

IMPLICATIONS OF SOLUTION-MINERAL EQUILIBRIA
ON THE EXPLOITATION OF THE S-NEGROS
GEOTHERMAL FIELD, PHILIPPINES

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ABSTRACT

Representative chemical data from production boreholes in the Southern Negros Geothermal Field, Philippines, were studied to assist in the evaluation of certain solute-mineral equilibria and the calibration and selection of the most applicable chemical geothermometer. Anhydrite and quartz are shown to be in equilibrium with the deep fluid, while calcite saturation seems to be governed by temperature and mixing with cooler fluid. A computer simulation of boiling and cooling processes suggests no potential calcite deposition, whereas problems of amorphous silica scaling could occur at a re-injection or production separation temperature of 150 °C and lower. Empirically derived relation of aqueous acid molecules and cation to proton ratios with temperatures, as obtained using chemical data in Iceland, fits well for the aqueous H_4SiO_4 and H_2CO_3 species in fluids of Southern Negros, but gives significant deviations for H_2S , H_2SO_4 , and the cation to proton ratios. The deviations are attributed to the mechanisms of boiling and degassing, a different rock leaching property, and mixing with cold water. Calibration of various chemical geothermometers using measured temperatures at main production zone as reference, shows that the pressure sensitive silica geothermometer of Ragnarsdottir and Walther (1982) and the Na/K geothermometer of Arnorsson, et al. 1982 c) using computed activities of Na and K yield the least mean and standard deviations. Applying the Chloride-enthalpy mixing models of Fournier (1977), suggests cold meteoric water recharge from the northwest, picking up heat from the rock before mixing with the outflowing hydrothermal fluid within the vicinity of Sogongon and Nasuji. True reservoir temperatures are estimated to be at least 300 °C.

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1 INTRODUCTION

1.1 Scope and objective of work

The Philippines is one of the many countries presently engaged in an accelerated development programme to harness geothermal energy as an indigenous power source. The revised energy development strategy calls for an installed 1,774 MW of electricity at the end of 1985. Contributory to this ambitious plan is the availment of foreign expertise and assistance, as well as the foreign training of selected Filipinos in the leading countries engaged in geothermal exploration, development, and utilization. In line with this, the author was awarded a United Nations University Fellowship to attend the 1982 UNU Geothermal Training Programme held at the National Energy Authority in Reykjavik, Iceland. He attended the specialized course in the geochemistry of thermal fluids.

After a five week introductory lecture series covering the various scientific, engineering, and economic aspects in geothermal exploration and exploitation, the author received specialized training in the following topics: water and steam sampling from natural thermal manifestations and production boreholes (3 weeks); analytical methods and procedures (3 weeks); rock-fluid chemistry and chemical thermodynamics (2 weeks); computer compilation of analytical results (2 weeks); and computer approach to the interpretation of geothermal well discharges (2 weeks). This report represents a research project carried out during the last month of the training programme.

Chemical data from production boreholes in the Southern Negros geothermal field, Philippines, were studied to assist in their overall evaluation. Since the study of aquifer fluid chemistry involves lengthy and complex equations, calculations, and especially iterative procedures, a computer programme developed at the University of Iceland and the National Energy Authority, was used (Arnorsson, et al. 1982 a). This topic was selected

to train the author in such a specialized field of studies commonly used both during the exploration and the exploitation stages in geothermal development.

1.2 Methods of collection and analysis

All analytical procedures were performed at the geochemical laboratories of PNO-C-EDC Geothermal Division at Manila and Southern Negros. The following is a brief description of the methods used in water and steam sampling.

Figure 1 shows the typical discharge test assembly of production boreholes in the Southern Negros geothermal field. The two-phase fluid from the well head is by-passed to a horizontal discharge pipe and allowed to discharge to atmospheric pressure through a twin-tower cyclone silencer equipped with a rectangular weir and a trapezoidal cippoletti weir to measure water flow rate. Two sampling probes are located on the horizontal by-pass pipe consisting of either a 3/4" or 1/2" inside diameter nipple and a gate valve to control fluid flow during sampling. Generally, water samples were collected at atmospheric pressure and local boiling point (i.e. in the weirbox), and at the low pressure probe.

Water samples could not be collected at the high pressure point since it was experienced that steam carry over with the liquid phase could not be avoided, probably due to the proximity of this probe to the T-junction of the wellhead, where turbulent flow of both water and steam phases makes effective separation impossible. A weber cyclone separator was used to carry out the steam and water sampling under pressure. Sampling pressures ranged from 2-12 bars absolute. The difference between fluid pressure within the pipe and the sampling pressure were kept to a minimum to reduce the possibility of steam flashing.

Water sample treatment consists of acidification to pH 2-3 for the analysis of silica and the main cations and storing in polyethylene bottles; untreated samples were analysed for boron, chloride, and sulfate; water samples collected into air-tight

