

**GEOCHEMICAL SAMPLING, LABORATORY ANALYSIS,
AND MINERAL EQUILIBRIA CALCULATIONS FOR
FLUIDS FROM SELECTED AREAS IN ICELAND**

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ABSTRACT

The work carried out during the project period of the geothermal training programme is described. Field work involved collection of samples of aqueous solutions from well SN-2 in the Seltjarnarnes geothermal area, well H-1 in the Reykjanes Isafjardardjup geothermal area, and from hot springs in the Borgarfjordur and Reykjanes Isafjardardjup geothermal areas. Steam and gas samples were collected from fumaroles G-27 and G-19, and the new well KG-24 in the Krafla geothermal area.

Laboratory work involved analysis of selected samples for pH, CO₂, H₂S, Cl, SiO₂, Al, Fe, and Mn by the methods used at the National Energy Authority of Iceland.

Mineral Equilibria calculations were carried out for selected low and high temperature geothermal fluids using the WATCH1 and WATCH3 computer programs, to obtain the temperatures at which the geothermal fluids were in equilibrium with the various expected alteration minerals. Out of 17 geothermal fluids for which the calculations were performed, 9 gave equilibrium temperatures within 10°C of the "expected" reservoir temperature, and 15 were within 20°C. The two for which the calculated temperatures differed from the expected reservoir temperatures by more than 20°C were high salinity fluids in open reservoirs.

There are however some differences between the results obtained using the WATCH program and those given by Reed and Spycher (1984) using the SOLVEQ program.

It is expected that this approach will give more accurate predictions when better thermodynamic data become available, particularly for aluminium aqueous species, and for clay minerals.

Cold waters which have never been in equilibrium with the surrounding rocks are easily recognised by the complete lack of equilibrium among the minerals.

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1. INTRODUCTION

Geochemistry has uses in both the exploration and exploitation stages of geothermal energy development. During the exploration stages, analysis of the chemical composition of hot springs and fumaroles can be used to infer the subsurface temperatures of the geothermal system, and hence the feasibility of exploitation of the system. The chemistry of the fluids also gives an indication of possible scaling and/or corrosion problems that may arise during exploitation. After exploitation is under way, monitoring of the chemistry of the well discharge fluids with time gives early indications of possible inflow of colder waters into the aquifer which may lead to the cooling of the system.

In order for predictions based on the chemistry of the fluids to be reliable, great diligence has to be exercised during collection and analysis of the geothermal samples. Subsequent interpretation is made easier by the use of computer programs, which allow for the evaluation of the distribution of aqueous species in solution at the high temperatures and pressures existing at depth, taking into account the sampling conditions and the various processes that may take place during the ascent of the fluids from the reservoir to the surface.

During the geothermal training program, the author had experience of applying the methods in use at the National Energy Authority of Iceland for collecting geothermal samples from springs, fumaroles and wells from both high and low temperature geothermal areas in Iceland. Analysis of selected chemical constituents was also carried out for various geothermal fluids. The WATCH computer program of Arnorsson et al (1982) was used to test the applicability of a proposed new approach to geothermometry which takes into account the attainment of overall chemical equilibrium between the geothermal fluid and all the possible mineral phases (Reed and Spycher, 1984).

2. COLLECTION OF GEOTHERMAL FLUID SAMPLES FOR CHEMICAL ANALYSIS

2.1 Introduction

Geothermal fluid samples for chemical analysis consist of aqueous solutions (water samples), steam (NaOH-collected condensate with gas), and gas samples (non-condensable gases, mainly CO₂ and H₂S). In low temperature geothermal areas, samples predominantly consist of aqueous solutions, although gases may also be sampled from hot springs and geothermal wells; while in high temperature geothermal areas, aqueous solutions, steam, and gases are important in wells and hot springs: in fumaroles, there may only be steam and gases that are available for collection.

Proper practice in the collection and preservation of samples is necessary to allow for appropriate interpretation of the resulting analytical data.

2.2 Field Work

During the Geothermal Training Programme practice in the collection of geothermal fluids was undertaken in three low temperature geothermal fields, and one high temperature geothermal field. Well samples in low temperature areas were collected from Seltjarnarnes well 2 and from Reykjanes, Isafjardardjup well 1. Low temperature hot spring samples were collected from England in the Borgarfjordur area, and from Reykjanes Isafjardardjup area. High temperature fluids were collected from both mixed water-steam wells and fumaroles in the Krafla geothermal field. The locations of these areas are shown in figure 1. The procedures used are those employed at the Icelandic National Energy Authority as described by Armannsson, (1985). Four samples were collected from each low temperature site as follows:

1. A raw untreated (Ru) sample portion in a 250ml glass tube for pH and CO₂ determination

2. A filtered untreated (Fu) sample portion in a 500ml plastic bottle for the determination of anions.
3. A filtered acidified (Fa) sample portion in a 100-500ml plastic bottle for the determination of cations.
4. A filtered diluted (Fd) sample portion in a 100ml plastic bottle for the determination of dissolved silica.

In some instances, a 100ml sample was collected, to which was added 1 or 2 pellets of solid NaOH for the determination of silica, CO₂ and H₂S for comparison with the results of the methods above. In most cases, each type of sample was collected in duplicate.

In the high temperature field in Krafla, two fumaroles, G-27 and G-19, as well as the new well KG-24 were sampled. Four sample portions were collected from each of the fumaroles as follows:

1. A steam sample portion (Rb) obtained by bubbling the fumarole discharge into an air-free flask containing 40% NaOH solution.
2. A condensate sample portion (Ruv), collected into a gas tube.
3. A noncondensable gases fraction (Run), collected into a gas tube.
4. A steam condensate sample portion collected into a McCartney bottle for the determination of the isotopic composition of the condensate (Rui).

Both steam and water fractions were sampled from the well. The samples collected from the steam fraction were Rb, Ruv, and Run. From the water fraction were collected the Fu, Fa, Fd, and Fp portions. The latter was a filtered sample, to which was added zinc acetate to precipitate out reduced sulfur species from solution to stop their oxidation to sulfate. The sample was used for the determination of sulfate in solution.

3. CHEMICAL ANALYSIS OF GEOTHERMAL WATER SAMPLES

3.1 Introduction

Complete analyses of sampled geothermal fluids is necessary for the calculation of geothermometer temperatures, and to assess how close the samples are to attaining equilibrium with minerals which may cause scaling problems during the exploitation of the geothermal field. An assessment of potential corrosion problems and an indication of the hydrological conditions of the field can also be obtained through careful interpretation of the analytical data. Methods that are both highly accurate and precise are therefore required for the ensuing interpretation to be valid.

3.2 Laboratory Work

During the Geothermal Training Programme, analyses of selected aqueous solutions for selected species - pH, and concentrations of CO₂, H₂S, Cl, SiO₂, Al, Fe, and Mn - were carried out using the methods employed at the National Energy Authority. The methods of analysis of pH, CO₂, H₂S, Cl, and SiO₂ are described by Armannsson (1985). Analysis of Al was done by determination of the fluorescence of the lumogallion complex (Hydes and Liss, 1976). Iron was determined spectrophotometrically after complexation with 2,4,6 tripyridyl - 1,3,5 triazine (TPTZ) (Grasshoff et al, 1983). Manganese was also determined spectrophotometrically after complexation with formaldoxime (Grasshoff et al, 1983).

The determination of pH, CO₂, and H₂S in the fumarole and well samples was carried out in the Krafla Power Station field laboratory. Analyses were carried out as soon as possible after sampling, not later than the next day.

3.3 Results of Chemical Analysis

The results of the laboratory analyses are summarised in Table 1. The complete results of analyses carried out by the laboratory staff of the National Energy Authority on the same samples are given in Table 2. The methods are summarised in Table 3.

4. MINERAL EQUILIBRIA CALCULATIONS FOR THE ASSESSMENT OF RESERVOIR TEMPERATURES AND/OR MIXING PROCESSES IN SELECTED HIGH AND LOW TEMPERATURE GEOTHERMAL WATERS IN ICELAND

4.1 Introduction

Among the main applications of the study of the chemistry of geothermal waters is the prediction of subsurface temperatures during exploration (i.e. before drilling starts), and the prediction of possible problems associated with scaling and corrosion during the exploitation of the geothermal field. As the reservoir is exploited, an influx of colder waters from the surroundings may lead to a decrease in the temperature of the geothermal system, which can be foretold by comparing changes in the reservoir temperatures indicated by the various chemical geothermometers.

Chemical geothermometers have been developed for use in the prediction of reservoir temperatures under different conditions (Fournier and Rowe, 1966; Fournier and Truesdell, 1973; Fournier, 1979; Fournier and Potter, 1982; Arnorsson, 1975; Arnorsson et al, 1983b). However, the use of these geothermometers has often required prior knowledge of the conditions of the geothermal field, such as subsurface geology. When applied simultaneously, the various geothermometers frequently yield appreciably different values for reservoir temperatures in the same field. It is clearly desirable to have a geothermometer which can give reliable indications of subsurface temperatures during the exploration and exploitation stages of geothermal investigations.

It has been proposed (Reed and Spycher, 1984) that by determining all the mineral phases that can be in equilibrium with a given geothermal fluid at a given temperature, the most probable reservoir temperature can be determined. This is the temperature at which the maximum number of mineral phases are in equilibrium with each other and with the geothermal fluid constituents. By definition, at equilibrium,

$Q_k = K_k$, where Q_k is the activity quotient for the aqueous species arising from the dissolution of the mineral, k , and K_k is the equilibrium constant for the dissolution reaction; ie.

$$K_k = \prod a_i^{\nu_{i,k}},$$

in which a_i is the activity and $\nu_{i,k}$ is the stoichiometric coefficient of species i in the equilibrium mass action expression for mineral k , written for the mineral dissolving (ie. mineral on the left hand side and aqueous components on the right hand side), and a_k is the activity of mineral k . Similarly,

$$Q_k = \prod a_{i,k}^{\nu_{i,k}}.$$

Since at equilibrium $Q_k = K_k$, $\log(Q/K) = 0$ at the equilibrium temperature.

In a plot of $\log(Q/K)$ versus temperature, all minerals in equilibrium with the solution will converge to a value of zero at the same temperature. Geothermal fluids in which there is mixing of solutions of different composition will show either (a) a shift in the positions at which the minerals are apparently in equilibrium to lower $\log(Q/K)$ values if the solutions that mix with the geothermal fluid are very dilute or (b) a complete lack of any identifiable equilibrium temperature if the solution mixing with the geothermal fluid is not dilute, or if the unmixed geothermal fluid is boiling.

Results of chemical analyses of selected geothermal waters from high and low temperature fields in Iceland have been used to calculate possible equilibrium temperatures and/or mixing indications of the fluids using the WATCH mineral equilibria programme of Arnorsson et al, (1982). The results of some of these computations have been compared to the published results of computations for the same samples obtained using the SOLVEQ programme (Reed and Spycher, 1984).

The aim of this work was to investigate the applicability of this concept of "complete mineral equilibria" in the determination of reservoir temperatures using chemical analytical data from both high and low temperature geothermal fields in Iceland.

4.2 Theory of Complete Mineral Equilibria Geothermometry

Reed and Spycher (1984) have proposed that by determining all the mineral phases that are in equilibrium with a given geothermal fluid at a given temperature, the most probable reservoir temperature can be determined. This is the temperature at which the maximum number of mineral phases are in equilibrium with the geothermal fluid, and with each other. As can be shown from considerations of the phase rule, the maximum number of mineral phases which can coexist in equilibrium is about 10 - 12 (see for example Arnorsson et al, 1982). In practice there are usually 5 - 6 minerals that coexist, say in a hand specimen. Since it is not known a priori which phases actually exist in a given geothermal system, the calculations are performed for as many "realistic" minerals as possible: the existing assemblages will be in equilibrium, while others will not be.

In a plot of $\log(Q/K)$ versus temperature, geothermal fluids in which there is mixing of solutions of different compositions will show either (a) a shift in the positions at which the minerals are apparently in equilibrium to negative values of $\log(Q/K)$ if the solutions that mix with the geothermal fluid are very dilute, or (b) a complete lack of any identifiable equilibrium temperature if the solution that is mixing with the geothermal fluid is not dilute.

Further, if the geothermal fluid is boiling prior to sampling, there is an unsystematic dispersion of the $\log(Q/K)$ plots, since the loss of gases (including H_2O vapour) will lead to (a) a simple concentration of elements in the residual aqueous solution (b) a change in the gas content of

the aqueous solution, which may lead to precipitation or dissolution of some mineral phases but not others (c) a change in the pH of the aqueous solution, which may also lead to dissolution or precipitation of some mineral phases but not others. These factors lead to complex changes in the apparent temperatures at which the various minerals are in equilibrium with the geothermal fluid.

Examples of expected plots of $\log (Q/K)$ versus temperature at equilibrium, after mixing, and after boiling are shown in Figures 2, 3, and 4.

4.3 Water Chemistry Data

Most of the data used in these computations was taken from Arnorsson et al (1983a). Additional data were obtained from the samples collected during the training programme, or supplied by Dr. Halldor Armannsson, from analyses carried out in the laboratory of the INEA, Orkustofnun. All the data used are presented in Tables 1 to 3. The sample locations are shown in Figure 1.

The wet-steam well data was interpreted using the WATCH1 programme, while the hot springs and hot water wells data was interpreted using the WATCH3 program of Arnorsson et al (1982).

4.4 Results of WATCH1 and WATCH3 Computations

Examples of outputs from the WATCH1 and WATCH3 programs are given in Tables 5 and 6. The programs are used to calculate the equilibrium constants at a desired temperature, mostly using data from Helgeson (1969) and Helgeson et al (1978) (Arnorsson et al, 1982). Where thermodynamic data are available only at 25°C, equilibria at elevated temperatures have been predicted by the methods of Helgeson (1967).

Examples of mineral equilibria diagrams ($\log(Q/K)$ versus temperature) obtained for the above data are presented in Figures 5 and 6. According to the WATCH program computations, the minerals attain equilibrium over a range of temperatures, rather than at one specific temperature, or a narrow temperature interval as indicated in the figures in Reed and Spycher (1984). The range of equilibrium temperatures, arbitrarily defined as the temperature interval over which at least two minerals attain equilibrium within 4 C° of each other, are presented in table 7. Table 7 also shows the temperature range over which most minerals appear to reach equilibrium (cluster equilibrium temperature), and the optimum equilibrium temperature, defined as the temperature at which the highest number of expected minerals are in equilibrium with the aqueous solution, and with each other. The measured temperature, and the chalcedony and/or quartz equilibrium temperatures, as well as the NaK temperature, are also presented for comparison.

The calculated equilibrium temperature for Reykholt well 1 is in excellent agreement with the measured downhole temperature. For this sample, the calculated equilibrium temperature of 135°C is only 2 C° off the measured downhole temperature of 133°C. The differences between calculated equilibria temperatures and measured downhole temperatures then vary from 7 to 8 C° in Urridavatn well 8 and Krafla well 9 to 46 C° for Reykjanes well 8. The calculated temperatures can be too low, as at the Svartsengi well 4, Reykjanes well 8, Hveragerdi well 4, or Urridavatn well 8 areas, but they are more typically higher than the measured temperatures. Two of the samples used by Reed and Spycher (1984) to support their proposed approach to geothermometry - Reykjanes well 8 and Hveragerdi well 4 - give calculated equilibrium temperatures with the WATCH program, that are 46 and 15°C too low, respectively, compared to the measured downhole temperatures.

The third, Namafjall well 8 gives a calculated temperature that is in very good agreement with the estimated reservoir

temperature at the time of sampling. Figures 7, 8, and 9 show comparisons between the plots given in Reed and Spycher (1984) from calculations with the SOLVEQ program, and those obtained in this study using the WATCH1 program.

4.5 Discussion

Measured temperatures do not necessarily show the reservoir temperatures within 1 or 2 degrees. There may be several feeding aquifers at different temperatures within a well. Also, temperature logging in wells is usually carried out shortly after drilling, when the effects of disturbances due to drilling might not have worn off, and may still be having an effect on the temperature distribution in the well. Temperature logs made long after drilling also indicate hybrid quasi-steady state temperatures, in cases where there are several aquifers feeding the well. The question therefore arises as to what should constitute an acceptable estimate of the temperature of the reservoir, since in the case of a multiple feed aquifers of different temperatures, such a temperature does not exist: it would be asking too much to expect a geothermometer to give the temperatures of all the individual aquifers. In the exploration stages, the aim of estimating reservoir temperatures is to give an indication as to what uses a particular geothermal field should be put to, e.g. whether it can be used for the generation of electricity, or not. In such instances, and considering the uncertainties mentioned above, a temperature estimate that is within 10 to 20°C may be acceptable.

The calculated equilibrium temperatures for Svartsengi and Reykjanes are unreasonably low. Both of these areas have very open systems, and fluid flow is probably fast. They are saline (sea water), and it is conceivable that the fast inflow of sea water does not allow for attainment of chemical equilibrium, even though the waters get heated to high temperatures. The WATCH program may also not adequately account for the thermodynamics of aqueous species at these high temperatures and high salinities.

