwater in the vicinity of the hole as a lens of 14°C hot water is found at the groundwater surface and lower temperature and increasing with depth is found immediately below. The nearest known geothermal field is at the fish farm at Laugar close to the diversion structure at Búði. Active fissures are very close to the fish farm and the simplest explanation of the temperature anomaly in NK-14 is that the fissure zone extends to the vicinity of NK-14 and significant inflow of hot water flows up and intermixes with groundwater.

Temperature anomalies as described above are positive indication of open fissures filled with warm upflowing water, most likely connected to active faulting. Most of geothermal fields in the Southern Lowlands in Iceland are connected to a set of earthquake fissures similar to those found in the project area.

#### Electrical conductivity

Electrical conductivity was measured in the holes. Substantial difference was found. The water in the Þjórsá lava has low conductivity but the water below has substantially higher values. The high values are probably from the silt formation as it was deposited in seawater. The difference in conductivity shows that the groundwater in the Þjórsá lava is a separate groundwater system.

#### Groundwater level

Groundwater observations date from the drilling campaign. Variations in groundwater level have not been monitored. The groundwater system consists of four main parts.

- 1 The Þjórsá lava
- 2 Sand and gravel formation on top of the lava in Árnes
- 3 The bedrock in Hamarinn at the Akbraut farm
- 4 The sand and gravel formation west of the Hamarinn.

Considerable groundwater flows past the project area within the Þjórsá lava at Árnes. The temperature profiles in the holes show consistent low groundwater temperature (4-5°C) and conductivity of the water supports this. The water level within the lava has higher elevation than the river below Hestafoss. Little water flows into the gorge of Þjórsá below the damsite. Only few springs have been observed amounting to a few tens of l/s of water. No water has been observed flowing into the Þjórsá river north of Árnes island or below Hestafoss but the conditions were not favourable as the water level of the river was usually too high to allow any observation likely to succeed. The main flow of the Þjórsá lava was along the Árnes Island to the west (see drawing 8) and it is likely that the main flow of groundwater in the lava follows the old rivercourse as it did provide wet areas to allow formation of scoria at the bottom of the lava during the flow of the lava. The groundwater in the lava seems to have two feeders besides precipitation. Firstly, groundwater flow in the lava through the Búði pass and secondly, infiltration into the sand and gravel formation on top of the lava from the Árneskvísl branch of Þjórsá river just south of Búðafoss. The Búði pass is narrow and carries limited amount of groundwater as can be seen just above the waterfall at Vindás where substantial groundwater flow (~0,5-1  $m^3$ /s) is forced to the surface because of flow restriction in the lava at Búði pass.

The Árneskvísl Branch of Þjórsá river has higher elevation than the groundwater in the sand and gravel formations overlying the lava at the east section of Árnes island a distance of 1,5 km. The groundwater is high near the river (very near the surface) but lowers with a gentle gradient towards the west (see drawing 7), indicating infiltration of water from the river to the groundwater. Water can be seen in old channels of the river in the area seeping towards the river just above Hestafoss. Part of the water could infiltrate the top scoria of the lava along the scoria margin, see drawing 7. The Árneskvísl branch will be diverted during the construction. The groundwater table in Árnes is expected to lower as a result of the diversion.

Permeability in Hamarinn and at the powerhouse site is rather low and the groundwater system in the area is somewhat isolated. It is not in contact with groundwater in the Þjórsá lava or the sand and gravel deposits west of the Hamarinn. Little inflow of groundwater is therefore expected in the powerhouse pit except from the sand and gravel deposits in the tailrace canal site.

Extensive sand and gravel deposits are west of the Hamarinn hill. The main branch of Þjórsá River flowed here until few centuries ago and deposited the sand and gravel below the gorge. Close to the Hamarinn the deposits are thicker than 15 m but further downstream they are only few meters. The sand and gravel is rather uniform. Layer of silt was found under the sand at holes NK-43 to NK-45. The groundwater in the gravel seems to be governed by the water level of Árneskvísl. The Árneskvísl will be diverted during construction and therefore flow of water in the tailrace canal should dwindle with time. The inflow is expected to be substantial at first.

### **5 Geological conditions of the project components**

#### **Diversion structures at Búðafoss**

The diversion structure will be constructed in Árneskvísl near Búðafoss. It will mostly be founded on the Þjórsá lava, but the easternmost section of the structure, just below the fish farm at Laugar, will probably rest on interglacial lavas.

The depression in the glacial formations at Búði is the main pass through which the Þjórsá lava has flowed to the sea at Eyrarbakki. It is therefore likely that the lava is reasonably solid at this place, but the thickness of the lava is still unknown and the lava near the fish farm is likely to be thin. The underlying strata, which could be encountered at/or near the lava margin, are most likely combination of glacial silt/tillite or some gravel and sand. The cube jointed basalt layers in Flagbjarnarholt could be found near the eastern part of the diversion structure.