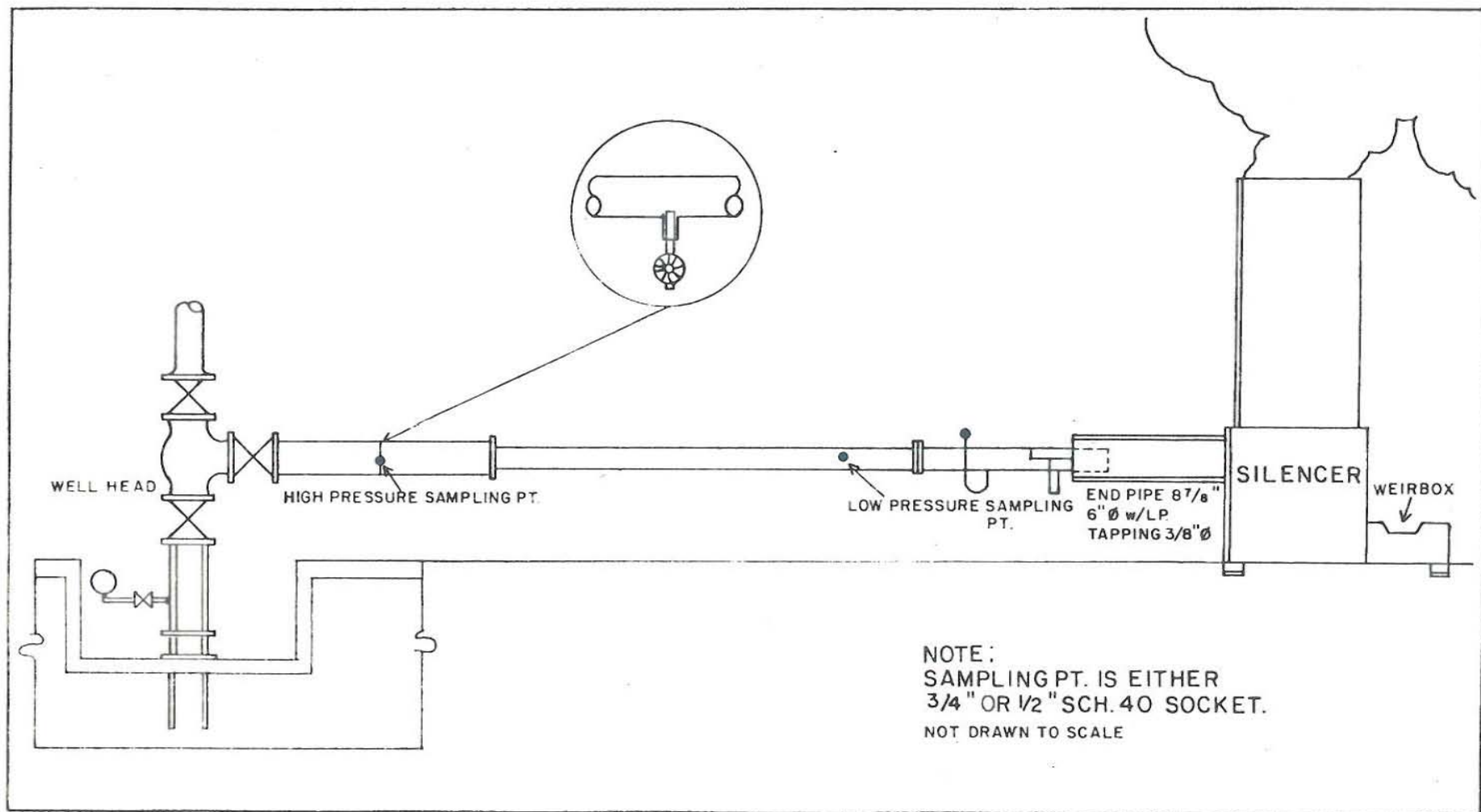


FIG. 1. DISCHARGE TEST ASSEMBLY IN THE PALINPINON FIELD, SOUTHERN NEGROS

(modified from Baltazar, 1980)

F. 20034

TABLE 1

Analytical methods for steam and water samples

| | Element | Method |
|------------------|-------------------------|--|
| Steam phase | CO ₂ | Potentiometric titration from pH 8.25-3.80 using standard 0.1N hydrochloric acid. |
| | H ₂ S | Back titration with 0.01N sodium thiosulfate in acid solution(+2N sulfuric acid) after addition of 0.01N iodine solution |
| | NH ₃ | Specific ion sensitive electrode |
| | Cl | Mohr titration using 0.01N silver nitrate with potassium chromate-dichromate buffer indicator |
| Water phase | pH | Glass electrode |
| | Li | Atomic absorption spectrophotometry |
| | Na | Atomic absorption spectrophotometry |
| | K | Atomic absorption spectrophotometry |
| | Ca | Atomic absorption spectrophotometry |
| | Mg | Atomic absorption spectrophotometry |
| | Rb | Atomic absorption spectrophotometry |
| | Cs | Atomic absorption spectrophotometry |
| | Cl | Mohr titration using 0.1N silver nitrate standard solution with potassium chromate/dichromate buffer indicator |
| | SO ₄ | Ultra-violet spectrophotometry by liberation of chromate ions through pptn. of sulfate using barium chromate |
| | B | Potentiometric titration with 0.025N sodium hydroxide to pH 7.30 after addition of excess mannitol |
| | SiO ₂ | Ultra violet spectrophotometry through colorimetric reaction using ammonium molybdate |
| | CO ₂ | Double potentiometric titration from pH 8.25-3.80 using 0.02N hydrochloric acid |
| H ₂ S | Same as for steam phase | |

glass bottles with a 6-8" long butyl rubber tubing equipped with aluminum clips (air free samples) were taken mainly for the analysis of pH, total sulfide, and total carbonate. Steam samples were collected both at the high and low pressure points, into evacuated 2-3 liters boro-silicate flasks containing a measured volume of 70 % sodium-hydroxide solution to bring the CO₂ and H₂S gases of the steam phase into solution. Separate untreated steam condensates were similarly collected into 3-liters evacuated flasks for ammonia analysis. All analytical procedures and methods used are summarized in Table 1. Relevant bore output data such as discharge enthalpies and total mass flow rates were also measured along with the time of sampling, using the critical lip pressure method developed by James (1962).