The measured temperature at Reykjaból (152°C) may be slightly lower than the true reservoir temperatures as it was measured soon after drilling (Benedikt Steingrímsson and Halldor

Armannsson, personal communication), and the calculated optimum equilibrium temperature of 161°C is very reasonable. It is conceivable that a trickle of cooler aquifer mixes with the warmer aquifer and cools it: this cooler flow may be responsible for the value of 126°C obtained.

The aquifer in well 9 at Krafla was originally at 300°C, but later cooled down. The temperature of 240°C was a compromise based on several measurements: it is conceivable that the hotter water did not lose all its imprints. The calculated value of 248°C therefore seems to be very good (Halldor Armannsson, personal communication).

The temperature data presented for well 7 at Krafla by Arnorsson et al (1983a) were measured soon after drilling. There is, for example, no longer any flow at 160°C, and the deepest aquifer has been completely blocked off (no effect on the flow). On the other hand, the equilibrium temperature at 1700m depths is approximately 325°C. It has been estimated that in 1979, when the sample was collected, the flow from the upper zone (206°C) contributed about 40% to the total flow, while the flow from the lower zone (325°C) contributed approximately 60% to the total flow (Halldor Armannsson, personal communication). Thus a temperature of approximately 277°C would be expected from the mixture, and the calculated temperature of 275°C is in excellent accord with this.

During the magmatic activities at Krafla in April and September of 1977, there was a southward flow of magma which affected the Namafjall field severely, where wells 4 and 8 are situated: well 4 actually erupted some magma. The area was extensively fractured, and a large amount of cold groundwater entered the geothermal system. It was difficult to obtain temperature data from these wells, but one profile obtained for well 4 in 1984 shows an inflow temperature of 180°C (Halldor Armannsson and Benedikt Steingrimsson, personal communication). The samples referred to in Arnorsson et al (1983a) were collected in 1979, but the temperature

data listed were obtained before 1977 (Halldor Armannsson, personal communication). The calculated equilibrium temperature value of 185°C for well 4 is therefore very reasonable, and the temperature of 160°C for well 8 (which was always cooler than well 4 (Halldor Armannsson, personal communication)) is probably close to the reservoir temperature at the time the samples were collected.

The Seltjarnarnes well 2 draws from relatively shallow depths, and conventional geothermometers suggest temperatures higher than those measured, with the waters probably coming from a hotter, deeper aquifer. Well 4, which is adjacent to well 2, draws from a deeper aquifer at 120°C, which is the primary reservoir water. There are indications that the two wells are hydrologically connected (Tulinus et al, 1987). The calculated values of 134°C and 138°C are therefore within 20°C of the deep reservoir water; the calculated 94°C temperature probably reflects water that has mixed with cooler, shallow water.

The temperature of 82°C calculated for England, Borgarfjordur, is a little lower than the measured temperature of 91°C. This can, however be taken to be a strong indication that temperatures substantially higher than the measured surface temperatures are unlikely to exist in the reservoir. The same argument probably applies for the calculated equilibrium temperature of 90°C for the Reykjanes, Isafjardardjup spring: the highest temperatures obtained in borehole 2 adjacent to (15m away from) the spring that was sampled, was 94.5°C at a depth of 240m, after which it decreased to 92°C.

The calculated equilibrium temperature for Urridavatn well 8 is 84°C, compared to a measured temperature of 77°C. This is considered to be reasonable enough, considering the uncertainties mentioned earlier. The sample from well 4 was collected from an artesian flow at a temperature of 60.2°C. This probably contained a considerable amount of a cooler

component from the lake (Benjaminsson and Gislason, 1986). When the well is pumped, the temperature of the flow increases to 64-65°C, which is probably close to the temperature of the main aquifer. The similarity between the range and cluster equilibrium temperatures for well 4 and 8 at Urridavatn suggests that the two wells probably draw fluids from the same aquifers, but in different proportions.

In the IRDP well, the main flow is 47°C, but a deeper one at 78°C has also been observed. The calculated value of 65°C is somewhat high, since the contribution from the lower aquifer is not likely to be large enough to cause such a large difference in temperature.

As expected, the cold spring at Borgarfjordur did not indicate any minerals to be in equilibrium with one another. If the 17 other samples are considered in light of the information discussed above, it is seen that there are 9 samples for which the optimum equilibrium temperature is within 10°C of the expected reservoir temperature, and 15 samples for which the optimum temperature is within 20°C of the expected reservoir temperature. Two samples out of 17 have calculated equilibrium temperatures differing from the expected reservoir temperatures by more than 20°C. These are the high temperature, highly saline waters in open reservoirs for which either equilibrium is not attained in the reservoir, or the WATCH program calculations do not adequately model.

The computations assume an equilibrium situation. Such equilibrium is only established where waters have been at the high reservoir temperatures for sufficiently long periods of time for water-rock reactions to have attained equilibrium at one temperature. This condition is violated in cases where waters come from different aquifers at different temperatures. This is the case with the discharges from Reykjanes well 8, in which aquifer temperatures range from 274° to 292°C. Equilibrium might also not be attained in

flowing systems, when there is relatively rapid movement of waters through aquifers as probably happens during exploitation of geothermal systems. Further, during ascent of geothermal fluids, some minerals may be precipitated from solution (as is observed at Krafla, Svartsengi and Hveragerdi, where calcite scaling frequently causes blocking of wells), while other minerals may be dissolved from the wall rocks into solution.

There are considerable uncertainties in the available thermodynamic data, especially for aqueous species of aluminium. This means that calculated equilibrium constants may not be very accurate, particularly at temperatures above 200°C (Arnorsson et al, 1982). Thermodynamic data for some solid solution minerals are currently not available. According to Reed and Spycher (1984) this introduces only small errors if end-member data are used instead of the actual solid solutions. However, for minerals with a very low slope in $\log(Q/K)$ versus temperature curve, small shifts in the value of $\log(Q/K)$ can result in large shifts in the intersection temperatures.

The WATCH program recalculates the deep water composition by assuming adiabatic steam loss. This assumption is probably true for low temperature systems in which boiling starts near the surface, but it may be less valid in the case of high temperature systems where there may be a continuous steam loss as the solution rises to the surface.

The mineral phases that are taken into consideration in the WATCH program calculations are those that are commonly observed in drill holes in Icelandic geothermal systems. As mentioned above, solid solutions and interlayering of clay minerals has not been taken into consideration. For accurate simulation of the geothermal system, the real conditions at depth must be portrayed as closely as possible, and these minerals need to be taken into account.

The quality of the chemical analyses is assumed to be very high. However, inaccuracies in the analyses, particularly for the gases could cause a wide scatter in the calculated equilibrium temperatures.

5. SUMMARY AND CONCLUSIONS

During the geothermal training program period, experience in collecting samples from hot springs, fumaroles, and wells from both high temperature and low temperature geothermal areas in Iceland following the methods used at the National Energy Authority of Iceland was gained. Analysis of pH, CO₂, H₂S, Cl, SiO₂, Al, Fe, and Mn was also carried out on selected geothermal samples.

Calculations were made to verify the new approach to geothermometry proposed by Reed and Spycher (1984), in which the total equilibrium of the water-rock system is examined, using computations with the WATCH programme of Arnorsson et al (1982). These computations showed that the approach can be used to give an indication of the reservoir temperature to within 20°C, particularly for relatively dilute waters. Calculated equilibrium temperatures for high temperature, high salinity fluids vary from the expected reservoir temperatures by more than 20°C. Thermodynamic data of better quality, and a better understanding of the processes which take place during the generation and subsequent ascent of the geothermal fluid are needed before this approach to geothermometry can be applied with confidence to all types of geothermal fluids. Waters in which equilibrium is not approached at all (as for example river waters) can be distinguished by a complete lack of equilibrium.

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Table 1. Results of chemical analyses of selected aqueous geothermal samples.

Spl No.	Locality	T/°C	pH/T°C	CO2 ppm	H2S ppm	Cl ppm	SiO2 ppm	Al ppm	Fe ppm	Mn ppm
880072	Seltjarnarnes well 2, 80m	31	8.85/18.5	23	0.02	1417	8	0.164	0.445	<0.00002
880073	" " " , 570m	78	8.45/20.8	11	0.01	1527	48	0.137	0.066	0.00002
880074	" " " , 725m	77	8.18/21.2	9	0.01	1395	147	0.278	0.237	<0.00002
880075	Borgarfjordur, cold spring	3.8	8.10/20.3	25	0.02	8	18	0.048	0.010	0.011
880076	England, hot spring	91.1	9.24/21.3	51	0.24	19	167	0.031	0.014	0.009
889015	Thorlakshofn, groundwater mixed with 1/2 sea water	27	8.09/22.0	16.1	<0.03	10066	23.4	0.022	0.131	0.402
889034	Reykjavik, cold groundwater	16	8.05/23.3	177	<0.03	387	21.4	0.095	0.069	0.122
889035	Reykjavik, cold lake	19	9.44/23.4	88.7	<0.03	418	18.2	0.058	0.131	0.108
880079	Urridavatn well 8	77	9.82/23.6	10.8	0.06	46.3	66.3	0.138	0.007	0.004
880080	Urridavatn well 4	60	9.81/23.7	21.0	0.07	29.0	54.4	0.106	0.023	0.016
880081	IRDP well, Areyjar	43	9.64/23.8	5.6	<0.03	70.5	59.3	0.036	0.002	0.005
881003	G-27 fumarole, Krafla steam fraction	93	nd	40860	792	nd	nd	nd	nd	nd
	condensate fraction	93	4.70/19.2	936	15.4	1.6	nd	nd	nd	nd
881004	G-19 fumarole, Krafla steam fraction	94.8	nd	8427	234	nd	nd	nd	nd	nd
	condensate fraction	94.8	4.50/21.6	630	44.6	0.4	nd	nd	nd	nd
881005	KG-24 well, Krafla steam fraction	130	nd	1251	123	nd	nd	nd	nd	nd
	condensate fraction	130	4.65/22.2	415	88.9	nd	nd	nd	nd	nd
	water fraction	130	9.31/20.5	33.8	12.3	nd	298	nd	nd	nd

Note: "nd" denotes "not determined"

Table 2. Results of complete chemical analyses of some of the samples in Table 1. Results for sample 60 were taken from Arnorsson et al (1983a)

Spl. No.	Temp. /°C	pH/T°C	Resist. Ωm/°C	SiO2 ppm	Na ppm	K ppm	Ca ppm	Mg ppm	CO2 ppm	SO4 ppm	H2S ppm	Cl ppm	F ppm	Al ppm	Fe ppm	Mn ppm	TDS ppm
880072	nd	8.85/19	3990/21.2	41.3	568.9	7.3	277	0.27	23	192.4	<0.01	1216	0.473	0.164	0.445	<0.00002	2327.4
880073	nd	8.45/21	4270/21.1	92.3	595.6	14.5	324.7	0.28	11	221.7	<0.01	1321	0.549	0.137	0.066	0.00002	2582.3
880074	nd	8.18/21	3850/21.1	85.9	537.2	10.1	285.4	0.22	9	203.2	<0.01	1167	0.527	0.278	0.237	<0.00002	2298.7
880075	3.8	8.10/20	90.0/20.9	15.8	9.1	0.34	5.87	2.12	25	4.8	<0.01	8.4	0.081	0.048	0.010	0.011	71.4
880076	91.1	9.27/20	344/20.8	164.6	73.8	2.5	2.30	0.04	51	55.6	0.24	18.0	1.113	0.031	0.014	0.009	368.9
880079	77.2	9.82/23.6	367/21.5	66.3	70.7	1.22	7.29	0	10.8	55.2	0.06	46.3	0.697	0.138	0.007	0.004	258.6
880080	60.2	9.81/23.7	268/21.3	54.4	53.6	0.78	4.15	0.009	21.0	28.9	0.07	29.0	0.467	0.106	0.023	0.016	163.1
880081	42.8	9.64/23.8	598/21.3	59.3	94.9	1.21	31.9	0	5.6	138.1	<0.03	70.5	2.056	0.036	0.002	0.0005	403.6
60	84	9.07/11	nd	70.9	184.2	4.34	67.6	0.094	9.0	42.1	<0.01	390.2	0.23	0.02	0.017	nd	965

Table 3. Summary of analytical methods

pH	pH-meter
ΣCO_2	Titration with 0.1N HCl using pH-meter
$\Sigma\text{H}_2\text{S}$	Titration with $\text{Hg}(\text{CH}_3\text{COO})_2$ using dithizone as indicator
Cl	Mohr titration
SiO_2	Spectrophotometric on yellow silicomolybdate complex
Al	Fluorimetric on pink lumogallion complex
Fe	Spectrophotometric on pink TPTZ complex
Mn	Spectrophotometric on red-brown formaldoxime complex

Table 4. Chemical composition of water and steam from selected wet-steam wells in Iceland (from Arnorsson et al 1983a).

Concentrations are in ppm. Sample numbers correspond to locality numbers shown in figure 1.

	Svartsengi	Reykjanes	Reykholt	Reykjabol	Hveragerdi	Krafla	Krafla	Namafjall	Namafjall
	well 4	well 8	well 1	well 1	well 4	well 9	well 7	well 4	well 8
Sample No.	08	09	28	30	32	48	49	51	52
Water sample (ppm)									
pH/°C	7.53/20	6.38/20	9.30/20	9.28/20	8.82/20	9.02/22	8.52/20	9.33/22	8.20/22
SiO ₂	534.6	631.1	179.8	252.8	281.0	575.5	766.8	442.8	446.3
B	8.65	8.72	0.35	0.65	0.62	0.71	1.50	1.48	1.66
Na	8037	11150	109.7	102.2	153.3	193.9	193.0	158.6	154.8
K	1245	1720	5.54	6.21	13.4	29.0	36.0	25.5	24.0
Ca	1343	1705	2.68	1.28	1.73	2.54	1.90	2.81	4.52
Mg	1.68	1.44	0.055	0.048	0.002	0.064	0.067	0.046	0.085
Fe	0.196	0.329	0.010	0.007	0.008	0.020	8.78	0.007	0.019
Al	0.07	0.07	0.17	0.42	0.14	0.12	0.21	0.10	0.10
ΣCO ₂	32.6	63.1	30.4	54.3	74.2	100.0	186.6	70.7	88.2
SO ₄	40.5	28.4	72.6	70.7	43.7	194.5	143.0	42.2	48.7
ΣH ₂ S	0.16	2.21	1.04	4.09	19.2	36.2	22.2	109.2	132.6
Cl	17010	22835	77.7	41.6	109.5	28.0	52.0	16.3	16.6
F	0.16	0.21	2.69	1.86	1.82	0.87	1.03	0.68	0.43
TDS	27470	39124	555	569	765	1237	1389	1003	902
Condensate (ppm)									
CO ₂	722	584	123	242	627	553	14680	190	172
H ₂ S	18.8	65.6	7.80	10.3	84.5	108	381	258	277
Gas w/condensate (vol%)									
CO ₂	97.7	96.2	18.0	33.2	84.5	88.6	-	27.2	36.8
H ₂ S	0.8	2.9	0.4	0.6	3.0	5.0	-	10.4	17.0
H ₂	0.1	0.2	0.4	0.4	2.8	3.4	2.5	48.6	37.4
CH ₄	0.1	0.1	0.3	0.3	0.4	0.1	0.1	2.3	2.9
N ₂	1.3	0.6	80.9	65.5	9.4	2.6	0.7	11.5	5.9
l gas/kg condensate/°C	0.67/20	2.63/20	1.92/20	3.85/20	1.06/20	2.08/20	8.50/20	0.86/20	6.25/20
sampling pressure/bars	3.8	20.0	2.0	4.0	6.8	7.9	5.5	8.6	9.8
Disch. Enth. MJ/Kg	1.028	1.151	0.559	0.628	0.782	0.962	1.602	1.117	1.092
Mass flow kg/s.	40	50	8	10	80	40	12	40	45

Table 5. Physical data on drillholes (modified from Arnorsson et al, 1983a, unless otherwise indicated).

Sample No.	Location	Depth (meters)	Discharge (kg/sec)	Fluid inflow depth/temp. (meters) (°C)
08	Svartsengi, 4	1703	40	1090/243, 1180, 1255, 1365, 1565/242.
09	Reykjanes, 8	1756	50	1010/274, 1155/279, 1235, 1280/280 1510/284, 1650, 1660/288, 1705, 1740/292.
28	Reykholt, 1	752	8	745/133.
30	Reykjaból, 1	820	10	773/152.
32	Hveragerdi, 4	692	80	630, 660, 670/181.
48	Krafla, 9	1263	40	1226/240.
49	Krafla, 7	2165	8	825, 1163/206, 1700/325.
51	Namafjall, 4	1136	40	640, 1040/180.
52	Namafjall, 8	1283	45	850, 1000/160.
72}	Seltjarnarnes, 2	856	5	80/31, 580/78, 760/77 a.
73}				
74}				
79	Urridavatn, 8	1066		700, 870/77 b.
80	Urridavatn, 4	1136		200, 450/64 c.
81	IRD	1919		517-628/47(main), 1710/78 d.

a Tomasson et al, 1977

b Benjaminsson, 1984

c Benjaminsson et al, 1982

d Benedikt Steingrímsson, pers. comm.