#### **Intake pond and lateral embankments**

The weir along the northern side of the intake pond will mostly be founded on the Þjórsá-lava, near the margin of the lava. The thickness of the lava varies (3 m in NK-40 and 8 m in NK-38, 16 m in NK-36, 21 m in NK-39 and it was not found in NK-37). The scoria on the lava is very thin and it seems that it has been eroded away by the river especially on the westernmost part near the damsite. A thin layer of silt was found on top of the lava in holes NK-36 and NK-37. A layer of gravel and sand (2-4 m thick) was found directly under the lava in most of the holes (not in NK-36). Under the gravel there is a layer of silt 10-15 m thick in most of the holes but only 3 m in NK-39. Apparent USC values are 4-10 MPa and permeability 10-15 LU. Below the silt is competent interglacial lava or reasonably sound bedrock which consists of basalt, conglomerate or dykes with low permeability (0,1-7 LU) and reasonably uniaxial strength even if it varies a lot (basalt and dykes~50-150 MPa and scoria and conglomerate 20-40 MPa) Several fissure zones strike the weir. Active fissures strike close to NK-39 and just west of NK-37 (found after the earthquake of June 17th 2000). Active fissures are most likely in the gorge just west of NK-40. Several small fissures at each site are more likely than just one larger fissure at each site.

#### **The dam**

The main dam and gate section will be constructed in the gorge east of Hamarinn. The west section of the dam will be founded on the highly jointed Akbraut Interglacial basalt and the east part on a layer of Holosen silt. The silt layer is found in the inclined hole NK-40 on the east bank of the gorge where it is 10 m thick. The silt has good core recovery in the upper part of the hole (~100%). The upper part of the layer is sandy and has high pebble content and could be classified as silty tillite but the lower part is horizontally layered and the core breaks along the layering. The apparent permeability of the silt is 10-15 LU except in NK-40 (test 22-27,5 m, possibly up to 70 LU) but the short leakage path to the gorge and nearby fissures just west of the test site could interfere with the results. The apparent USC value for the silt is 4-10 MPa A 4 m thick layer of sand was found in NK-40 and in other cored holes to the east but it is not found in the gorge or in NK-36. The sand layer is permeable and the leakage path is very short near the gorge. The highest permeability of the sand (and the bottom scoria of the lava) was detected in NK-38 (~200 LU). Below the silt and the sand, the Akbraut Interglacial basalt was found. It is sound basalt except where intersected by tectonic fissures. Typical point load figures are in the range of 160 to 200 MPa. The permeability was in the range of 5-20 LU. The basalt from the Skarðsfjall Tholeiite Group (STG) was found in NK-36, NK-39 and NK-40. It is mostly of sound basalt and conglomerates intersected with dykes. The permeability of the older basalt is near 10 LU but the tests are few. Fissures which opened in the earthquake of June 17th 2000 were found south of the damsite. They seem to strike in the gorge at the foundation of the main dam, see drawing 9 and 14. The fissures and faults at the damsite must therefore be classified as active.

#### **Approach channel, penstock and powerhouse**

The excavations for approach channel, concrete penstocks and powerhouse will be performed mostly in Akbraut interglacial basalt (AIB). The AIB layer forms Hamarinn itself. It is sound olivine basalt and little altered. Joints and fissures are covered with dark glossy clay minerals. It is highly jointed in at the east flank of Hamarinn and also just west of NK-41. The joints indicate probably fissure zones, one at the proposed powerhouse site and another (possibly several others) on the east flank of Hamarinn and in the gorge of Þjórsá (see drawings 10-12). The powerhouse will be founded on the scoracious lower part of the AIB formation and/or in the olivine basalts and conglomerate of the STG formation. The scoraceous part of AIB was only found in one hole NK-41, thus it is quite conceivable that the scoria thins out to the east and the area just east of the present powerhouse site seems to be considerably less fractured than the present site. The apparent uniaxial strength of the older rocks varies considerably. The basalt layers from 50-150 MPa and the scoria and conglomerate appear to have strength of 20-40 MPa. The AIB has high apparent USC strength of 180-220 MPa and the scoria is competent with apparent USC of 20-40 MPa. Despite the fissure zones the permeability is in the range of 10-60 LU where the most common values are 15-20 LU. Magnetic and VFL survey were carried out at the powerhouse site and they indicated an anomaly at the present powerhouse site. The anomaly was interpreted as a fault zone and the presence of it was confirmed by drilling NK-41. The most likely location of the fault zone is shown in drawings 10-12 and 14.



*Figure 5-1. Highly jointed basalt in Hamarinn*

Active fissures were found to the south of the Akbraut farm (see drawing 9) and they seem to strike along the east flank of Hamarinn and/or in the Þjórsá gorge at the damsite. The highest section of Hamarinn seems to be the least faulted site for the powerhouse.