2 THE SOUTHERN NEGROS GEOTHERMAL FIELD

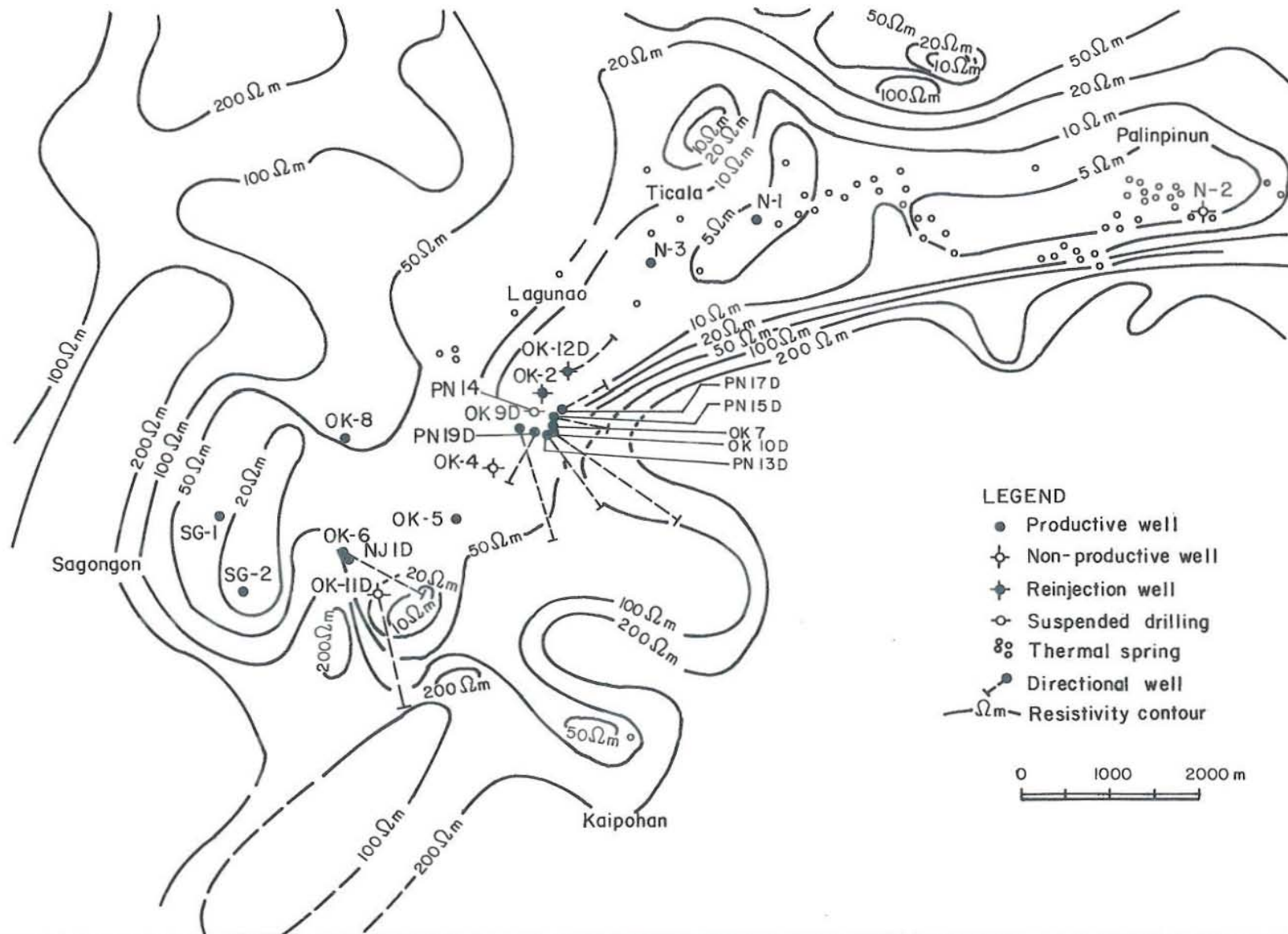
2.1 Introduction

2.1.1 Regional geological setting

The Southern Negros Geothermal Field (SNGF) covers almost one half of the peninsular arm of Negros Island, and encompasses a total reservation area of approximately 133,000 hectares (Fig. 2). The SNGF is situated within a localized graben that forms the Okoy river-valley, located in between the merging flanks of two dormant andesitic volcanoes, namely the Cuernos de Negros to the south, and mount Balinsasayao to the north. The whole field is dissected by a series of NNW trending right lateral faults, NE trending left lateral faults, and a system of step faults striking NW and NE. The first two fault patterns are believed to be sympathetic to the Philippine fault, while the third pattern could be due to a failure enhanced by local coupling stresses that could have been initiated by lateral movements along shear planes (Bagamasbad, 1979; Reyes and Tolentino, 1981).

The SNGF consists of two promising thermal areas, namely the Palinpinon field, located on the eastern section of the Okoy river valley, and the Baslay de Dauin area, situated within a ring plain southeast of the Cuernos volcano. The distribution of thermal springs and resistivity contours are shown in Fig. 3. These manifestations, in the form of numerous hot and warm springs, alteration patches, and steaming vents, are broadly contained within the two thermal areas, but are more numerous and intense along the 9 km length of the Okoy river, the southern tributaries of which drain the Northern flank of the large Plio-Pleistocene composite volcano of Cuernos de Negros. Dipole-dipole and Schlumberger resistivity surveys with electrode spacings of AB/2 equals 250 and 500 m identified four distinct anomalies (5, 10, 20, and 50 ohm-meter), all

Figure 3 Map of Palinpinon field showing resistivity contours, distribution of thermal springs, and location of boreholes



converging towards the Cuernos volcano, suggesting that this may be the ultimate heat source of the field (KRTA, 1977).

2.1.2 Brief status of geothermal development.

The Palinpinon field has been given priority over Baslay de Dauin in exploration and production drilling due to the more impressive thermal areas and encouraging geological, geophysical, and geochemical results. In April, 1982 a total of 26 wells had been drilled (vertical and directional), while four others are currently being drilled. Due to the steep topography in Puhagan which severely limits the number of suitable sites for vertical wells, directional drilling has become common practice. Thus all targets can be reached from drillsites close to the location of the power station. These wells penetrate a series of intermediate volcanics and volcanoclastics before encountering a fossiliferous sedimentary formation and/or a diorite-monzonite pluton (Reyes and Tolentino, 1981). Fig. 4 shows the main structures, the location of all wellheads, and the bottomholes of the directionally drilled production wells that will be considered in this report.

2.2 Review of previous geochemical studies.

The chemistry of surface thermal manifestations and early exploration wells has been described by previous workers (Glover, 1975; Camales, 1979; Ruaya, 1980). Production borehole geochemistry studies were described by Jordan (1981 a,b,c) and Clemente (1982 a,b). A brief summary will be given of the general interpretations that have been made.

2.2.1 Geochemistry of surface thermal manifestations.

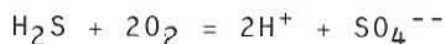
Figures 3 and 4 show the distribution of hot and warm springs of the SNGF. On the basis of their chemical composition, three distinct types have been identified:

- 1) Low lying, near neutral to slightly alkaline, high sodium chloride springs of Palinpinon and Cambucal.
- 2) Near neutral to acidic, high sulfate (up to 3,000 ppm) waters with significant chloride (780 ppm) of Tagbac-Magaso, Baslay and Nagpantaw at moderate elevations.
- 3) Acid, low chloride-high sulfate steam heated springs of Lagunao, Sogongon, and Kaipohan, at high elevations.