Table 6. Example of a printout from the WATCH1 program

UNU Geothermal Training Programme

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08 Svartsengi Well 4

PROGRAM WATCH1.

WATER SAMPLE (PPM)		STEAM SAMPLE				
PH/DEG.C	7.53/20.0	GAS (VOL.%)		REFERENCE TEMP.	DEGREES C	242.0 (ARBITRARY)
SI02	534.60	CO2	97.70			
NA	8037.00	H2S	.80	SAMPLING PRESSURE	BARS ABS.	4.8
K	1245.00	H2	.10	DISCHARGE ENTHALPY	MJ/OL/RG	1.046 (CALCULATED)
CA	1343.00	O2	.01	DISCHARGE	KG/SEC.	40.0
MG	1.620	CH4	.10			
CO2	32.60	N2	1.30	MEASURED TEMPERATURE	DEGREES C	242.0
SO4	40.50			RESISTIVITY/TEMP.	OHMM/DEG.C	.0/ .0
H2S	.16			BH/TEMP.	MV/DEG.C	.000/ .0
CL	17010.00					
F	.16	LITERS GAS PER KG				
DISS.SOLIDS	27470.00	CONDENSATE/DEG.C	.67/20.0	MEASURED DOWNHOLE TEMP.	DEGREES C/METERS	FLUID INFLOW DEPTH (METERS)
AL	.0700					
B	8.6500	CONDENSATE (PPM)		1090.0	243.0	1090.0
FB	.1960	PH/DEG.C	.00/ .0	1180.0	242.0	1180.0
NH3	.0000	CO2	722.00	1255.0	242.0	1255.0
		H2S	18.80	1365.0	242.0	1365.0
		NA	.00	1565.0	242.0	1565.0
				.0	.0	.0
				.0	.0	.0
				.0	.0	.0
		CONDENSATE WITH NAOH (PPM)		.0	.0	.0
		CO2	722.00	.0	.0	.0
		H2S	18.80	.0	.0	.0

IONIC STRENGTH = .49691 IONIC BALANCE : CATIONS (MOL.EQ.) .44664270
 ANIONS (MOL.EQ.) .47954240
 DIFFERENCE (%) -7.10

DEEP WATER (PPM)		DEEP STEAM (PPM)		GAS PRESSURES (BARS ABS.)	
SI02	430.13	CO2	167.36	CO2	.387E-02
NA	6465.97	H2S	3.80	H2S	.384E-04
K	1001.58	H2	.01	H2	.304E-04
CA	1080.48	O2	.02	O2	.260E-05
MG	1.303	CH4	.09	CH4	.445E-04
SO4	32.58	N2	1.98	N2	.470E-03
CL	13683.81	NH3	.00	NH3	.000E+00
F	.13			H2O	.347E+02
DISS.S.	22100.30			TOTAL	.347E+02
AL	.0563				
B	6.9585			H2O (%)	.00
FB	.1577			BOILING PORTION	.00

ACTIVITY COEFFICIENTS IN DEEP WATER

H+	.630	KSO4-	.538	FE++	.106	FECL+	.487
OH-	.466	F-	.466	FE+++	.020	AL+++	.020
H3SIO4-	.487	CL-	.444	FEOH+	.524	ALOH++	.092
H2SIO4--	.092	NA+	.487	FE(OH)3-	.524	AL(OH)2+	.538
H2BO3-	.419	K+	.444	FE(OH)4--	.083	AL(OH)4-	.507
HCO3-	.487	CA++	.106	FEOH++	.083	ALSO4+	.507
CO3--	.073	MG++	.151	FE(OH)2+	.538	AL(SO4)2-	.507
HS-	.466	CAHCO3+	.557	FE(OH)4-	.538	ALF++	.092
S--	.083	MGHCO3+	.487	FEHO4+	.524	ALF2+	.538
HSO4-	.507	CAOH+	.557	FECL++	.083	ALF4-	.507
SO4--	.062	MGOH+	.571	FECL2+	.524	ALF5--	.073
NASO4-	.538	NH4+	.419	FECL4-	.487	ALF6---	.003

CHEMICAL COMPONENTS IN DEEP WATER (PPM AND LOG MOLE)

H+ (ACT.)	.00	-5.751	MG++	1.27	-4.282	FE(OH)3	.15	-5.860
OH-	.17	-4.990	NACL	1555.70	-1.575	FE(OH)4-	.18	-5.847
H4SIO4	686.50	-2.146	KCL	115.35	-2.810	FECL+	.00	-7.833
H3SIO4-	.74	-5.108	NASO4-	6.12	-4.289	FECL2	.00	-11.008
H2SIO4--	.00	-10.012	KSO4-	4.12	-4.516	FECL++	.00	-17.536
NAH3SIO4	.98	-5.083	CASO4	17.21	-3.898	FECL2+	.00	-18.260
H3BO3	39.76	-3.192	MGSO4	.09	-6.113	FECL3	.00	-19.421
H2BO3-	.04	-6.134	CACO3	.09	-6.056	FECL4-	.00	-20.521
H2CO3	203.56	-2.484	MGCO3	.00	-9.862	FEHO4	.00	-11.345
HCO3-	6.31	-3.985	CAHCO3+	42.12	-3.380	FEHO4+	.00	-20.085
CO3--	.00	-8.517	MGHCO3+	.01	-7.085	AL+++	.00	-19.951
H2S	3.60	-3.976	CAOH+	.41	-5.144	ALOH++	.00	-13.468
HS-	.20	-5.221	MGOH+	.02	-6.322	AL(OH)2+	.00	-8.014
S--	.00	-13.710	NH4OH	.00	.000	AL(OH)3	.16	-5.683
H2SO4	.00	-12.210	NH4+	.00	.000	AL(OH)4-	.00	-8.742
HSO4-	.37	-5.417	FE++	.00	-8.402	ALSO4+	.00	-21.370
SO4--	12.13	-3.899	FE+++	.00	-22.150	AL(SO4)2-	.00	-23.642
HF	.02	-6.004	FEOH+	.00	-8.650	ALF++	.00	-16.642
F-	.11	-5.238	FE(OH)2	.00	-9.666	ALF2+	.00	-15.255
CL-	12685.27	-.446	FE(OH)3-	.00	-10.958	ALF3	.00	-15.765
NA+	5852.59	-.594	FE(OH)4--	.00	-16.613	ALF4-	.00	-17.952
K+	939.90	-1.619	FE(OH)++	.00	-14.724	ALF5--	.00	-21.119
CA++	1058.39	-1.578	FE(OH)2+	.00	-8.890	ALF6---	.00	-25.426

IONIC STRENGTH = .37167 IONIC BALANCE : CATIONS (MOL.BQ.) .33195230
 ANIONS (MOL.BQ.) .35822360
 DIFFERENCE (%) -7.61

CHEMICAL GEOTHERMOMETERS DEGREES C 1000/T DEGREES KELVIN = 1.94

QUARTZ 233.9
 CHALCEDONY 999.9
 NAE 244.1

OXIDATION POTENTIAL (VOLTS) : EH H2S= -.444 EH CH4= -.510 EH H2= -.459 EH NH3= 99.999

LOG SOLUBILITY PRODUCTS OF MINERALS IN DEEP WATER

	TEOR.	CALC.		TEOR.	CALC.		TEOR.	CALC.
ADULARIA	-14.433	-14.386	ALBITE LOW	-13.978	-13.321	ANALCIME	-11.524	-11.174
ANHYDRITE	-7.959	-7.660	CALCITE	-12.565	-12.212	CHALCEDONY	-2.027	-2.146
MG-CHLORITE	-84.280	-86.475	FLUORITE	-10.915	-13.692	GORHITE	1.834	-.795
LAUMONTITE	-24.532	-23.091	MICROCLINE	-15.126	-14.386	MAGNETITE	-17.696	-21.612
CA-MONTMOR.	-72.625	-69.571	K-MONTMOR.	-34.035	-35.480	MG-MONTMOR.	-74.122	-72.118
NA-MONTMOR.	-34.308	-34.415	MUSCOVITE	-17.853	-15.694	PRHNNITE	-37.070	-34.143
PYRRHOTITE	-25.519	-47.686	PYRITE	-47.057	-60.469	QUARTZ	-2.140	-2.146
WAIKAITITE	-24.408	-23.091	WOLLASTONITE	7.619	6.800	ZOISITE	-37.611	-34.797
EPIDOTE	-37.233	-34.938	MARCASITE	-30.030	-60.469			

Table 6. (continued)

UNU Geothermal Training Programme
SAMPLE = 08

DEEP WATER COOLED TO 250.0 DEGREES C.

ACTIVITY COEFFICIENTS IN DEEP WATER

H+	.625	KS04-	.530	FE++	.099	FECL+	.479
OH-	.458	F-	.458	FE+++	.018	AL+++	.018
H3SIO4-	.479	CL-	.435	PROH+	.517	ALOH++	.086
H2SIO4--	.086	NA+	.479	FE(OH)3-	.517	AL(OH)2+	.530
H2BO3-	.410	K+	.435	FE(OH)4--	.078	AL(OH)4-	.499
HCO3-	.479	CA++	.099	PROH++	.078	ALSO4+	.499
CO3--	.068	MG++	.144	FE(OH)2+	.530	AL(SO4)2-	.499
HS-	.458	CAHCO3+	.549	FE(OH)4-	.530	ALF++	.086
S--	.078	MGHCO3+	.479	FESO4+	.517	ALF2+	.530
HSO4-	.499	CAOH+	.549	FECL++	.078	ALF4-	.499
SO4--	.058	MGOH+	.564	FECL2+	.517	ALF5--	.068
NASO4-	.530	NH4+	.410	FECL4-	.479	ALF6---	.002

CHEMICAL COMPONENTS IN DEEP WATER (PPM AND LOG MOLE)

H+ (ACT.)	.00	-5.804	MG++	1.27	-4.283	FE(OH)3	.13	-5.911
OH-	.21	-4.908	NACL	1758.62	-1.522	FE(OH)4-	.20	-5.799
H4SIO4	686.48	-2.146	KCL	128.75	-2.763	FECL+	.00	-8.443
H3SIO4-	.76	-5.100	NASO4-	6.08	-4.292	FECL2	.00	-11.298
H2SIO4--	.00	-9.991	KS04-	4.35	-4.493	FECL++	.00	-18.041
NAH3SIO4	.98	-5.081	CASO4	17.49	-3.891	FECL2+	.00	-18.788
H3BO3	39.75	-3.192	MGSO4	.09	-6.137	FECL3	.00	-19.939
H2BO3-	.05	-6.092	CACO3	.10	-6.010	FECL4-	.00	-21.009
H2CO3	203.68	-2.484	MGCO3	.00	-9.857	FESO4	.00	-11.996
HCO3-	5.89	-4.015	CAHCO3+	42.62	-3.375	FESO4+	.00	-20.630
CO3--	.00	-8.556	MGHCO3+	.01	-7.092	AL+++	.00	-20.557
H2S	3.61	-3.975	CAOH+	.53	-5.031	ALOH++	.00	-13.820
HS-	.19	-5.245	MGOH+	.03	-6.151	AL(OH)2+	.00	-8.137
S--	.00	-13.616	NH4OH	.00	.000	AL(OH)3	.16	-5.682
H2SO4	.00	-12.186	NH4+	.00	.000	AL(OH)4-	.00	-8.867
HSO4-	.40	-5.382	FE++	.00	-9.035	ALSO4+	.00	-21.956
SO4--	11.78	-3.911	FE+++	.00	-22.818	AL(SO4)2-	.00	-24.222
HF	.02	-5.981	PROH+	.00	-9.162	ALF++	.00	-17.123
F-	.11	-5.242	FE(OH)2	.00	-10.016	ALF2+	.00	-15.643
CL-	12555.81	-.451	FE(OH)3-	.00	-11.106	ALF3	.00	-16.089
NA+	5772.77	-.600	FE(OH)4--	.00	-16.696	ALF4-	.00	-18.235
K+	932.80	-1.622	FE(OH)++	.00	-15.172	ALF5--	.00	-21.371
CA++	1058.02	-1.578	FE(OH)2+	.00	-9.109	ALF6---	.00	-25.671

IONIC STRENGTH = .36800 IONIC BALANCE : CATIONS (MOL.EQ.) .32828750
ANIONS (MOL.EQ.) .35455960
DIFFERENCE (%) -7.69

OXIDATION POTENTIAL (VOLTS) : BH H2S= -.462 BH CH4= -.528 BH H2= -.470 BH NH3= 99.999

LOG SOLUBILITY PRODUCTS OF MINERALS IN DEEP WATER

	TEOR.	CALC.		TEOR.	CALC.		TEOR.	CALC.
ADULARIA	-14.402	-14.405	ALBITE LOW	-13.956	-13.341	ANALCINE	-11.528	-11.195
ANHYDRITE	-8.106	-7.732	CALCITE	-12.771	-12.308	CHALCEDONY	-1.995	-2.146
MG-CHLORITE	-84.886	-86.008	FLUORITE	-10.957	-13.743	GORTHITE	2.273	-.827
LAUMONTITE	-24.572	-23.131	MICROCLINE	-15.073	-14.405	MAGNETITE	-16.930	-22.185
CA-MONTMOR.	-72.599	-70.585	K-MONTMOR.	-33.974	-35.986	MG-MONTMOR.	-74.100	-73.129
NA-MONTMOR.	-34.252	-34.922	MUSCOVITE	-17.843	-15.875	PREHNITE	-37.322	-34.060
PYRRHOTITE	-21.030	-47.731	PYRITE	-41.351	-60.895	QUARTZ	-2.105	-2.146
WAIKAITI	-24.549	-23.131	WOLLASTONITE	7.484	6.880	ZOISITE	-37.954	-34.795
EPIDOTE	-37.493	-34.887	MARCASITE	-24.586	-60.895			

Table 6. (continued)

UNU Geothermal Training Programme
 SAMPLE = 08

DEEP WATER COOLED TO 200.0 DEGREES C.

ACTIVITY COEFFICIENTS IN DEEP WATER

H+	.658	HSO4-	.573	FE++	.137	FECL+	.527
OH-	.507	F-	.507	FE+++	.032	AL+++	.032
H3SIO4-	.527	CL-	.486	FEOH+	.561	ALOH++	.122
H2SIO4--	.122	NA+	.527	FE(OH)3-	.561	AL(OH)2+	.573
H2BO3-	.463	K+	.486	FE(OH)4--	.112	AL(OH)4-	.545
HCO3-	.527	CA++	.137	FEOH++	.112	ALSO4+	.545
CO3--	.099	MG++	.187	FE(OH)2+	.573	AL(SO4)2-	.545
HS-	.507	CAHCO3+	.591	FE(OH)4-	.573	ALF++	.122
S--	.112	MGHCO3+	.527	FEHO4+	.561	ALF2+	.573
HSO4-	.545	CAOH+	.591	FECL++	.112	ALF4-	.545
SO4--	.087	MGOH+	.604	FECL2+	.561	ALF5--	.099
NASO4-	.573	NH4+	.463	FECL4-	.527	ALF6---	.006

CHEMICAL COMPONENTS IN DEEP WATER (PPM AND LOG MOLE)

H+ (ACT.)	.00	-5.521	MG++	1.28	-4.279	FE(OH)3	.03	-6.613
OH-	.06	-5.439	NACL	1031.32	-1.753	FE(OH)4-	.01	-7.106
H4SIO4	686.69	-2.146	KCL	66.77	-3.048	FECL+	.16	-5.767
H3SIO4-	.64	-5.173	NASO4-	6.00	-4.297	FECL2	.00	-10.817
H2SIO4--	.00	-10.127	HSO4-	3.32	-4.610	FECL++	.00	-16.022
NAH3SIO4	.87	-5.132	CASO4	14.87	-3.962	FECL2+	.00	-16.664
H3BO3	39.77	-3.192	MGSO4	.11	-6.048	FECL3	.00	-17.896
H2BO3-	.03	-6.314	CACO3	.05	-6.269	FECL4-	.00	-19.178
H2CO3	203.11	-2.485	MGCO3	.00	-9.867	FEHO4	.00	-9.022
HCO3-	9.46	-3.810	CAHCO3+	37.69	-3.429	FEHO4+	.00	-18.366
CO3--	.00	-8.265	MGHCO3+	.01	-7.039	AL+++	.00	-16.873
H2S	3.54	-3.984	CAOH+	.10	-5.763	ALOH++	.00	-11.688
HS-	.26	-5.110	MGOH+	.00	-7.180	AL(OH)2+	.00	-7.404
S--	.00	-14.175	NH4OH	.00	.000	AL(OH)3	.16	-5.690
H2SO4	.00	-12.465	NH4+	.00	.000	AL(OH)4-	.00	-8.147
HSO4-	.21	-5.655	FE++	.04	-6.154	ALSO4+	.00	-18.424
SO4--	14.59	-3.818	FE+++	.00	-19.784	AL(SO4)2-	.00	-20.724
HF	.01	-6.172	FEOH+	.01	-7.059	ALF++	.00	-14.184
F-	.12	-5.215	FE(OH)2	.00	-8.924	ALF2+	.00	-13.261
CL-	13026.40	-.435	FE(OH)3-	.00	-11.274	ALF3	.00	-14.078
NA+	6058.92	-.579	FE(OH)4--	.00	-17.278	ALF4-	.00	-16.460
K+	965.61	-1.607	FE(OH)++	.00	-13.460	ALF5--	.00	-19.759
CA++	1061.07	-1.577	FE(OH)2+	.00	-8.793	ALF6---	.00	-24.067