#### **Tailrace canal and riverbed downstream of Akbraut**

The upper part of the tailrace canal will be excavated in similar formations as the powerhouse site but the lower part of the canal and the riverbed below the end of the canal will be excavated in the sand and gravel deposits at the river. A layer of fine sand and silt was found in NK-44 and NK-45, near the bottom of the proposed tailrace canal. The elevation of the silt is near 45 and 48 m a.s.l. respectively or just above the elevation of the proposed tailrace canal at NK-45. The canal site was not investigated farther downstream Nátthagahólar by drilling. According to study of areal photos the lava margin of Þjórsá lava lies to the north of the Árneskvísl branch of Þjórsá river from Hamarinn and down below the junction of Árneskvísl branch and the main branch of Þjórsá river. This interpretation is supported by Cobra soundings on the west bank of the river.

Conclusion: The tailrace will mostly be excavated in loose gravel and sand in the river and on the south bank of the Þjórsá river.

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## **Drawings:**



Drawing 16 Powerhouse site, geological section F-F'













**Chan** 

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Hagaey

Karlsnes

Vesturhraun

Austurhraun

Vesturjaðar

Austurjaðar Skarðsheiði

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Bæjarnes

Austurhraun

Bæjarnessporður



Fossnes

Hagi

Vesturjaðar

Austurjabar

Skarðsheiði

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Karlsnes

Vesturhraun



**Hagaey** 

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 $NK-3$ 



#### Fellsmúla 26 S:580 8100 av@almenna.is - - - 108 Reykjavík Fax: 580 8101 www.almenna.is Áritanir á teikningu eru á ábyrgð Almennu verkfræðistofunnar hf. kt. 470671-0179 Design AOT Checked SPS Scale SPS  $Date$  Dec. 2007 No. DRAWING 8 HOLTAVIRKJUN HYDROELECTRIC PROJECT ÞJÓRSÁ LAVA MARGINS AND MAIN FLOW DIRECTION **Legend** • Core hole • Percussion hole V<sub>/</sub> Overflowed by the river Main channel of the lava flow approx. Flood deposits **bjórsá lava Silt -** Project components [16] cale 1:15.000 1227.111-G-008 0 100 200 300 400<br>**mm = mm = m**m Kt. Appr. Map source(s): National Land Survey of Iceland Hnit hf.



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Nautavað

- Fissures & faults mapped by HI 2001
- Active fissures mapped by AV

### Hofsheiði

NK-14NK-13



Vindás

Fissure <sub>St</sub>

**NK-9** 

Samþ. SPS Dags. Dec.  $2007$ <sup>N</sup>

NK-3

- Core hole
- Percussion hole
- Inclined core hole
- June 17, 2001 fissures mapped by HÍ



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Kt. Dec. 2007 1227.111-G-009



Fissures mapped by AV

Map source(s): National Land Survey of Iceland Hnit hf.







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## Núpsvirkjun NK-22 1-2 af 2

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# **KODAK Color Control Patches** ta | Núpsvirkjun NK-23 1-2 af 2  $1190$  $f_{\rm{obs}}$ fund ago ITE PS IPED TOPPED SAS m ilicio mais 142 iriu  $\frac{1}{2}$ **REPORT OF**  $22 - 12$  $24$ **DAS**








































# Holtavirkjun NK-43  $1$  af 1





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# Legend

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Tephra sand and gravel

Þjórsárhraun lava

**Silt/siltstone** 

Hreppar Interglacial tholeiite

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#### Permeability of NK-36



Þjórsár-lava





Olivine basalt

#### Permeability of NK-38







#### Permeability of NK-39



Þjórsá-lava

#### Permeability of NK-40



Þjórsá-lava, sand & gravel and stratified silt



Akbraut interglacial basalt



Akbraut interglacial basalt













#### Permeability of NK-41



Akbraut interglacial basalt





Dykes and olivine basalt



Akbraut interglacial basalt







#### Permeability of NK-42









Permeability of NK-45



## Holtavirkjun hydroelectrical project Temperature readings in boreholes



Akbraut

## Holtavirkjun hydrelectrical project Temperature readings in boreholes



## Holtavirkjun hydroelectrical project Electrical conductivity readings in boreholes



### Holtavirkjun hydrelectrical project Electrical conductivity readings in boreholes



# Temperature and conductivity in boreholes



#### Holtavirkjun hydroelectric powerplant Temperature readings in test pits



Date: 2.11.2006



# **Groundwater level in Holtavirkjun**

# **Point load tests**



#### **NK-22**





\* Tested by Icelandic building research institude

#### **Uniaxial Compressive Strength**





# **Soil and scoria in Þjórsá-lava**





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