The first type is attributed to a direct leakage from the aquifer fluid, and hence, believed to reflect underground hydrothermal conditions. Chemical geothermometry and mixing models yield maximum and minimum subsurface temperatures of 279-173°C (KRTA, 1979).

The second type is believed to be associated with the hot chloride water but could have been diluted with near surface ground water. The neutral springs of this type could be formed when the hydrothermal fluids migrate upward via a formation of restricted permeability, where neutralization reactions between rock and fluid are enhanced by the slow flow rates, resulting in neutral emissions at the surface. The acid springs could be formed when steam separates from the water table of the hot chloride body and dilutes near surface ground water during upward flow.

The third type is evidently steam heated surface ground water and does not reflect deep conditions. The high sulfate and low pH could be quantitatively explained through the oxidation of hydrogen sulfide gas in the steam to sulfuric acid upon contact with oxygen containing meteoric water:



Glover (1975) concluded that the type of discharge could be controlled by elevation and permeability, and Barnett(1980) suggests a steam cap to be responsible for the steam heated springs.

2.2.2 Geochemistry of exploration and step-out wells.

Early exploration and step-out wells (N1 to N3, OK-2 to OK-5) intercepted higher saline fluids than the least diluted springs of Palinpinon and Cambucal, a consistent sodium chloride water containing variable concentrations of sulfate with a mineralization of about seven percent of total dissolved solids. The absolute chemical components vary slightly for each well and yet show relatively uniform ratios of non-reactive elements e.g. Cl/B, suggesting a single source of parent water undergoing varying degrees of mixing and dilution with cold ground water during ascent. An evidence for mixing is the linear relation between the absolute chloride content and quartz calculated temperatures (KRTA, 1977). Temperature dependent and reactive constituents e.g. Na, K, SiO₂ vary in direct proportion to the temperature of the deep water at any point, which is believed to range from 170°C to possibly 300°C. Wells N1 and N2 however, encountered problems with calcite deposition and output testing did not last long enough to allow any geochemical monitoring with time. Such deposition problems were correlated to the presence of abundant calcite that has crystallized in open fractures as reported from the petrological logs of cores that were cut between the major loss zones of N1. A similar problem was predicted for N3, (Glover, 1975, Barnett, 1980), but no indications of drastic decline in total mass flow rates have been observed. Consequently, no problems of calcite scaling have been encountered in this well. The maximum measured temperatures for N1, N2, N3, OK-2 and OK-5 were 195°C, 163°C, 238°C, 247°C and 310°C, respectively. All

geochemical data point to an easterly outflow of hot chloride water from west to east. Each well intercepted this flow regime and chemically calculated temperatures which were generally higher than measured values at conditions of stable flow, indicated that higher temperatures could exist towards the southwest. Preliminary isochemical plots of Obusan (1980) suggests a major upwelling zone within the vicinity of OK-5.

2.2.3 Geochemistry of production wells

Information on the subsurface chemistry of the area has been updated with the successful discharge of deep production wells. From 1980 to the first half of 1982, a total of 14 wells had undergone medium and long term discharge tests. These wells initially undergo a 1-2 month clearing discharge period, followed by a successive throttle series with the primary objective of examining variations in bore output characteristics in response to changes in wellhead pressures and end pipe diameters. Wells OK-5, PN13D, and OK-7 are currently utilized to drive the 3x1.5MW non-condensing turbines, which started operating in 1981 and 1982, while PN17D, PN15D, PN19D, NJ1D, and OK-8 are undergoing or have just completed their respective medium term discharge tests. OK-12D and OK-2 are used as reinjection wells. The boreholes in Southern Negros show slight two-phase conditions, except for SG1, OK-12D, PN13D, which appear to draw production from a single phase fluid. These wells have a discharge enthalpy ranging from 1050 to 1350 kJ/kg at wellhead pressures of 4-12 bars absolute; only OK-5 shows extreme two-phase conditions with an average discharge enthalpy of 1900 kJ/Kg under fully opened conditions. Most wells encounter two major permeable zones at depth which is referred to as the main and secondary production zones (EDC-KRTA, 1982). Table 2 shows the physical measurements data of all production wells in S-Negros and Fig. 5 shows a cross section map of the SNGF showing these production zones.

Table 2.

Physical measurements, data and interpretations of production wells in S- Negros, showing production zones and measured temperatures

| Well No. | Depth (m) | Loss zones (m)** main /secondary | Main Production *** zones(m) | Measured temperatures(°C) | | | Power rating MWe |
|----------|-----------|-------------------------------------|---------------------------------|---------------------------|-----|-----|---------------------|
| | | | | a | b | c | |
| OK-7 | 2883 | 2600-2882/1500-1700 | 2600 | 268 | 318 | 318 | 12.5 |
| OK-9D | 2889 | 2400-2600/1600-1800 | 2400 | 306 | 332 | 291 | 6.4 |
| OK-10D | 3019 | 1500-1750/2900-3019 | 1600 | 226 | 270 | 259 | 6.8 |
| OK-12D | 2798 | 1400-1700/2000 | 2013 | 243 | 282 | 250 | 4.4 |
| PN13D | 2854 | 1600-1900/2000-2854 | 2000 | 236 | 317 | 268 | 4.6 |
| PN15D | 2807 | 1500-1600/2200-2806 | 1600 | 265 | 276 | 274 | 7.5 |
| PN17D | 3043 | 2000-2150/1600-1780 | 2100 | 281 | 290 | 245 | 4.5 |
| OK-5 | 1975 | 1450-1550/1100-1200 | 1550 | 310 | 290 | 264 | 8.2 |
| OK-6 | 2771 | 2300-2770/1340-1550 | 2550 | 254 | 285 | 285 | 10.1 |
| OK-8 | 2982 | 2300-2400/1700-1900 | 2450 | 282 | 282 | 244 | 3.0 |
| N31D | 2848 | *2050-2400/1325-1610 | ---- | 243 | --- | --- | --- |
| SG-1 | 2763 | 2550-2650/1550-1650 | 2650 | 278 | 277 | 277 | 5.5 |
| SG-2 | 2945 | 2050-2150/1400-1600 | 1400 | 275 | 276 | 275 | 10.1 |

* Data available only from drilling logs

** Loss zones indicate zones of water losses deduced from pressure and temperature logs during completion testing of the well.