IONIC STRENGTH = .38149 IONIC BALANCE : CATIONS (MOL.EQ.) .34167250
 ANIONS (MOL.EQ.) .36794220
 DIFFERENCE (%) -7.40

OXIDATION POTENTIAL (VOLTS) : EH H2S= -.362 EH CH4= -.421 EH H2= -.409 EH NH3= 99.999

LOG SOLUBILITY PRODUCTS OF MINERALS IN DEEP WATER

	TEOR.	CALC.		TEOR.	CALC.		TEOR.	CALC.
ADULARIA	-14.776	-14.303	ALBITE LOW	-14.271	-13.240	ANALCIME	-11.648	-11.094
ANHYDRITE	-7.207	-7.320	CALCITE	-11.526	-11.709	CHALCEDONY	-2.217	-2.146
MG-CHLORITE	-81.685	-89.234	FLUORITE	-10.718	-13.459	GORTHITE	-.350	-1.614
LAUMONTITE	-24.602	-22.913	MICROCLINE	-15.613	-14.303	MAGNETITE	-21.553	-21.712
CA-MONTMOR.	-73.343	-64.068	K-MONTMOR.	-34.669	-32.734	MG-MONTMOR.	-74.794	-66.635
NA-MONTMOR.	-34.902	-31.671	MUSCOVITE	-18.071	-14.724	PREHNITE	-36.095	-34.675
PYRRHOTITE	-48.482	-55.348	PYRITE	-76.641	-66.074	QUARTZ	-2.340	-2.146
WAIKAKITE	-23.902	-22.913	WOLLASTONITE	8.387	6.455	ZOISITE	-36.177	-34.885
EPIDOTE	-37.293	-36.289	MARCASITE	-58.102	-66.074			

Table 6. (continued)

UNU Geothermal Training Programme
SAMPLE = 08

DEEP WATER COOLED TO 150.0 DEGREES C.

ACTIVITY COEFFICIENTS IN DEEP WATER

H+	.687	KSO4-	.609	FE++	.173	FECL+	.566
OH-	.547	F-	.547	FE+++	.047	AL+++	.047
H3SIO4-	.566	CL-	.528	FE(OH)+	.598	ALOH++	.156
H2SIO4--	.156	NA+	.566	FE(OH)3-	.598	AL(OH)2+	.609
H2BO3-	.506	K+	.528	FE(OH)4--	.144	AL(OH)4-	.582
HCO3-	.566	CA++	.173	FE(OH)++	.144	ALSO4+	.582
CO3--	.130	MG++	.227	FE(OH)2+	.609	AL(SO4)2-	.582
HS-	.547	CAHCO3+	.625	FE(OH)4-	.609	ALF++	.156
S--	.144	MGNCO3+	.566	FESO4+	.598	ALF2+	.609
HSO4-	.582	CAOH+	.625	FECL++	.144	ALF4-	.582
SO4--	.116	MGOH+	.637	FECL2+	.598	ALF5--	.130
NASO4-	.609	NH4+	.506	FECL4-	.566	ALF6---	.011

CHEMICAL COMPONENTS IN DEEP WATER (PPM AND LOG MOLE)

H+ (ACT.)	.00	-5.339	MG++	1.28	-4.278	FE(OH)3	.00	-9.817
OH-	.02	-6.035	NACL	794.59	-1.867	FE(OH)4-	.00	-10.885
H4SIO4	687.09	-2.146	KCL	35.86	-3.318	FECL+	.14	-5.828
H3SIO4-	.44	-5.333	NASO4-	5.32	-4.350	FECL2	.00	-13.611
H2SIO4--	.00	-10.374	KSO4-	2.65	-4.708	FECL++	.00	-16.764
NAH3SIO4	.62	-5.278	CASO4	11.19	-4.085	FECL2+	.00	-17.349
H3BO3	39.78	-3.192	MGSO4	.10	-6.094	FECL3	.00	-18.702
H2BO3-	.02	-6.472	CACO3	.03	-6.499	FECL4-	.00	-20.233
H2CO3	202.75	-2.486	MGCO3	.00	-9.857	FESO4	.00	-8.693
HCO3-	15.60	-3.592	CAHCO3+	28.13	-3.556	FESO4+	.00	-18.875
CO3--	.00	-7.945	MGNCO3+	.01	-6.977	AL+++	.00	-13.463
H2S	3.49	-3.989	CAOH+	.02	-6.567	ALOH++	.00	-9.728
HS-	.30	-5.043	MGOH+	.00	-8.126	AL(OH)2+	.01	-6.762
S--	.00	-14.737	NH4OH	.00	.000	AL(OH)3	.15	-5.724
H2SO4	.00	-13.039	NH4+	.00	.000	AL(OH)4-	.00	-7.611
HSO4-	.09	-6.042	FE++	.07	-5.882	ALSO4+	.00	-15.193
SO4--	18.35	-3.719	FE+++	.00	-19.541	AL(SO4)2-	.00	-17.531
HF	.01	-6.459	FE(OH)+	.00	-7.618	ALF++	.00	-11.406
F-	.12	-5.192	FE(OH)2	.00	-10.509	ALF2+	.00	-10.964
CL-	13184.71	-.430	FE(OH)3-	.00	-14.141	ALF3	.00	-12.089
NA+	6152.23	-.573	FE(OH)4--	.00	-20.612	ALF4-	.00	-14.662
K+	982.01	-1.600	FE(OH)++	.00	-14.404	ALF5--	.00	-18.074
CA++	1066.01	-1.575	FE(OH)2+	.00	-11.047	ALF6---	.00	-22.337

IONIC STRENGTH = .38628

IONIC BALANCE : CATIONS (MOL.EQ.) .34630230
ANIONS (MOL.EQ.) .37258310
DIFFERENCE (%) -7.31

OXIDATION POTENTIAL (VOLTS) : BH H2S= -.275 BH CH4= -.328 BH H2= -.358 BH NH3= 99.999

LOG SOLUBILITY PRODUCTS OF MINERALS IN DEEP WATER

	TEOR.	CALC.		TEOR.	CALC.		TEOR.	CALC.
ADULARIA	-15.664	-14.231	ALBITE LOW	-15.067	-13.173	ANALCIME	-12.168	-11.027
ANHYDRITE	-6.366	-6.993	CALCITE	-10.399	-11.169	CHALCEDONY	-2.492	-2.146
MG-CHLORITE	-80.047	-93.252	FLUORITE	-10.561	-13.245	GOETHITE	-2.669	-4.804
LAUMONTITE	-25.402	-22.752	NICROCLINE	-16.731	-14.231	MAGNETITE	-25.789	-28.845
CA-MONTMOR.	-76.836	-56.804	K-MONTMOR.	-36.807	-29.111	MG-MONTMOR.	-78.177	-59.389
NA-MONTMOR.	-36.953	-28.053	MUSCOVITE	-18.991	-13.468	PREHNITE	-35.828	-35.536
PYRRHOTITE	-74.605	-82.269	PYRITE	-111.470	-90.329	QUARTZ	-2.683	-2.146
WAIRAKITE	-23.901	-22.752	WOLLASTONITE	9.469	6.195	ZOISITE	-35.407	-35.155
EPIDOTE	-39.779	-40.340	NARCASITE	-90.765	-90.329			

Table 6. (continued)

UNU Geothermal Training Programme
SAMPLE = 08

DBBP WATER COOLED TO 100.0 DEGREES C.

ACTIVITY COEFFICIENTS IN DBBP WATER

H+	.716	HSO4-	.645	Fe++	.213	FeCL+	.605
OH-	.588	F-	.588	Fe+++	.068	AL+++	.068
H3SiO4-	.605	CL-	.569	FeOH+	.634	ALOH++	.195
H2SiO4--	.195	NA+	.605	Fe(OH)3-	.634	AL(OH)2+	.645
H2BO3-	.549	K+	.569	Fe(OH)4--	.183	AL(OH)4-	.620
HCO3-	.605	CA++	.213	FeOH++	.183	ALSO4+	.620
CO3--	.167	MG++	.270	Fe(OH)2+	.645	AL(SO4)2-	.620
HS-	.588	CAHCO3+	.659	Fe(OH)4-	.645	ALP++	.195
S--	.183	MGHCO3+	.605	FeSO4+	.634	ALP2+	.645
HSO4-	.620	CAOH+	.659	FeCL++	.183	ALP4-	.620
SO4--	.151	MGOH+	.670	FeCL2+	.634	ALP5--	.167
NASO4-	.645	NH4+	.549	FeCL4-	.605	ALP6---	.019

CHEMICAL COMPONENTS IN DBBP WATER (PPM AND LOG MOL%)

H+ (ACT.)	.01	-5.251	MG++	1.29	-4.276	Fe(OH)3	.00	-13.028
OH-	.00	-6.778	NaCL	456.81	-2.107	Fe(OH)4-	.00	-14.694
H4SiO4	687.55	-2.146	KCL	20.07	-3.570	FeCL+	.07	-6.089
H3SiO4-	.22	-5.646	NASO4-	4.45	-4.427	FeCL2	.00	-17.307
H2SiO4--	.00	-10.895	HSO4-	2.05	-4.819	FeCL++	.00	-17.671
NAH3SiO4	.34	-5.546	CASO4	8.37	-4.211	FeCL2+	.00	-18.190
H3BO3	39.79	-3.191	MGSO4	.07	-6.252	FeCL3	.00	-19.684
H2BO3-	.01	-6.621	CACO3	.02	-6.776	FeCL4-	.00	-21.496
H2CO3	202.33	-2.487	MGCO3	.00	-9.902	FeSO4	.00	-8.446
HCO3-	23.12	-3.422	CAHCO3+	16.36	-3.791	FeSO4+	.00	-19.536
CO3--	.00	-7.744	MGHCO3+	.01	-6.939	AL+++	.00	-10.298
H2S	3.53	-3.985	CAOH+	.00	-7.492	ALOH++	.00	-7.939
HS-	.27	-5.092	MGOH+	.00	-9.089	AL(OH)2+	.04	-6.234
S--	.00	-15.413	NH4OH	.00	.000	AL(OH)3	.11	-5.839
H2SO4	.00	-13.905	NH4+	.00	.000	AL(OH)4-	.00	-7.388
HSO4-	.03	-6.525	Fe++	.11	-5.699	ALSO4+	.00	-12.181
SO4--	21.54	-3.649	Fe+++	.00	-19.517	AL(SO4)2-	.00	-14.547
HF	.00	-6.823	FeOH+	.00	-8.347	ALP++	.00	-8.717
F-	.13	-5.179	Fe(OH)2	.00	-12.340	ALP2+	.00	-8.637
CL-	13397.14	-.423	Fe(OH)3-	.00	-17.365	ALP3	.00	-9.982
NA+	6285.34	-.563	Fe(OH)4--	.00	-24.460	ALP4-	.00	-12.700
K+	990.46	-1.596	Fe(OH)++	.00	-15.451	ALP5--	.00	-16.196
CA++	1071.52	-1.573	Fe(OH)2+	.00	-13.336	ALP6---	.00	-20.404

IONIC STRENGTH = .39263 IONIC BALANCE : CATIONS (MOL.EQ.) .35246830

ANIONS (MOL.EQ.) .37876140

DIFFERENCE (%) -7.19

OXIDATION POTENTIAL (VOLTS) : BH H2S= -.198 BH CH4= -.245 BH H2= -.313 BH NH3= 99.999

LOG SOLUBILITY PRODUCTS OF MINERALS IN DBBP WATER

	TEOR.	CALC.		TEOR.	CALC.		TEOR.	CALC.
ADULARIA	-17.222	-14.165	ALBITE LOW	-16.485	-13.106	ANALCINE	-13.199	-10.961
ANHYDRITE	-5.607	-6.714	CALCITE	-9.438	-10.765	CHALCEDONY	-2.841	-2.146
MG-CHLORITE	-80.302	-98.500	FLUORITE	-10.538	-13.063	GOETHITE	-4.674	-7.877
LAUMONTITE	-27.182	-22.602	MICROCLINE	-18.598	-14.165	MAGNETITE	-29.703	-36.139
CA-MONTMOR.	-85.559	-47.778	K-MONTMOR.	-41.648	-24.608	MG-MONTMOR.	-86.708	-50.378
NA-MONTMOR.	-41.660	-23.549	MUSCOVITE	-21.151	-11.925	PREHNITE	-36.778	-36.717
PYRRHOTITE	-99.806	-108.424	PYRITE	-147.012	-113.241	QUARTZ	-3.098	-2.146
WAIRAKITE	-24.710	-22.602	WOLLASTONITE	10.824	6.112	ZOISITE	-35.893	-35.597
EPIDOTE	-43.603	-44.593	MARCASITE	-123.575	-113.241			

Table 6. (continued)

UNU Geothermal Training Programme
SAMPLE = 08

DEEP WATER COOLED TO 50.0 DEGREES C.

ACTIVITY COEFFICIENTS IN DEEP WATER

H+	.738	HSO4-	.672	FE++	.248	FECL+	.635
OH-	.619	F-	.619	FE+++	.088	AL+++	.088
H3SIO4-	.635	CL-	.602	FE(OH)+	.662	AL(OH)++	.229
H2SIO4--	.229	NA+	.635	FE(OH)3-	.662	AL(OH)2+	.672
H2BO3-	.583	K+	.602	FE(OH)4--	.216	AL(OH)4-	.649
HCO3-	.635	CA++	.248	FE(OH)++	.216	ALSO4+	.649
CO3--	.199	MG++	.307	FE(OH)2+	.672	AL(SO4)2-	.649
HS-	.619	CAHCO3+	.686	FE(OH)4-	.672	ALP++	.229
S--	.216	MGHCO3+	.635	FE(SO4)+	.662	ALP2+	.672
HSO4-	.649	CAOH+	.686	FECL++	.216	ALP4-	.649
SO4--	.182	MGOH+	.696	FECL2+	.662	ALP5--	.199
NASO4-	.672	NH4+	.583	FECL4-	.635	ALP6---	.028

CHEMICAL COMPONENTS IN DEEP WATER (PPM AND LOG MOLE)

H+ (ACT.)	.01	-5.295	MG++	1.29	-4.274	FE(OH)3	.00	-16.361
OH-	.00	-7.763	NACL	131.92	-2.646	FE(OH)4-	.00	-18.705
H4SIO4	687.87	-2.145	KCL	10.99	-3.831	FECL+	.02	-6.604
H3SIO4-	.07	-6.144	NASO4-	3.38	-4.547	FECL2	.00	-22.291
H2SIO4--	.00	-11.775	HSO4-	1.43	-4.977	FECL++	.00	-18.712
NAH3SIO4	.12	-5.976	CASO4	6.14	-4.346	FECL2+	.00	-19.178
H3BO3	39.79	-3.191	MGSO4	.03	-6.562	FECL3	.00	-20.858
H2BO3-	.01	-6.768	CACO3	.01	-7.183	FECL4-	.00	-23.011
H2CO3	201.94	-2.487	MGCO3	.00	-10.077	FE(SO4)	.00	-8.283
HCO3-	29.43	-3.317	CAHCO3+	6.55	-4.189	FE(SO4)+	.00	-20.314
CO3--	.00	-7.706	MGHCO3+	.01	-6.956	AL+++	.00	-7.573
H2S	3.62	-3.973	CAOH+	.00	-8.654	ALOH++	.01	-6.591
HS-	.18	-5.276	MGOH+	.00	-10.260	AL(OH)2+	.04	-6.163
S--	.00	-16.213	NH4OH	.00	.000	AL(OH)3	.03	-6.433
H2SO4	.00	-15.179	NH4+	.00	.000	AL(OH)4-	.00	-7.954
HSO4-	.01	-7.129	FE++	.14	-5.590	ALSO4+	.00	-9.607
SO4--	24.48	-3.594	FE+++	.00	-19.643	AL(SO4)2-	.00	-11.995
HF	.00	-7.341	FE(OH)+	.00	-9.310	ALP++	.02	-6.349
F-	.11	-5.222	FE(OH)2	.00	-14.585	ALP2+	.02	-6.552
CL-	13598.55	-.416	FE(OH)3-	.00	-21.224	ALP3	.00	-8.073
NA+	6413.40	-.554	FE(OH)4--	.00	-29.182	ALP4-	.00	-10.922
K+	995.40	-1.594	FE(OH)++	.00	-16.580	ALP5--	.00	-14.504
CA++	1076.07	-1.571	FE(OH)2+	.00	-15.710	ALP6---	.00	-18.668

IONIC STRENGTH = .39861

IONIC BALANCE : CATIONS (MOL.EQ.) .35829780
ANIONS (MOL.EQ.) .38460230
DIFFERENCE (%) -7.08

OXIDATION POTENTIAL (VOLTS) : EH H2S= -.130 EH CH4= -.172 EH H2= -.274 EH NH3= 99.999

LOG SOLUBILITY PRODUCTS OF MINERALS IN DEEP WATER

	TBOB.	CALC.		TBOB.	CALC.		TBOB.	CALC.
ADULARIA	-19.664	-14.119	ALBITE LOW	-18.715	-13.056	ANALCIME	-14.883	-10.911
ANHYDRITE	-4.966	-6.510	CALCITE	-8.720	-10.583	CHALCEDONY	-3.297	-2.145
MG-CHLORITE	-82.885	-105.881	FLUORITE	-10.737	-13.038	GORTHITE	-6.429	-10.906
LAUMONTITE	-30.213	-22.494	MICROCLINE	-21.446	-14.119	MAGNETITE	-33.553	-43.951
CA-MONTHOR.	-101.366	-35.869	K-MONTHOR.	-50.171	-18.661	MG-MONTHOR.	-102.214	-38.480
NA-MONTHOR.	-49.974	-17.598	MUSCOVITE	-25.002	-9.912	PREHNITE	-39.270	-38.468
PYRRHOTITE	-125.035	-134.618	PYRITE	-185.595	-135.240	QUARTZ	-3.641	-2.145
WAIBAKITE	-26.535	-22.494	WOLLASTONITE	12.603	6.268	ZOISITE	-37.927	-36.365
EPIDOTE	-47.739	-49.374	MARCASITE	-158.572	-135.240			

Table 6. (continued)

UNU Geothermal Training Programme
SAMPLE = 08

DEEP WATER COOLED TO .0 DEGREES C.