***Refers to the zone which dominates production at maximum discharge rate of the well; deduced from temperature and pressure logs at flowing conditions.

a Maximum measured well temp. before discharge

b Maximum measured well temp. after discharge

c Maximum measured temp. of main production zone in flowing conditions

The typical fluid chemistry is presented in Table 3. The wells are grouped on the basis of their respective locations, i.e. the Puhagan, Nasuji, and Sogongon areas.

Obtaining representative chemical data from a well is hindered by several factors. Fig. 6 shows the typical variations of chemically calculated subsurface temperatures, deep water enthalpy, chloride in total discharge, and the gas components (CO_2 and H_2S) in total flow with the normal duration of a discharge test of a well. Because of the large amount of fluids injected into the geothermal reservoir during drilling and completion tests, a well fluid is not representative until after approximately two months of full discharge. Another important factor to be taken into consideration is the mixing between aquifers. The existence of two production zones at a temperature difference of about $40\text{-}50^\circ\text{C}$ both contributing to the total flow results in a mixed fluid at the surface sampling point.

The criteria for selecting representative data from individual wells are thus a) a fully stabilized chemistry and b) a sample point where only the major production zone dominates bulk of the discharge, which, according to experience from temperature pressure logs at flowing conditions, occurs at throttled conditions. These criteria will also minimize the effects of excess enthalpies. All chemical data in Table 3 is taken under these conditions. The fluids are relatively neutral (pH 6-7) but contain high amounts of sodium, potassium, silica, and chloride. Significant amounts of the alkali metals are also present (Li, Rb, Cs), but magnesium, bicarbonate, and sulfate, though variable, are low. In the steam phase, carbon dioxide comprises more than 90% of the total non condensible gases, whereas hydrogen sulfide and ammonia comprises 2-4 %. Following is a summary of the interpretations about the reservoir fluid that have been made:

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Table 3

Chemical composition of production well fluids at S-Negros field (all concentrations in parts per million)

| Well No. | H ₂ S kg/kg | SP bars abs | Steam CO ₂ | phase at SP H ₂ S | SP bars abs. | pH/25 C | Li | Na | K | Ca | Mg | Water phase at SP Rb | Cs | Cl | SO ₄ | B | SiO ₂ | CO ₂ | H ₂ S |
|----------|---------------------------|----------------|--------------------------|---------------------------------|-----------------|---------|------|------|-----|------|------|-------------------------|------|------|-----------------|------|------------------|-----------------|------------------|
| 08-7 | 1326 | 9.30 | 34,296 | 695.7 | 9.30 | 6.97 | 14.8 | 3161 | 774 | 118 | 0.13 | 4.27 | 3.00 | 5788 | 26.0 | 69.3 | 867 | 32.7 | 0.10 |
| 08-9D | 1786 | 4.52 | 21,599 | 493.5 | 4.52 | 7.43 | 17.4 | 2580 | 555 | 75.5 | 0.19 | 3.56 | 2.71 | 4308 | 34.7 | 52.0 | 734 | 43.6 | - |
| 08-10D | 1413 | 2.59 | 7072 | 165.8 | 2.59 | 6.69 | 14.0 | 3167 | 634 | 122 | 0.85 | 3.79 | 3.27 | 5474 | 51.5 | 76.1 | 717 | 2.6 | - |
| 08-12D | 1089 | 11.00 | 31,729 | 435.3 | 11.00 | 6.70 | 10.7 | 2352 | 457 | 76.0 | 0.36 | 2.53 | 2.10 | 4361 | 32.0 | 42.2 | 532 | 43.0 | 82.6 |
| PH14D | 1703 | 11.50 | 23,082 | 69.0 | 11.50 | 5.95 | 10.1 | 2472 | 551 | 58.3 | 0.38 | 3.17 | 2.33 | 4773 | 24.9 | 61.3 | 675 | 18.3 | 10.2 |
| PH15D | 1152 | 4.80 | 23,293 | 317.9 | 4.80 | 6.01 | 10.7 | 2800 | 487 | 63.1 | 0.26 | 2.70 | 2.31 | 4965 | 25.6 | 57.0 | 757 | 154 | 5.96 |
| PH17D | 1094 | 6.60 | 19,154 | 327.5 | 6.60 | 6.21 | 10.0 | 2965 | 421 | 68.6 | 0.42 | 2.41 | 2.16 | 4617 | 30.6 | 53.3 | 559 | 40.2 | 7.16 |
| 08-5 | 1818 | 9.30 | 10,805 | 356.3 | 1.00 | 7.27 | 17.5 | 3347 | 820 | 35.1 | 0.04 | 5.31 | 3.58 | 5993 | 61.5 | 73.0 | 1073 | 68.5 | 9.88 |
| 08-6 | 1263 | 11.10 | 17,738 | 204.4 | 11.10 | 6.60 | 11.6 | 2911 | 480 | 31.2 | 0.03 | 3.87 | 2.81 | 4436 | 32.5 | 52.0 | 726 | 6.60 | - |
| 08-8 | 1262 | 7.62 | 11,480 | 192.5 | 1.00 | 7.55 | 14.3 | 2771 | 636 | 33.6 | 0.05 | 4.27 | 2.87 | 5262 | 36.1 | 64.4 | 864 | 6.98 | 6.48 |
| 0810 | 1046 | 2.00 | 10,218 | 1102.0 | 1.00 | 5.20 | 6.88 | 1230 | 323 | 16.8 | 5.25 | 1.77 | 1.80 | 3489 | 155 | 46.9 | 600 | 2.54 | 2.39 |
| SG-1 | 1147 | 6.20 | 9842 | 154.4 | 6.20 | 7.30 | 10.7 | 2250 | 487 | 34.4 | 0.06 | 2.97 | 2.15 | 4148 | 39.8 | 51.1 | 626 | 32.1 | 3.07 |
| SG-2 | 1280 | 9.20 | 14,691 | 182.6 | 9.20 | 5.91 | 10.4 | 2139 | 452 | 21.4 | 0.06 | 2.99 | 2.07 | 3957 | 23.2 | 48.6 | 660 | 25.8 | 3.55 |

H measured discharge enthalpy (in KJ/kg)

US total sulfides (H₂S, HS⁻, S²⁻)

SP sampling pressure (in bars absolute)

--- analysis not available

CO₂ total carbonates (HCO₃⁻, CO₃²⁻, H₂CO₃) reported as CO₂

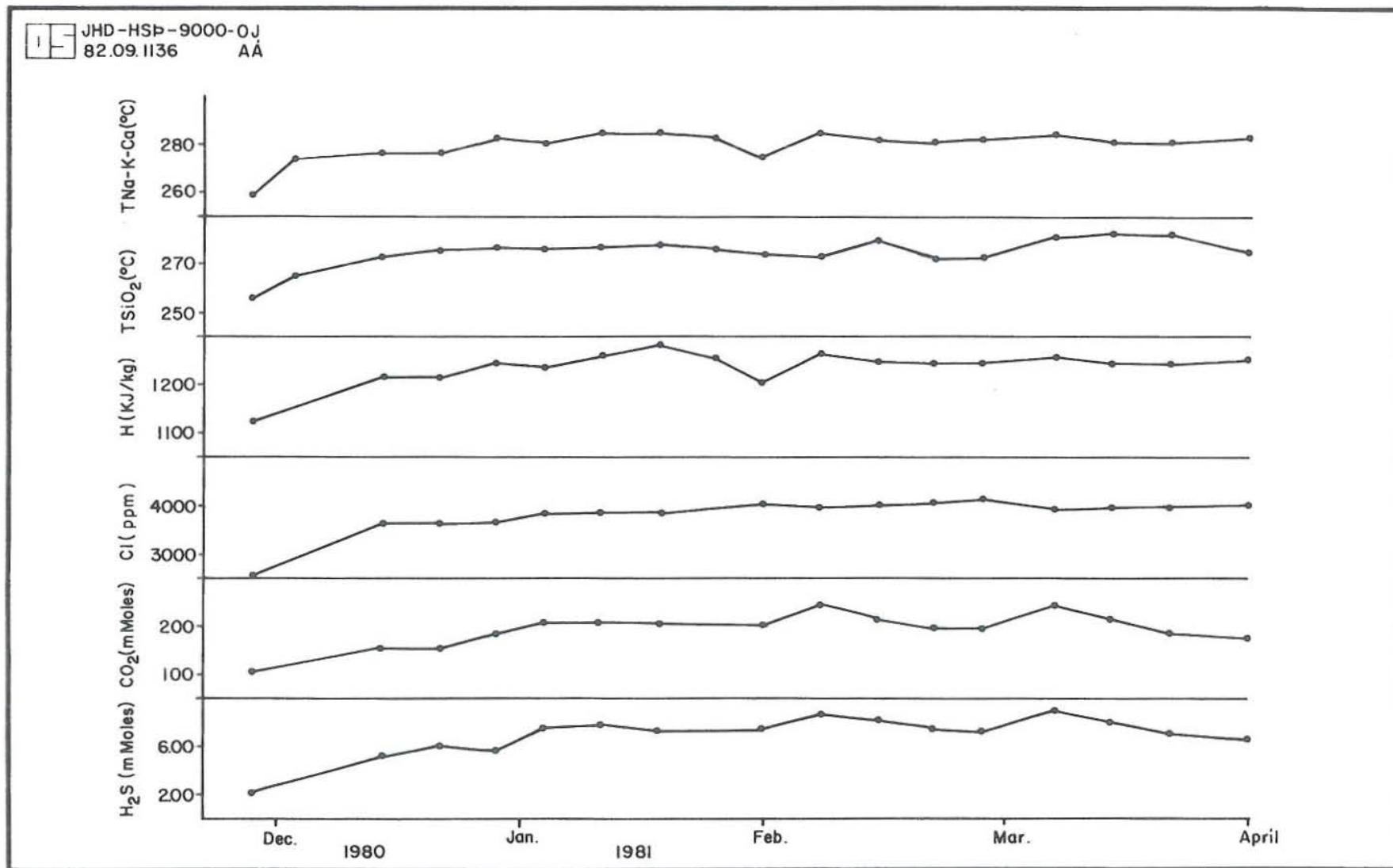


Figure 6 Typical variation with time of chemically calculated temperatures, deep water enthalpy, deep water chloride, and CO₂ and H₂S in total discharge of a production well (OK-9D).

1) Two production zones are deduced from the variation of total discharge chemistry in response to differing wellhead pressures, namely, a deep, highly saline fluid containing typically 3,900 ppm of chloride at a mean temperature of 280°C (T Na/K), and a less mineralized but gas rich fluid at shallower levels. Variations with time and changing wellhead pressures are interpreted as an intermittent contribution from the upper feed zone (Jordan, 1981b,c, Clemente, 1982a,b). This also attributes to the wellhead pressure cycling that was noted in wells OK-5 and N3 (Camales 1979). This deduction is supported by fluid losses during completion testing and the temperature and pressure logs at flowing conditions, where at least two producing horizons are recognized. The top and bottom of these exploitable zones are believed to be at 300m bsl and 2600m bsl with mean temperatures of 230°C and 320°C respectively (KRTA 1981).

2) The Puhagan wells intercept the least diluted fluid, whereas the Nasuji and Sogongon fluid chemistry has been shown to be modified by dilution with cooler fluid, and oxidation with oxygen containing meteoric water. A common reservoir, however has been consistently indicated by the uniformity of alkali metal ratios, and the chloride to boron equivalence (Jordan, 1981,b,c).