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ACTIVITY COEFFICIENTS IN DEEP WATER

H+	.746	HSO4-	.682	FE++	.259	FECL+	.646
OH-	.630	F-	.630	FE+++	.095	AL+++	.095
H3SiO4-	.646	CL-	.614	FE(OH)+	.672	AL(OH)++	.240
H2SiO4--	.240	NA+	.646	FE(OH)3-	.672	AL(OH)2+	.682
H2BO3-	.595	K+	.614	FE(OH)4--	.227	AL(OH)4-	.660
HCO3-	.646	CA++	.259	FE(OH)++	.227	ALSO4+	.660
CO3--	.210	MG++	.318	FE(OH)2+	.682	AL(SO4)2-	.660
HS-	.630	CAHCO3+	.695	FE(OH)4-	.682	ALP++	.240
S--	.227	MGHCO3+	.646	FESO4+	.672	ALP2+	.682
HSO4-	.660	CAOH+	.695	FECL++	.227	ALP4-	.660
SO4--	.193	MGOH+	.705	FECL2+	.672	ALP5--	.210
NASO4-	.682	NH4+	.595	FECL4-	.646	ALP6---	.031

CHEMICAL COMPONENTS IN DEEP WATER (PPM AND LOG MOLE)

H+ (ACT.)	.00	-5.585	MG++	1.30	-4.272	FE(OH)3	.00	-20.320
OH-	.00	-9.141	NaCl	59.01	-2.996	FE(OH)4-	.00	-23.498
H4SiO4	688.01	-2.145	KCl	5.72	-4.115	FECL+	.00	-7.475
H3SiO4-	.01	-6.847	NASO4-	2.13	-4.748	FECL2	.00	-29.230
H2SiO4--	.00	-13.082	HSO4-	.80	-5.229	FECL++	.00	-20.208
NAH3SiO4	.03	-6.591	CASO4	4.40	-4.491	FECL2+	.00	-20.657
H3BO3	39.79	-3.191	MGSO4	.01	-7.076	FECL3	.00	-22.608
H2BO3-	.01	-6.895	CACO3	.00	-7.765	FECL4-	.00	-25.200
H2CO3	201.68	-2.488	MGCO3	.00	-10.413	FESO4	.00	-8.228
HCO3-	32.64	-3.272	CAHCO3+	1.66	-4.784	FESO4+	.00	-21.523
CO3--	.00	-7.836	MGHCO3+	.01	-7.041	AL+++	.00	-6.897
H2S	3.72	-3.962	CAOH+	.00	-10.222	AL(OH)++	.00	-7.391
HS-	.08	-5.590	MGOH+	.00	-11.944	AL(OH)2+	.00	-8.359
S--	.00	-17.082	NH4OH	.00	.000	AL(OH)3	.00	-9.381
H2SO4	.00	-17.156	NH4+	.00	.000	AL(OH)4-	.00	-11.309
HSO4-	.00	-7.953	FE++	.16	-5.555	ALSO4+	.00	-9.119
SO4--	27.19	-3.548	FE+++	.00	-20.182	AL(SO4)2-	.00	-11.530
HF	.00	-8.104	FE(OH)+	.00	-10.616	ALP++	.06	-5.871
F-	.09	-5.314	FE(OH)2	.00	-17.501	ALP2+	.04	-6.256
CL-	13645.29	-.415	FE(OH)3-	.00	-26.146	ALP3	.00	-7.876
NA+	6442.34	-.552	FE(OH)4--	.00	-35.344	ALP4-	.00	-10.818
K+	998.35	-1.593	FE(OH)++	.00	-18.112	ALP5--	.00	-14.456
CA++	1078.52	-1.570	FE(OH)2+	.00	-18.587	ALP6---	.00	-18.544

IONIC STRENGTH = .40012

IONIC BALANCE : CATIONS (MOL.BQ.) .35970790
ANIONS (MOL.BQ.) .38602530
DIFFERENCE (%) -7.06

OXIDATION POTENTIAL (VOLTS) : BH H2S= -.074 BH CH4= -.112 BH H2= -.244 BH NH3= 99.999

LOG SOLUBILITY PRODUCTS OF MINERALS IN DEEP WATER

	TEOR.	CALC.		TEOR.	CALC.		TEOR.	CALC.
ADULARIA	-23.276	-14.102	ALBITE LOW	-22.008	-13.039	ANALCIME	-17.390	-10.894
ANHYDRITE	-4.511	-6.419	CALCITE	-8.378	-10.670	CHALCEDONY	-3.921	-2.145
MG-CHLORITE	-88.387	-116.737	FLUORITE	-11.323	-13.185	GORTHITE	-8.225	-14.323
LAUMONTITE	-34.804	-22.459	MICROCLINE	-25.573	-14.102	MAGNETITE	-38.146	-53.470
CA-MONTMOR.	-121.263	-19.312	K-MONTMOR.	-61.001	-10.383	MG-MONTMOR.	-121.654	-21.926
NA-MONTMOR.	-60.441	-9.321	MUSCOVITE	-30.074	-7.142	PREHNITE	-43.653	-41.153
PYRRHOTITE	-152.446	-164.666	PYRITE	-231.987	-159.601	QUARTZ	-4.382	-2.145
WAIRAKITE	-29.576	-22.459	WOLLASTONITE	15.039	6.868	ZOISITE	-41.763	-37.673
EPIDOTE	-54.230	-55.476	MARCASITE	-200.020	-159.601			

Table 7. Example of a printout from the WATCH3 program

UNU Geothermal Training Programme

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073                seltjarnarnes well sn-2 570m depth

PROGRAM WATCH2.

WATER SAMPLE (PPM)          STEAM SAMPLE

PH/DEG.C          8.45/21.0  GAS (VOL.%)
SI02              92.30      CO2
NA               595.60      H2S
K                14.50      H2
CA              324.70      O2
MG               .280      CH4
CO2             11.00      N2
SO4             221.70
H2S              .10
CL             1321.00
F                .55
DISS.SOLIDS     2582.00
AL               .1370
B                .0000
FE              .0660
NH3             .0000
Mn              0.00

LITERS GAS PER KG
CONDENSATE/DEG.C

CONDENSATE (PPM)
PH/DEG.C
CO2
H2S
NA

CONDENSATE WITH NaOH (PPM)
CO2
H2S

REFERENCE TEMP.          DEGREES C          .0 (CHA)
SAMPLING PRESSURE       BARS ABS.
DISCHARGE ENTHALPY     MJ/OL/KG
DISCHARGE                KG/SEC.          .0
MEASURED TEMPERATURE   DEGREES C          .0
RESISTIVITY/TEMP.     OHMM/DEG.C    4270.0/21.1
BH/TEMP.                MV/DEG.C       .000/ .0

MEASURED DOWNHOLE TEMP.  FLUID INFLOW
DEGREES C/METERS        DEPTH (METERS)

.0      .0      .0
.0      .0      .0
.0      .0      .0
.0      .0      .0
.0      .0      .0
.0      .0      .0
.0      .0      .0
.0      .0      .0
.0      .0      .0
.0      .0      .0
.0      .0      .0
.0      .0      .0
.0      .0      .0
.0      .0      .0
.0      .0      .0
.0      .0      .0
.0      .0      .0
.0      .0      .0
.0      .0      .0

IONIC STRENGTH = .05133      IONIC BALANCE :  CATIONS (MOL.EQ.) .04176022
ANIONS (MOL.EQ.) .04146050
DIFFERENCE (%)      .72

DEEP WATER (PPM)          DEEP STEAM (PPM)          GAS PRESSURES (BARS ABS.)

SI02          92.31          CO2          11.00          CO2          .00          CO2          .138E-04
NA           595.60          H2S           .10          H2S          .00          H2S          .103E-06
K            14.50          H2            .00          H2           .00          H2           .000E+00
CA          324.70          O2            .00          O2           .00          O2           .000E+00
MG           .280          CH4           .00          CH4          .00          CH4          .000E+00
SO4         221.70          N2            .00          N2           .00          N2           .000E+00
CL         1320.89          NH3           .00          NH3          .00          NH3          .000E+00
F            .55
DISS.S.     2582.00
AL           .1370
B            .0000
FE          .0660

H2O (%)          .00
BOILING PORTION  .00
TOTAL            .105E+01
    
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Table 7. (continued)

SAMPLE = 073

ACTIVITY COEFFICIENTS IN DEEP WATER

H+	.829	KS04-	.800	FE++	.427	FECL+	.785
OH-	.780	F-	.780	FE+++	.192	AL+++	.192
H3SIO4-	.785	CL-	.774	FEOH+	.796	ALOH++	.415
H2SIO4--	.415	NA+	.785	FE(OH)3-	.796	AL(OH)2+	.800
H2BO3-	.768	K+	.774	FE(OH)4--	.407	AL(OH)4-	.791
HCO3-	.785	CA++	.427	FEOH++	.407	ALSO4+	.791
CO3--	.397	MG++	.463	FE(OH)2+	.800	AL(SO4)2-	.791
HS-	.780	CAHCO3+	.805	FE(OH)4-	.800	ALF++	.415
S--	.407	MGHCO3+	.785	FESO4+	.796	ALF2+	.800
HSO4-	.791	CAOH+	.805	FECL++	.407	ALF4-	.791
SO4--	.387	MGOH+	.810	FECL2+	.796	ALF5--	.397
NASO4-	.800	NH4+	.768	FECL4-	.785	ALF6---	.126

CHEMICAL COMPONENTS IN DEEP WATER (PPM AND LOG MOLE)

H+ (ACT.)	.00	-7.423	MG++	.21	-5.067	FE(OH)3	.00	-8.845
OH-	.33	-4.711	NACL	7.62	-3.885	FE(OH)4-	.00	-8.420
H4SIO4	141.41	-2.832	KCL	.05	-6.145	FECL+	.01	-7.230
H3SIO4-	5.15	-4.266	NASO4-	8.11	-4.167	FECL2	.00	-19.126
H2SIO4--	.00	-7.553	KS04-	.60	-5.352	FECL++	.00	-21.275
NAH3SIO4	1.28	-4.966	CASO4	86.30	-3.198	FECL2+	.00	-22.418
H3BO3	.00	.000	MGSO4	.35	-5.538	FECL3	.00	-24.683
H2BO3-	.00	.000	CACO3	.88	-5.057	FECL4-	.00	-27.477
H2CO3	.85	-4.864	MGCO3	.00	-8.489	FESO4	.01	-7.346
HCO3-	11.00	-3.744	CAHCO3+	4.71	-4.332	FESO4+	.00	-20.756
CO3--	.04	-6.159	MGHCO3+	.00	-7.815	AL+++	.00	-17.418
H2S	.01	-6.523	CAOH+	.13	-5.640	ALOH++	.00	-12.737
HS-	.09	-5.579	MGOH+	.00	-7.534	AL(OH)2+	.00	-8.598
S--	.00	-13.942	NH4OH	.00	.000	AL(OH)3	.03	-5.924
H2SO4	.00	-16.963	NH4+	.00	.000	AL(OH)4-	.37	-5.411
HSO4-	.00	-7.527	FE++	.04	-6.166	ALSO4+	.00	-17.687
SO4--	153.54	-2.796	FE+++	.00	-22.369	AL(SO4)2-	.00	-18.788
HF	.00	-8.223	FEOH+	.03	-6.418	ALF++	.00	-14.944
F-	.55	-4.539	FE(OH)2	.00	-8.116	ALF2+	.00	-10.859
CL-	1316.24	-1.430	FE(OH)3-	.00	-11.037	ALF3	.00	-14.342
NA+	590.79	-1.590	FE(OH)4--	.00	-16.196	ALF4-	.00	-16.400
K+	14.30	-3.437	FE(OH)++	.00	-16.007	ALF5--	.00	-13.464
CA++	296.98	-2.130	FE(OH)2+	.00	-11.438	ALF6---	.00	-23.304

IONIC STRENGTH = .04980

IONIC BALANCE :

CATIONS (MOL.EQ.) .04095064

ANIONS (MOL.EQ.) .04061699

DIFFERENCE (%) .82

CHEMICAL GEOTHERMOMETERS DEGREES C

1000/T DEGREES KELVIN = 2.67

QUARTZ	124.7
CHALCEDONY	101.0
NAR	82.7

OXIDATION POTENTIAL (VOLTS) : EH H2S= -.366 EH CH4= 99.999 EH H2= 99.999 EH NH3= 99.999

LOG SOLUBILITY PRODUCTS OF MINERALS IN DEEP WATER

	TEOR.	CALC.		TEOR.	CALC.		TEOR.	CALC.
ADULARIA	-17.181	-17.441	ALBITE LOW	-16.448	-15.588	ANALCINE	-13.171	-12.756
ANHYDRITE	-5.621	-5.709	CALCITE	-9.456	-9.059	CHALCEDONY	-2.832	-2.832
MG-CHLORITE	-80.275	-84.852	FLUORITE	-10.536	-11.794	GOETHITE	-4.636	-3.636
LAUMONTITE	-27.133	-24.622	MICROCLINE	-18.550	-17.441	MAGNETITE	-29.623	-25.576
CA-MONTMOR.	-85.307	-82.527	K-MONTMOR.	-41.511	-43.562	MG-MONTMOR.	-86.461	-85.423
NA-MONTMOR.	-41.526	-41.709	MUSCOVITE	-21.090	-18.595	PREHNITE	-36.744	-35.328
PYRRHOTITE	-99.285	-94.726	PYRITE	-146.252	-121.007	QUARTZ	-3.086	-2.632
WATERITE	-24.684	-24.622	WOLLASTONITE	10.792	9.515	ZOISITE	-35.868	-34.505
EPIDOTE	-43.519	-37.626	MARCASITE	-122.879	-121.007			

Table 7. (continued)

UNU Geothermal Training Programme
SAMPLE = 073

DEEP WATER COOLED TO 200.0 DEGREES C.

ACTIVITY COEFFICIENTS IN DEEP WATER

H+	.782	KS04-	.745	FB++	.324	FECL+	.726
OH-	.719	F-	.719	FB+++	.114	AL+++	.114
H3SI04-	.726	CL-	.711	FBOH+	.740	ALOH++	.312
H2SI04--	.312	NA+	.726	FE(OH)3-	.740	AL(OH)2+	.745
H2BO3-	.703	K+	.711	FE(OH)4--	.304	AL(OH)4-	.733
HCO3-	.726	CA++	.324	FBOH++	.304	ALSO4+	.733
CO3--	.293	MG++	.362	FE(OH)2+	.745	AL(SO4)2-	.733
HS-	.719	CAHCO3+	.752	FE(OH)4-	.745	ALF++	.312
S--	.304	MGHCO3+	.726	FBSO4+	.740	ALF2+	.745
HSO4-	.733	CAOH+	.752	FECL++	.304	ALF4-	.733
SO4--	.283	MGOH+	.757	FECL2+	.740	ALF5--	.293
NASO4-	.745	NH4+	.703	FECL4-	.726	ALF6---	.064

CHEMICAL COMPONENTS IN DEEP WATER (PPM AND LOG MOLE)

H+ (ACT.)	.00	-7.067	MG++	.16	-5.173	FE(OH)3	.01	-6.915
OH-	1.53	-4.045	NACL	20.18	-3.462	FE(OH)4-	.13	-5.975
H4SI04	143.52	-2.826	KCL	.21	-5.552	FECL+	.00	-10.159
H3SI04-	3.40	-4.446	NASO4-	12.08	-3.994	FECL2	.00	-15.902
H2SI04--	.00	-8.123	KS04-	1.07	-5.102	FRCL++	.00	-22.229
NAH3SI04	.85	-5.141	CASO4	174.01	-2.893	FECL2+	.00	-23.389
H3RO3	.00	.000	MGSO4	.52	-5.363	FECL3	.00	-25.334
H2BO3-	.00	.000	CACO3	.72	-5.143	FECL4-	.00	-27.588
H2CO3	3.69	-4.226	MGCO3	.00	-9.123	FBSO4	.00	-11.149
HCO3-	4.38	-4.144	CAHCO3+	11.27	-3.953	FBSO4+	.00	-22.133
CO3--	.00	-7.385	MGHCO3+	.00	-7.980	AL+++	.00	-21.707
H2S	.04	-5.988	CAOH+	1.62	-4.546	ALOH++	.00	-14.831
HS-	.06	-5.720	MGOH+	.02	-6.340	AL(OH)2+	.00	-8.706
S--	.00	-13.522	NH4OH	.00	.000	AL(OH)3	.36	-5.332
H2SO4	.00	-14.264	NH4+	.00	.000	AL(OH)4-	.04	-6.372
HSO4-	.09	-6.037	FB++	.00	-9.947	ALSO4+	.00	-21.540
SO4--	87.90	-3.039	FB+++	.00	-25.278	AL(SO4)2-	.00	-22.548
HF	.00	-6.893	FROH+	.00	-9.053	ALF++	.00	-18.048
F-	.55	-4.541	FE(OH)2	.00	-9.252	ALF2+	.00	-16.004
CL-	1308.55	-1.433	FE(OH)3-	.00	-10.175	ALF3	.00	-15.884
NA+	585.16	-1.594	FE(OH)4--	.00	-14.948	ALF4-	.00	-17.570
K+	14.08	-3.444	FE(OH)++	.00	-17.288	ALF5--	.00	-20.386
CA++	267.57	-2.175	FE(OH)2+	.00	-10.754	ALF6---	.00	-24.440

IONIC STRENGTH = .04674

IONIC BALANCE : CATIONS (MOL.EQ.) .03931884
ANIONS (MOL.EQ.) .03897264
DIFFERENCE (%) .88

OXIDATION POTENTIAL (VOLTS) : EH H2S= -.505 EH CH4= 99.999 EH H2= 99.999 EH NH3= 99.999

LOG SOLUBILITY PRODUCTS OF MINERALS IN DEEP WATER

	TEOR.	CALC.		TEOR.	CALC.		TEOR.	CALC.
ADULARIA	-14.776	-17.498	ALBITE LOW	-14.271	-15.640	ANALCIME	-11.646	-12.814
ANHYDRITE	-7.207	-6.252	CALCITE	-11.526	-10.582	CHALCEDONY	-2.217	-2.826
MG-CHLORITE	-81.685	-80.911	FLUORITE	-10.718	-12.033	GORTHITE	-.350	-1.916
LAUMONTITE	-24.602	-24.827	MICROCLINE	-15.613	-17.498	MAGNETITE	-21.553	-22.643
CA-MONTHOR.	-73.343	-90.588	K-MONTHOR.	-34.669	-47.553	MG-MONTHOR.	-74.794	-93.537
NA-MONTHOR.	-34.902	-45.695	MUSCOVITE	-18.071	-19.981	PREHNITE	-36.095	-33.041
PYRRHOTITE	-48.482	-74.001	PYRITE	-76.641	-100.350	QUARTZ	-2.340	-2.826
WATRAKITE	-23.902	-24.827	WOLLASTONITE	8.387	8.643	ZOISITE	-36.177	-34.280
EPIDOTE	-37.293	-34.957	MARCASITE	-58.102	-100.350			

Table 7. (continued)

UNU Geothermal Training Programme
 SAMPLE = 073

DEEP WATER COOLED TO 150.0 DEGREES C.

ACTIVITY COEFFICIENTS IN DEEP WATER

H+	.810	KS04-	.778	FE++	.385	FECL+	.763
OH-	.757	F-	.757	FE+++	.157	AL+++	.157
H3ST04-	.763	CL-	.751	FE0H+	.774	AL0H++	.373
H2ST04--	.373	NA+	.763	FE(OH)3-	.774	AL(OH)2+	.778
H2B03-	.744	K+	.751	FE(OH)4--	.365	AL(OH)4-	.769
HCO3-	.763	CA++	.385	FE0H++	.365	ALSO4+	.769
CO3--	.355	MG++	.421	FE(OH)2+	.778	AL(SO4)2-	.769
HS-	.757	CAHCO3+	.785	FE(OH)4-	.778	ALF++	.373
S--	.365	MGHCO3+	.763	FESO4+	.774	ALF2+	.778
HSO4-	.769	CAOH+	.785	FECL+	.365	ALF4-	.769
SO4--	.344	MGOH+	.789	FECL2+	.774	ALF5--	.355
NASO4-	.778	NH4+	.744	FECL4-	.763	ALF6---	.098

CHEMICAL COMPONENTS IN DEEP WATER (PPM AND LOG MOLE)

H+ (ACT.)	.00	-7.172	MG++	.18	-5.135	FE(OH)3	.01	-6.869
OH-	.77	-4.343	NACL	14.50	-3.605	FE(OH)4-	.08	-6.211
H4ST04	142.08	-2.830	RCL	.10	-5.853	FECL+	.00	-7.513
H3ST04-	4.61	-4.315	NASO4-	10.51	-4.054	FECL2	.00	-16.014
H2ST04--	.00	-7.771	KS04-	.83	-5.209	FECL++	.00	-20.567
NAH3ST04	1.13	-5.019	CASO4	129.36	-3.022	FECL2+	.60	-21.710
H3B03	.00	.000	MGS04	.49	-5.390	FECL3	.00	-23.800
H2B03-	.00	.000	CACO3	.83	-5.079	FECL4-	.00	-26.310
H2CO3	1.95	-4.502	MGCCO3	.00	-8.796	FESO4	.00	-8.107
HCO3-	7.58	-3.906	CAHCO3+	8.67	-4.067	FESO4+	.00	-20.247
CO3--	.01	-6.731	MGHCO3+	.00	-7.880	AL+++	.00	-19.273
H2S	.02	-6.261	CAOH+	.49	-5.062	AL0H++	.00	-13.566
HS-	.08	-5.622	MGOH+	.00	-6.975	AL(OH)2+	.00	-8.495
S--	.00	-13.746	NH4OH	.00	.000	AL(OH)3	.24	-5.517
H2SO4	.00	-15.413	NH4+	.00	.000	AL(OH)4-	.19	-5.632
HSO4-	.02	-6.704	FE++	.01	-6.936	ALSO4+	.00	-13.313
SO4--	120.94	-2.900	FE+++	.00	-22.611	AL(SO4)2-	.00	-20.360
HF	.00	-7.499	FE0H+	.02	-6.604	ALF++	.00	-16.283
F-	.55	-4.540	FE(OH)2	.00	-7.549	ALF2+	.00	-14.776
CL-	1312.04	-1.432	FE(OH)3-	.00	-9.460	ALF3	.00	-15.001
NA+	587.65	-1.592	FE(OH)4--	.00	-14.389	ALF4-	.00	-16.902
K+	14.20	-3.440	FE(OH)++	.00	-15.525	ALF5--	.00	-19.836
CA++	282.50	-2.152	FE(OH)2+	.00	-10.040	ALF6---	.00	-23.820

IONIC STRENGTH = .04827

IONIC BALANCE :

CATIONS (MOL.EQ.) .04013072

ANIONS (MOL.EQ.) .03978148

DIFFERENCE (%) .87

OXIDATION POTENTIAL (VOLTS) :

BH H2S= -.430

BH CH4= 99.999

BH H2= 99.999

BH NH3= 99.999

LOG SOLUBILITY PRODUCTS OF MINERALS IN DEEP WATER

	TEOR.	CALC.		TEOR.	CALC.		TEOR.	CALC.
ADULARIA	-15.664	-17.464	ALBITE LOW	-15.067	-15.609	ANALCIME	-12.168	-12.779
ANHYDRITE	-6.366	-5.930	CALCITE	-10.399	-9.748	CHALCEDONY	-2.432	-2.830
MG-CHLORITE	-80.047	-82.572	FLUORITE	-10.561	-11.888	GORTHITE	-2.669	-1.857
LAUMONTITE	-25.402	-24.705	MICROCLINE	-16.731	-17.464	MAGNETITE	-25.789	-19.991
CA-MONTMOR.	-76.836	-86.991	K-MONTMOR.	-36.807	-45.776	MG-MONTMOR.	-78.177	-89.935
NA-MONTMOR.	-36.953	-43.922	MUSCOVITE	-18.991	-19.354	PRRHNITE	-35.828	-33.569
PYRRHOTITE	-74.605	-76.763	PYRITE	-111.470	-103.199	QUARTZ	-2.683	-2.830
WAIKAKITE	-23.901	-24.705	WOLLASTONITE	9.469	8.947	ZOISITE	-35.407	-34.314
EPIDOTE	-39.779	-35.225	MARCASITE	-90.765	-103.199			

Table 7. (continued)

UNU Geothermal Training Programme
SAMPLE = 073

DEEP WATER COOLED TO 100.0 DEGREES C.

ACTIVITY COEFFICIENTS IN DEEP WATER

H+	.831	KSO4-	.802	FE++	.432	FECL+	.788
OH-	.783	F-	.783	FE+++	.196	AL+++	.196
H3STO4-	.788	CL-	.777	FROH+	.798	ALOH++	.421
H2STO4--	.421	NA+	.788	FE(OH)3-	.798	AL(OH)2+	.802
H2BO3-	.771	K+	.777	FE(OH)4--	.413	AL(OH)4-	.794
HCO3-	.788	CA++	.432	FROH++	.413	ALSO4+	.794
CO3--	.403	MG++	.467	FE(OH)2+	.802	AL(SO4)2-	.794
HS-	.783	CAHCO3+	.808	FE(OH)4-	.802	ALF++	.421
S--	.413	MGHCO3+	.788	FESO4+	.798	ALF2+	.802
HSO4-	.794	CAOH+	.808	FECL++	.413	ALF4-	.794
SO4--	.392	MGOH+	.812	FECL2+	.798	ALF5--	.403
NASO4-	.802	NH4+	.771	FECL4-	.788	ALF6---	.130

CHEMICAL COMPONENTS IN DEEP WATER (PPM AND LOG MOLE)

H+ (ACT.)	.00	-7.432	MG++	.21	-5.067	FE(OH)3	.00	-8.892
OH-	.32	-4.721	NACL	7.51	-3.891	FE(OH)4-	.00	-8.473
H4STO4	141.41	-2.832	KCL	.05	-6.148	FECL+	.01	-7.234
H3STO4-	5.15	-4.267	NASO4-	8.11	-4.167	FECL2	.00	-19.209
H2STO4--	.00	-7.553	KSO4-	.60	-5.352	FECL++	.00	-21.305
NAH3STO4	1.29	-4.963	CASO4	86.75	-3.196	FECL2+	.00	-22.442
H3BO3	.00	.000	MGSO4	.35	-5.538	FRCL3	.00	-24.709
H2BO3-	.00	.000	CACO3	.89	-5.052	FECL4-	.00	-27.509
H2CO3	.83	-4.873	MGCO3	.00	-8.480	FESO4	.01	-7.336
HCO3-	11.04	-3.743	CAHCO3+	4.66	-4.336	FESO4+	.00	-20.777
CO3--	.04	-6.151	MGHCO3+	.00	-7.813	AL+++	.00	-17.392
H2S	.01	-6.528	CAOH+	.13	-5.650	ALOH++	.00	-12.723
HS-	.09	-5.579	MGOH+	.00	-7.544	AL(OH)2+	.00	-8.604
S--	.00	-13.948	NH4OH	.00	.000	AL(OH)3	.09	-5.933
H2SO4	.00	-17.001	NH4+	.00	.000	AL(OH)4-	.37	-5.408
HSO4-	.00	-7.547	FE++	.04	-6.162	ALSO4+	.00	-17.658
SO4--	153.23	-2.797	FE+++	.00	-22.382	AL(SO4)2-	.00	-18.758
HF	.00	-8.239	FROH+	.03	-6.423	ALF++	.00	-14.923
F-	.55	-4.539	FE(OH)2	.00	-8.136	ALF2+	.00	-13.840
CL-	1316.31	-1.430	FE(OH)3-	.00	-11.080	ALF3	.00	-14.326
NA+	590.83	-1.590	FE(OH)4--	.00	-16.247	ALF4-	.00	-16.386
K+	14.30	-3.437	FE(OH)++	.00	-16.031	ALF5--	.00	-13.333
CA++	296.87	-2.130	FE(OH)2+	.00	-11.476	ALF6---	.00	-23.236

IONIC STRENGTH = .04979

IONIC BALANCE : CATIONS (MOL.EQ.) .04094628
ANIONS (MOL.EQ.) .04061263
DIFFERENCE (%) .82

OXIDATION POTENTIAL (VOLTS) : BH H2S= -.365 BH CH4= 99.999 BH H2= 99.999 BH NH3= 99.999

LOG SOLUBILITY PRODUCTS OF MINERALS IN DEEP WATER

	TEOR.	CALC.		TEOR.	CALC.		TEOR.	CALC.
ADULARIA	-17.222	-17.438	ALBITE LOW	-16.485	-15.585	ANALCINE	-13.199	-12.753
ANHYDRITE	-5.607	-5.699	CALCITE	-9.438	-9.041	CHALCEDONY	-2.841	-2.852
MG-CHLORITE	-80.302	-84.892	FLUORITE	-10.538	-11.786	GORTHITR	-4.674	-3.741
LAUMONTITE	-27.182	-24.614	MICROCLINE	-18.598	-17.438	MAGNETITE	-29.703	-23.664
CA-MONTMOR.	-85.559	-82.406	K-MONTMOR.	-41.648	-43.502	MG-MONTMOR.	-86.708	-85.366
NA-MONTMOR.	-41.660	-41.649	MUSCOVITE	-21.151	-18.573	PREHNITE	-36.778	-33.931
PYRRHOTITE	-99.806	-95.133	PYRITE	-147.012	-121.401	QUARTZ	-3.098	-2.832
WATRAKITE	-24.710	-24.614	WOLLASTONITE	10.824	9.536	ZOISITE	-35.833	-34.433
EPIDOTE	-43.603	-37.672	MARCASITE	-123.575	-121.401			

Table 7. (continued)

UNU Geothermal Training Programme
SAMPLE = 073

DEEP WATER COOLED TO 50.0 DEGREES C.