3) The iso-chemical and isothermal contours (Harper and Clemente, 1982) suggest a narrow flow of thermal fluids to the northeast of Puhagan, where the contours close into the Okoy valley. This could be directly associated to a structural control of the Ticala fault which may confine the outflow in a narrow, northeasterly direction. A broad flow to the west through Nasuji and further into Sogongon is also evident, since the contours show a broader extension. This could be due to secondary permeability by thermal cracking, and not directly associated with faulting.

4) The main upflow zone which is postulated to lie beneath the Cuernos volcano, could probably be in temperatures of greater than 300°C. This was estimated by extrapolation of plots of deep chloride versus silica calculated temperatures (Harper and Clemente, 1982, Jordan, 1981 b,c).

3 SOLUTION-MINERAL EQUILIBRIA

3.1 Introduction

The study of solute/mineral equilibria has been shown to be one of the valuable aides in the overall geochemical evaluation of a geothermal system. This is based on the fact that fluid chemistry is primarily governed by the solubility of minerals, and thus an understanding of the thermodynamic factors which affect such equilibria is necessary for the proper interpretation of chemical data. The use of chemical geothermometers, and prediction of mineral scaling and deposition tendencies, for example, are shown to be very dependent on the saturation state that has been attained between rock and fluid.

Previous geochemical studies in the S-Negros Field, have mainly been focused on describing individual wells and field models. Only a limited attention has been given to a quantitative evaluation of the solute-mineral equilibria. It is the objective of this report to attempt such an evaluation investigation, and furthermore, to:

- (1) evaluate the implications of solution mineral equilibria on the applicability and limitations of previous and current chemical geothermometers,
- (2) anticipate mineral deposition tendencies, and,
- (3) predict physical processes (dilution, conductive cooling, precipitation) across the geothermal field.

3.1.1 Applications of solute-mineral equilibria studies

Previous workers, (Fournier 1977; Ellis and Mahon, 1977; Arnorsson, et al. 1982b) have demonstrated that most high temperature geothermal waters attain an overall equilibrium between solutes and alteration minerals. It is known that

chemical species is controlled by mineral solubility i.e. SiO_2 is controlled by the solubility of quartz and chalcedony, Na and K by sodium and potassium feldspars, calcium by calcite and calcium aluminum silicates, and sulfate by anhydrite and iron minerals. The state of saturation, however, is highly dependent on the physical processes which occur during the transport of the fluid from the reservoir to the surface. Intensive studies have shown (Arnorsson et al., 1982 b), that boiling, cooling, degassing, dilution with cooler fluid and the kinetics of water-rock interactions, are the major processes which cause some minerals to deviate from the original saturation state. First level boiling with maximum degassing, for example, has been shown to produce a supersaturated fluid with respect to calcite, and, progressive boiling and cooling to significantly lower temperatures yields amorphous silica saturated fluids, (Ellis and Mahon, 1977; Arnorsson, 1981). It is thus, highly desirable for the proper interpretation of chemical data, that such factors be taken into consideration.

The use of chemical geothermometers in predicting underground temperatures is based on the assumption that equilibrium between fluid and a particular mineral has been attained, and that re-equilibration during conductive cooling is insignificant. Many different chemical geothermometers have been used to estimate reservoir temperatures, among these are silica (Mahon, 1966), Na/K (White 1965; Ellis and Mahon, 1977; White, 1970; Truesdell, 1976a; Fournier, 1979a; Arnorsson et al. 1982 c) and Na-K-Ca (Fournier and Truesdell, 1973). They are based on the experimentally determined solubility of quartz and the variation of Na^+ and K^+ in natural geothermal waters as a function of temperature. The Na-K-Ca geothermometer was developed to deal with calcium rich waters that give anomalously high results when using the Na/K geothermometers, and is empirically calibrated.

It is important to consider the sensitivity of each chemical species (that is used as a temperature function), to physical and chemical changes accompanying the passage of fluid from the reservoir to the surface sampling point. The equilibria governing sodium and potassium feldspars are least affected by cooling, and the use of the ratios of sodium to potassium eliminates boiling and dilution effects, and hence, temperatures estimated from these geothermometers would tend to be higher than those of quartz. The quartz geothermometer, on the other hand, uses directly absolute concentrations of total silica, and would thus readily be affected by dilution. Furthermore, the equilibrium governing quartz and silica is more sensitive to conductive cooling and boiling and generally yields conservative values. Arnorsson et al. (1982 c) showed that the rate of quartz precipitation is positively related to water salinity. The rate of re-equilibration upon cooling between quartz, feldspar and solution is about the same for dilute solutions (1,000 ppm) for water temperatures of about 250°C (Namafjall, NE- Iceland Arnorsson, 1977). For waters of high salinity (13,000 ppm, Svartsengi SW-Iceland), conductive cooling by 10-20°C is fully detected by the quartz geothermometer, but not in the case of Na/K geothermometer.

The Na-K-Ca geothermometer is empirically calibrated involving the addition of two temperature dependent ratios (Na^+/K^+ and $\text{Ca}^{++}/\text{Na}^+$). This geothermometer however, (unlike quartz and Na-K) is not in any way related to a specific mineral assemblage and the major reason why it works in some cases is that the free energy that is involved in the dissolution of a particular mineral and the formation of new ones are negligible so that no significant change in water composition is detectable. Furthermore, the work of Arnorsson et al. (1982 c) shows a larger scatter of data points for the ratio $\text{Ca}^{++}/\text{Na}^+$ than for Na^+/K^+ indicating that the scatter could be due to equilibrium with more than one mineral assemblage, whereas Na^+ and K^+ activities seem to be governed by equilibrium between low-albite and microcline

over the whole temperature range. This makes the Na-K-Ca geothermometer inferior to the Na/K geothermometer whatever calibration curve or beta value is used.

Temperature estimates of the Na-K-Ca geothermometer are more conservative than those of the Na-K geothermometer since large monovalent cations take a long time to equilibrate, whereas divalent, trivalent, and smaller ions re-equilibrate more rapidly. Hence, the $\text{Ca}^{++}/\text{Na}^{+}$ would respond faster to cooling than the Na/K ratio. Another limitation to the application of the Na-K-Ca geothermometer is the precipitation of the calcium ion in significant amounts, especially in zones of intense boiling. This would then result in anomalously high calculated temperatures (Arnorsson et al. 1982 c).