ACTIVITY COEFFICIENTS IN DEEP WATER

H+	.846	HSO4-	.820	FE++	.469	FECL+	.807
OH-	.802	F-	.802	FE+++	.230	AL+++	.230
H3SIO4-	.807	CL-	.797	FEOH+	.816	ALOH++	.458
H2SIO4--	.458	NA+	.807	FE(OH)3-	.816	AL(OH)2+	.820
H2BO3-	.791	K+	.797	FE(OH)4--	.450	AL(OH)4-	.812
HCO3-	.807	CA++	.469	FEOH++	.450	ALSO4+	.812
CO3--	.441	MG++	.503	FE(OH)2+	.820	AL(SO4)2-	.812
HS-	.802	CAHCO3+	.825	FE(OH)4-	.820	ALF++	.458
S--	.450	MGHCO3+	.807	FESO4+	.816	ALF2+	.820
HSO4-	.812	CAOH+	.825	FECL++	.450	ALF4-	.812
SO4--	.430	MGOH+	.829	FECL2+	.816	ALF5--	.441
NASO4-	.820	NH4+	.791	FECLA-	.807	ALF6---	.159

CHEMICAL COMPONENTS IN DEEP WATER (PPM AND LOG MOLE)

H+ (ACT.)	.00	-7.964	MG++	.24	-4.998	FE(OH)3	.00	-11.329
OH-	.11	-5.207	NACL	1.99	-4.467	FE(OH)4-	.00	-11.090
H4SIO4	141.31	-2.833	KCL	.03	-6.441	FECL+	.00	-7.746
H3SIO4-	5.14	-4.267	NASO4-	5.56	-4.331	FECL2	.00	-24.220
H2SIO4--	.00	-7.425	RSO4-	.38	-5.549	FECL++	.00	-22.897
NAH3SIO4	1.41	-4.924	CASO4	56.59	-3.381	FECL2+	.00	-24.025
H3RO3	.00	.000	MGSO4	.17	-5.839	FRCL3	.00	-26.506
H2BO3-	.00	.000	CACO3	.95	-5.024	FECL4-	.00	-29.654
H2CO3	.25	-5.398	MGCO3	.00	-8.159	FESO4	.01	-7.198
HCO3-	13.27	-3.663	CAHCO3+	1.68	-4.778	FESO4+	.00	-22.147
CO3--	.14	-5.624	MGHCO3+	.00	-7.812	AL+++	.00	-15.944
H2S	.01	-6.810	CAOH+	.03	-6.333	ALOH++	.00	-12.179
HS-	.09	-5.556	MGOH+	.00	-8.176	AL(OH)2+	.00	-8.867
S--	.00	-14.031	NH4OH	.00	.000	AL(OH)3	.03	-6.382
H2SO4	.00	-19.284	NH4+	.00	.000	AL(OH)4-	.44	-5.331
HSO4-	.00	-8.663	FE++	.05	-6.014	ALSO4+	.00	-16.428
SO4--	176.86	-2.735	FE+++	.00	-23.033	AL(SO4)2-	.00	-17.584
HF	.00	-9.215	FEOH+	.01	-6.879	ALF++	.00	-13.810
F-	.55	-4.539	FE(OH)2	.00	-9.394	ALF2+	.00	-13.003
CL-	1319.67	-1.429	FE(OH)3-	.00	-13.455	ALF3	.00	-13.641
NA+	593.47	-1.588	FE(OH)4--	.00	-18.972	ALF4-	.00	-15.792
K+	14.37	-3.435	FE(OH)++	.00	-17.205	ALF5--	.00	-18.826
CA++	306.98	-2.116	FE(OH)2+	.00	-13.433	ALF6---	.00	-22.607

IONIC STRENGTH = .05089

IONIC BALANCE : CATIONS (MOL.EQ.) .04153923
ANIONS (MOL.EQ.) .04122734
DIFFERENCE (%) .75

OXIDATION POTENTIAL (VOLTS) : BH H2S= -.311 BH CH4= 99.999 BH H2= 99.999 BH NH3= 99.399

LOG SOLUBILITY PRODUCTS OF MINERALS IN DEEP WATER

	TEOR.	CALC.		TEOR.	CALC.		TEOR.	CALC.
ADULARIA	-19.664	-17.416	ALBITE LOW	-18.715	-15.564	ANALCIME	-14.863	-12.731
ANHYDRITE	-4.966	-5.545	CALCITE	-8.720	-8.424	CHALCEDONY	-3.297	-2.833
MG-CHLORITE	-82.885	-88.171	FLUORITE	-10.737	-11.714	GORTHITP	-6.429	-5.874
LAUNONTITE	-30.213	-24.545	NICROCLINE	-21.446	-17.416	MAGNETITE	-33.553	-28.636
CA-MONTMOR.	-101.366	-76.521	K-MONTMOR.	-50.171	-40.571	MG-MONTMOR.	-102.214	-79.373
NA-MONTMOR.	-49.974	-38.719	MUSCOVITE	-25.002	-17.581	PREHNITE	-39.270	-34.762
PYRRHOTITE	-125.035	-115.786	PYRITE	-185.595	-141.638	QUARTZ	-3.641	-2.833
WAIKAKITE	-26.535	-24.545	WOLLASTONITE	12.603	10.650	ZOISITE	-37.927	-34.844
EPIDOTE	-47.739	-40.636	MARCASITE	-158.572	-141.638			

Table 7. (continued)

UNU Geothermal Training Programme
 SAMPLE = 073

DEEP WATER COOLED TO .0 DEGREES C.

ACTIVITY COEFFICIENTS IN DEEP WATER

H+	.850	KSO4-	.825	FB++	.481	FECL+	.813
OH-	.808	F-	.808	FB+++	.241	AL+++	.241
H3SIO4-	.813	CL-	.803	FEOH+	.822	ALOH++	.470
H2SIO4--	.470	NA+	.813	FB(OH)3-	.822	AL(OH)2+	.825
H2BO3-	.798	K+	.803	FE(OH)4--	.463	AL(OH)4-	.818
HCO3-	.813	CA++	.481	FEOH++	.463	ALSO4+	.818
CO3--	.453	MG++	.514	FB(OH)2+	.825	AL(SO4)2-	.818
HS-	.808	CAHCO3+	.830	FB(OH)4-	.825	ALF++	.470
S--	.463	MGHCO3+	.813	FESO4+	.822	ALF2+	.825
HSO4-	.818	CAOH+	.830	FECL++	.463	ALF4-	.818
SO4--	.443	MGOH+	.834	FECL2+	.822	ALF5--	.453
NASO4-	.825	NH4+	.798	FECL4-	.813	ALF6---	.169

CHEMICAL COMPONENTS IN DEEP WATER (PPM AND LOG MOLE)

H+ (ACT.)	.00	-8.906	MG++	.27	-4.957	FB(OH)3	.00	-14.103
OH-	.02	-5.927	NACL	.87	-4.828	FB(OH)4-	.00	-14.042
H4SIO4	141.73	-2.831	KCL	.01	-6.735	FECL+	.00	-8.615
H3SIO4-	4.64	-4.312	NASO4-	3.31	-4.556	FECL2	.00	-31.167
H2SIO4--	.00	-7.417	KSO4-	.20	-5.825	FECL++	.00	-25.163
NAH3SIO4	1.52	-4.890	CASO4	38.34	-3.550	FECL2+	.00	-26.287
H3RO3	.00	.000	MGSO4	.05	-6.342	FECL3	.00	-29.047
H2RO3-	.00	.000	CACO3	1.04	-4.984	FECL4-	.00	-32.637
H2CO3	.05	-6.081	MGCO3	.00	-7.839	FESO4	.01	-7.161
HCO3-	13.86	-3.644	CAHCO3+	.40	-5.402	FESO4+	.00	-24.148
CO3--	.46	-5.120	MGHCO3+	.00	-7.889	AL+++	.00	-14.500
H2S	.00	-7.129	CAOH+	.00	-7.247	ALOH++	.00	-11.561
HS-	.09	-5.544	MGOH+	.00	-9.171	AL(OH)2+	.00	-8.998
S--	.00	-13.915	NH4OH	.00	.000	AL(OH)3	.02	-6.615
H2SO4	.00	-22.589	NH4+	.00	.000	AL(OH)4-	.46	-5.316
HSO4-	.00	-10.158	FB++	.06	-5.966	ALSO4+	.00	-15.202
SO4--	191.77	-2.700	FB+++	.00	-24.334	AL(SO4)2-	.00	-16.403
HF	.00	-10.542	FEOH+	.00	-7.524	ALF++	.00	-12.473
F-	.55	-4.539	FE(OH)2	.00	-11.000	ALF2+	.00	-11.772
CL-	1320.35	-1.429	FB(OH)3-	.00	-16.412	ALF3	.00	-12.426
NA+	594.32	-1.588	FB(OH)4--	.00	-22.509	ALF4-	.00	-14.579
K+	14.43	-3.433	FB(OH)++	.00	-18.848	ALF5--	.00	-17.574
CA++	312.83	-2.108	FB(OH)2+	.00	-15.774	ALF6---	.00	-21.180

IONIC STRENGTH = .05153

IONIC BALANCE : CATIONS (MOL.EQ.) .04185931
 ANIONS (MOL.EQ.) .04156729
 DIFFERENCE (%) .70

OXIDATION POTENTIAL (VOLTS) : BH H2S= -.269 BH CH4= 99.999 BH H2= 99.999 BH NH3= 99.999

LOG SOLUBILITY PRODUCTS OF MINERALS IN DEEP WATER

	TEOR.	CALC.		TEOR.	CALC.		TEOR.	CALC.
ADULARIA	-23.276	-17.404	ALBITE LOW	-22.008	-15.553	ANALCINE	-17.390	-12.722
ANHYDRITE	-4.511	-5.478	CALCITE	-8.378	-7.889	CHALCEDONY	-3.321	-2.831
MG-CHLORITE	-88.387	-93.644	FLUORITE	-11.323	-11.688	GORTHITE	-8.225	-8.106
LAUMONTITE	-34.804	-24.514	MICROCLINE	-25.573	-17.404	MAGNETITE	-38.146	-34.535
CA-MONTHOR.	-121.263	-67.824	K-MONTHOR.	-61.001	-36.227	MG-MONTHOR.	-121.654	-70.645
NA-MONTHOR.	-60.441	-34.377	MUSCOVITE	-30.074	-16.128	PREHNITE	-43.653	-36.147
PYRRHOTITE	-152.446	-137.996	PYRITE	-231.987	-162.799	QUARTZ	-4.382	-2.831
WAIKAKITE	-29.576	-24.514	WOLLASTONITE	15.039	12.556	ZOISITE	-41.763	-35.516
EPIDOTE	-54.230	-44.253	MARCASITE	-200.020	-162.799			

Table 8. Summary of results of mineral equilibria calculations using the WATCH1 and WATCH3 computer programs

Sample No.	Locality	Measured temp. °C	Chalced temp. °C	Quartz temp. °C	NaK temp. °C	Range equil. temp. °C	Cluster equil. temp. °C	Best equil. temp. °C
08	Svartsengi,4	242	na	234	244	214 - >250	214 - 250	214
09	Reykjanes,8	274 - 292	na	247	259	194 - 281	201 - 272	246
28	Reykholt,1	133	134	156	135	85 - >150	131 - >150	135
30	Reykjabol,1	152	157	178	151	106 - >200	123 - 188	126/161
32	Hveragerdi,4	181	169	190	187	104 - 234	121 - 166	166
48	Krafla,9	240	na	246	234	142 - >250	219 - 250	248
49	Krafla,7	206 - 325	na	265	251	134 - >350	217 - 301	275
51	Namafjall,4	180	na	214	251	113 - 288	138 - 215	185
52	Namafjall,8	160(est.)	na	220	247	87 - 300	141 - 226	160
60	Isafjardard,sp.	84?	85	110	82	10 - >150	57 - 92	90
43 72	Seltj.,2, 80m	31	59	85	52	62 - 119	62 - 97	94
73	Seltj.,2, 570m	78	101	125	83	92 - 166	92 - 134	134
74	Seltj.,2, 725m	77	98	122	69	77 - >200	98 - 138	138
75	Borgarfjord.,cold	3.8	25	51	112	10 - >40	10 - 29	none
76	England, hot	91	128	150	104	59 - 167	81 - 117	82
79	Urridavatn,8	77	60	86	66	39 - >200	49 - 87	84
80	Urridavatn,4	64	50	76	58	38 - >200	43 - 88	42
81	IRD P	47/78	62	87	52	2 - >100	44 - 67	65

Note: ? denotes temperatures measured at the surface

na denotes "not applicable"

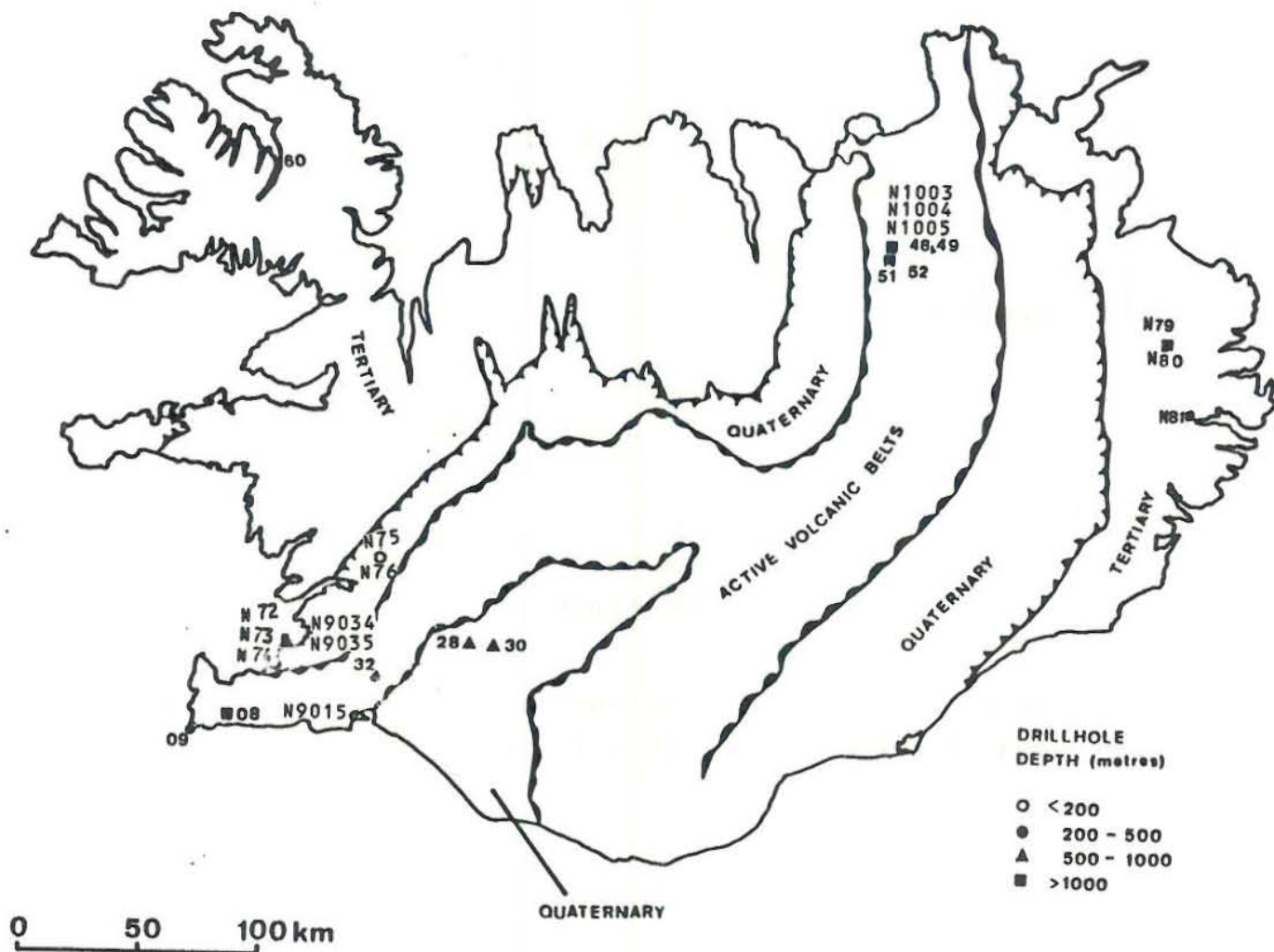


Figure 1. Map of Iceland showing localities from which samples were collected, analysed or interpreted. Samples designated with N (e.g. N73) were collected by the National Energy Authority in 1988. Other samples are from Arnorsson et al (1983a).

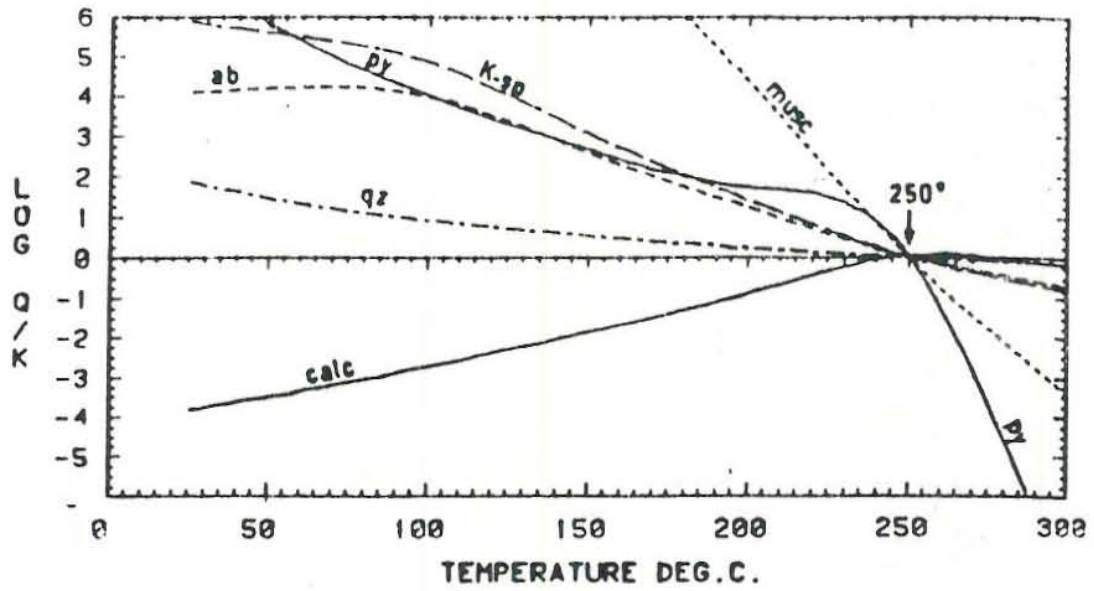


Figure 2. Plot of $\text{Log}(Q/K)$ versus temperature for a geothermal fluid at equilibrium at 250°C (Reed and Spycher, 1984).

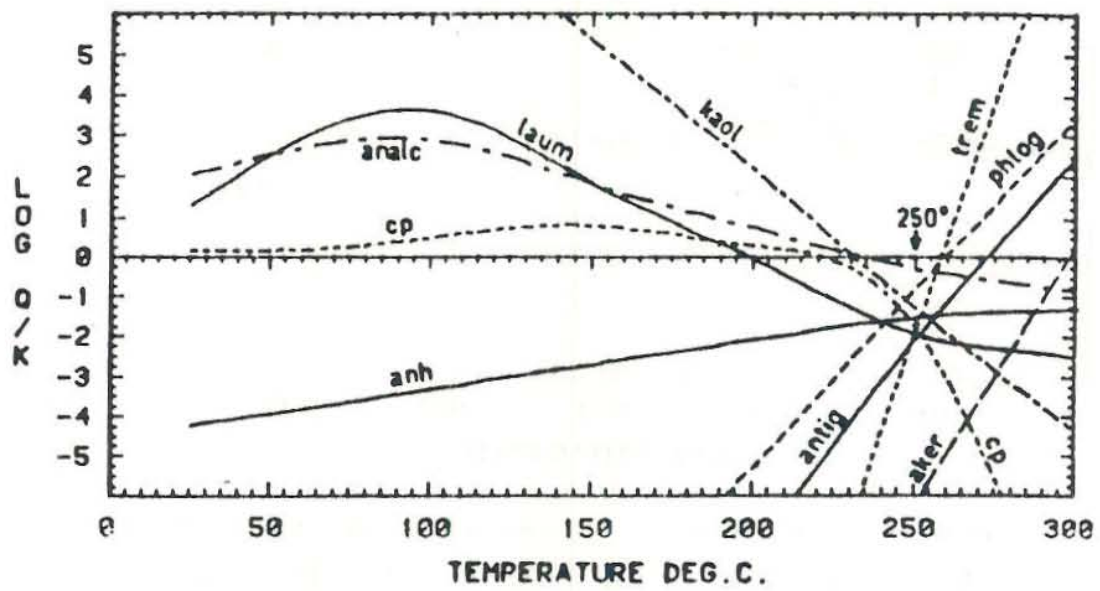


Figure 3. Plot of $\text{Log}(Q/K)$ versus temperature for a geothermal fluid mixed with highly dilute water (Reed and Spycher, 1984).

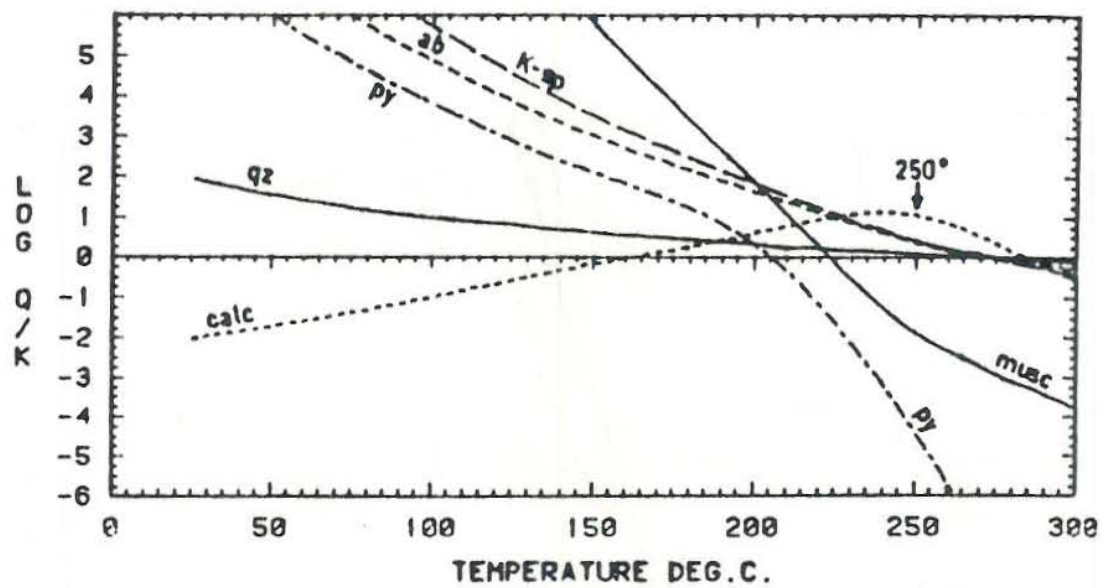


Figure 4. Plot of $\text{Log}(Q/K)$ versus temperature for a geothermal fluid that has boiled (Reed and Spycher, 1984).

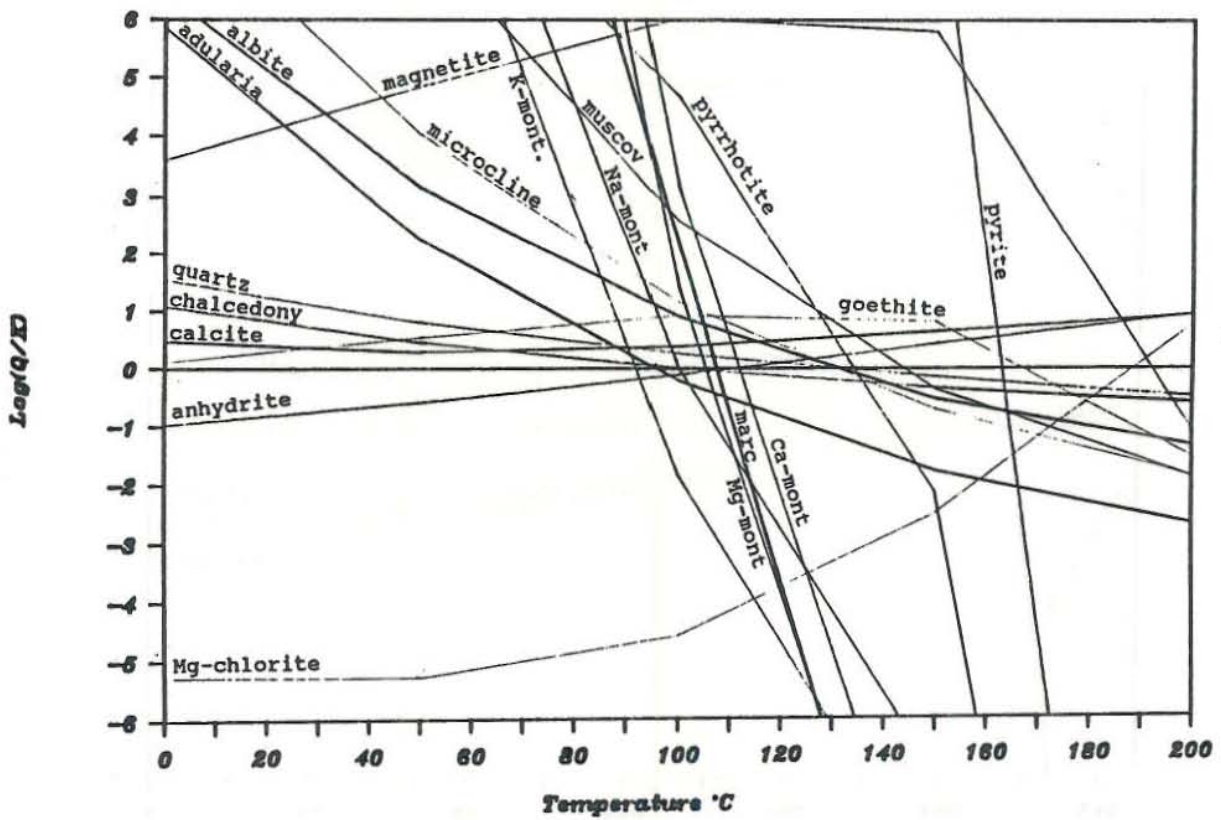


Figure 5. Mineral equilibrium diagram for sample from Seljarnarnes well 2 at 570m depth. All likely primary and secondary minerals have been included. Calculated on WATCH3.

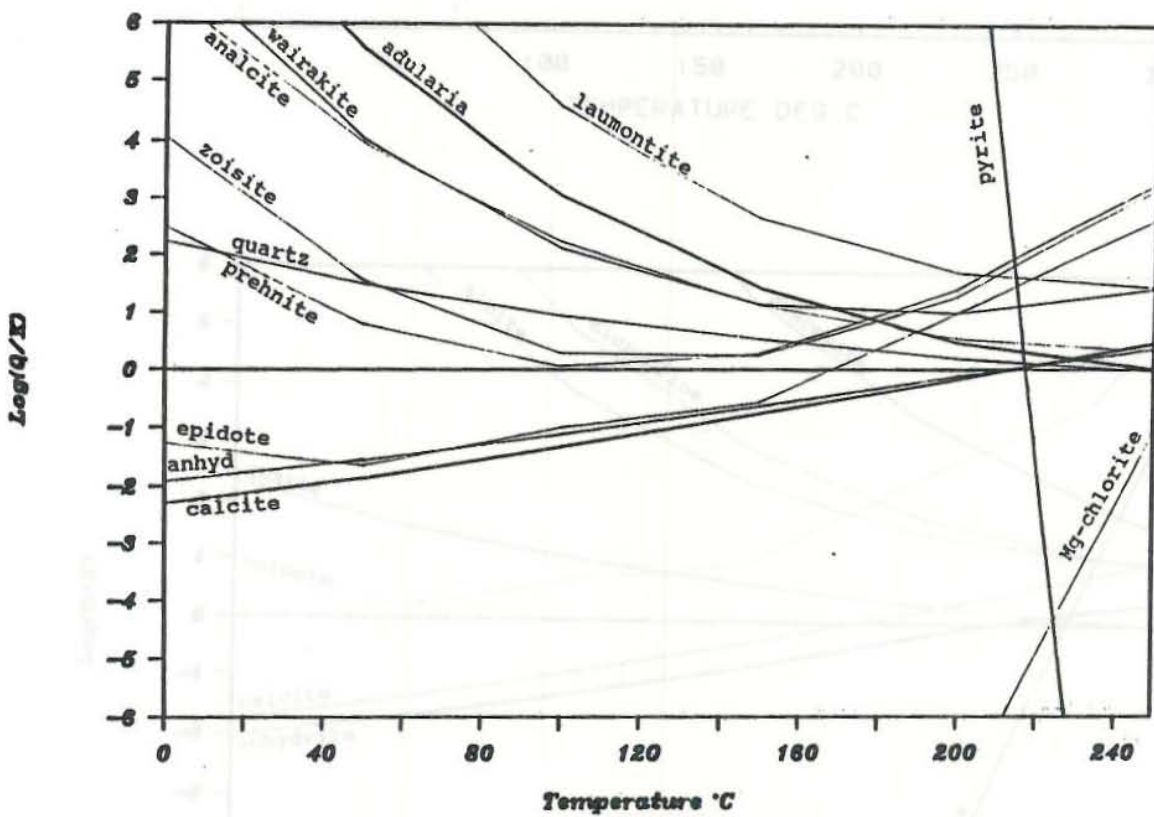


Figure 6. Mineral equilibrium diagram for Svartsengi Well 4. All likely primary and secondary minerals have been included. Calculated on WATCH1.

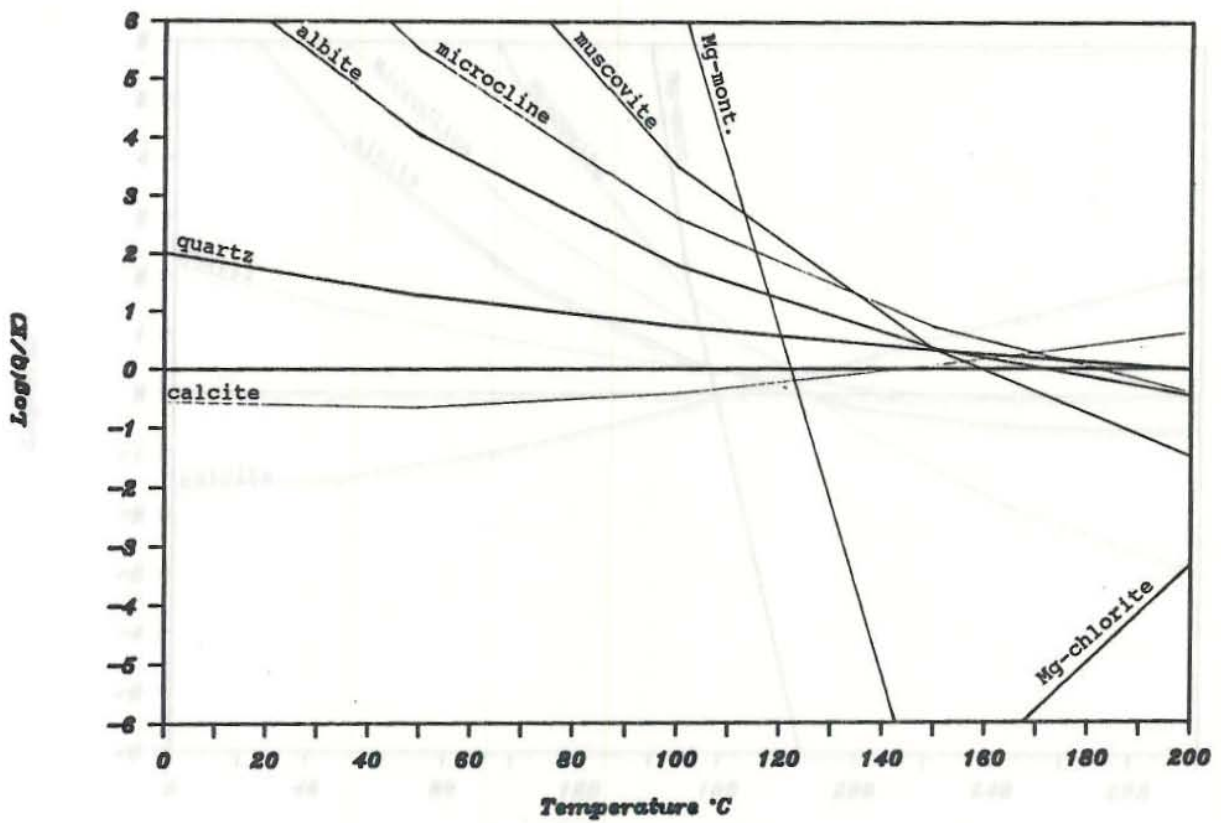
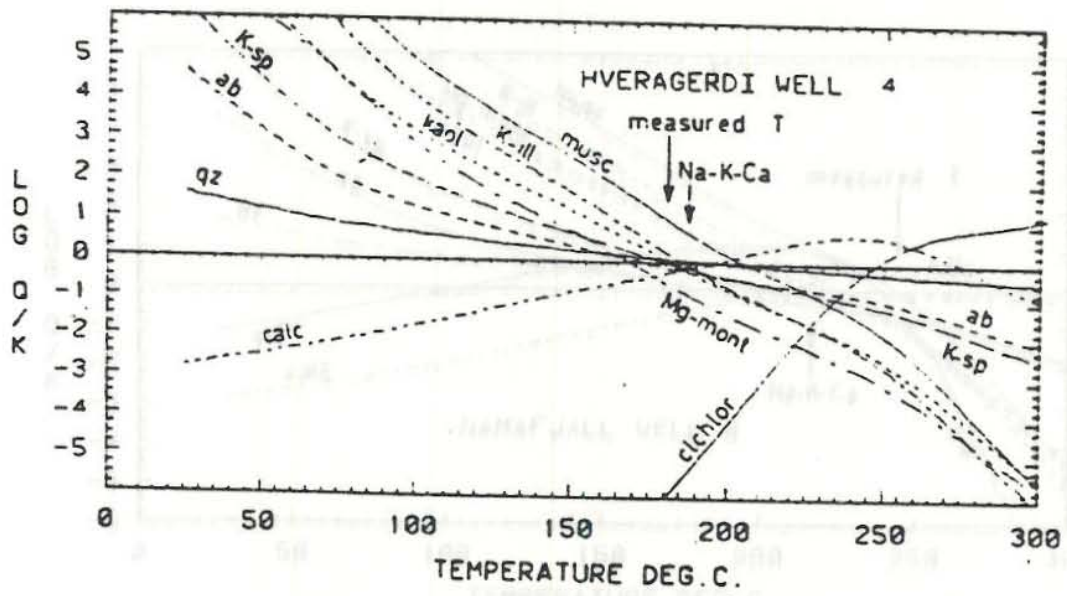


Figure 8. Comparison between mineral equilibria diagrams from the SOLVEQ (upper figure) and WATCH (lower figure) programs for Hveragerdi well 4.