Gas geothermometers have also been used, (Ellis, 1979; D'Amore and Panichi, 1980) which involves the temperature dependence of gas-gas equilibria and the relative concentrations of CO_2 , H_2S , H_2 , and CH_4 in geothermal steam. The calibration is empirical and shown to be in a reasonable agreement with measured temperatures from nearby drillholes. However, re-equilibration between gases during upflow to the surface, has not been duly considered. The apparent discrepancy in temperature values obtained by various geothermometers thus often aid in identifying processes, such as:

- a) dilution with cooler fluid,
- b) conductive cooling,
- c) precipitation and/or deposition along upflow zones accompanying boiling.

It is obvious that although several calibration curves are used in deriving temperature functions for chemical geothermometers, the most common practice in its use is three-fold, namely:

- (1) Complexing between solutes is ignored and analytical concentrations are assumed to represent activities.
- (2) Solubility measurements are mainly done at 25°C and just extrapolated to higher temperatures; effects of fluid pressure on mineral solubility is regarded insignificant, and,
- (3) For gas geothermometers, effects of re-equilibration between various gases has been neglected.

The first case may be true for dilute solutions (<3,000 ppm chloride), but in most geothermal fields, where salinity is a significant factor, e.g. Reykjanes, Iceland (32,000 ppm Cl), Tongonan, Philippines (13,000- 18,000 ppm Cl), Salton Sea, California (33,000 ppm Cl), this assumption often gives misleading results. Arnorsson, et al. (1982 c) demonstrate that temperatures calculated for waters of high salinity and high sulfate from analytical sodium and potassium values, do indeed deviate more significantly from measured temperatures, than those from computed activities. In the second case, Ragnarsdottir and Walther (1982) demonstrated in their quartz solubility measurements at 250°C and 250, 500, and 1000 bars that pressure poses a significant effect on the silica concentrations especially at pressures above 500 bars. These experiments, together with previously reported quartz solubility measurements, yielded a "pressure sensitive silica geothermometer" that takes into account the effect of fluid pressure on the solubility of quartz. The temperature obtained from this geothermometer shows best agreement with measured temperatures in many different geothermal areas (e.g. in Iceland and Kenya) and as will be discussed in the preceding chapters, in S-Negros, Philippines.

Experience, however, shows that fluid pressure in geothermal wells does not reach as high as 350 bars and the main reason why this geothermometer seems to show best agreement with measured temperatures is that the solubility

measurements are done at 250°C, which represents close to actual reservoir conditions, whereas the other quartz solubility measurements are done at 25°C and just extrapolated to higher temperatures.

For the gas geothermometers, Arnorsson et al. (1982 c) showed that gas-steam geothermometers are better than gas-reaction geothermometers since boiling and reaction with rock have less effect on the gas-steam ratios than on the gas-gas ratios. Furthermore, the previous gas geothermometers, which involve H₂, H₂S, and CH₄, are not known to be controlled by any mineral assemblages. The CO₂ gas geothermometer, proposed by Arnorsson et al. (1982 c), is believed to be controlled by the zoisite-epidote-phrenite-quartz-calcite assemblage and this assemblage has been found in many geothermal systems (Browne, 1978).

Arnorsson et al. (1982 b) demonstrated that the ratios of $a_{\text{Na}^+}/a_{\text{H}^+}$, $a_{\text{K}^+}/a_{\text{H}^+}$, $\sqrt{a_{\text{Ca}^{++}}}/a_{\text{H}^+}$, $\sqrt{a_{\text{Mg}^{++}}}/a_{\text{H}^+}$, and acid molecules e.g. H₂CO₃, H₂SO₄, H₂S, HF, H₄SiO₄ are always fixed at a particular temperature, provided equilibrium has been obtained between rock and fluid. This known relation enabled the derivation of empirical equations which relate the ratios and concentrations of acid molecules to fluid temperatures. These temperature functions combined with other solubility curves of certain minerals (e.g. anhydrite, quartz, calcite, for this study) would aid in precise determinations of whether or not equilibrium has been attained between rock and fluid. This would furthermore check the reliability of the use of chemical geothermometers, and would help in evaluating what processes could be responsible for non-attainment of equilibrium, e.g. mixing, precipitation, etc.

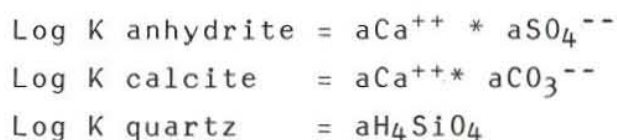
An attempt to apply such an empirically derived relation will be made on the S-Negros field in this report. It is believed to be applicable to define water composition in other areas for two reasons:

- (1) Free energies accompanying hydrothermal reactions are too small to be detected in water chemistry,
- (2) Variations in alteration mineralogy at elevated temperatures are not significant in volcanic rocks ranging from basaltic to rhyolitic compositions, Browne (1978).

Mineral deposition, corrosion, and scaling can cause a major constraint in the exploitation of geothermal fields. The most common scales are calcite, aragonite, amorphous silica, and iron oxides. Deposition is directly associated with temperature dependent mineral/solute equilibria controlling water chemistry. Mineral deposition, as a rule, occurs when the fluid attains a supersaturated state with a particular mineral, which usually follows intensive boiling and cooling. Thus, an understanding of how particularly troublesome minerals deviate from equilibrium in response to physical changes is highly desirable in order to anticipate and control scaling problems. As will be discussed in later sections, a computer simulation of cooling, boiling and degassing was used to follow mineral saturation states in response to the above mentioned processes.

3.1.2 Mineral saturation state of the deep fluid

Figures 7, 8 and 12a show the state of anhydrite, calcite, and quartz saturation of the deep water, with reference to the measured temperatures at the main production zone. The solubility values for each was calculated from the thermodynamic data of Helgeson (1969), and Helgeson et al. (1978), and the activities of the species from analytical values, using a computer programme described by Arnorsson et al. (1982 a). The saturation state of each mineral is expressed as:



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Figure 7 State of anhydrite saturation of the deep water at measured temperatures.

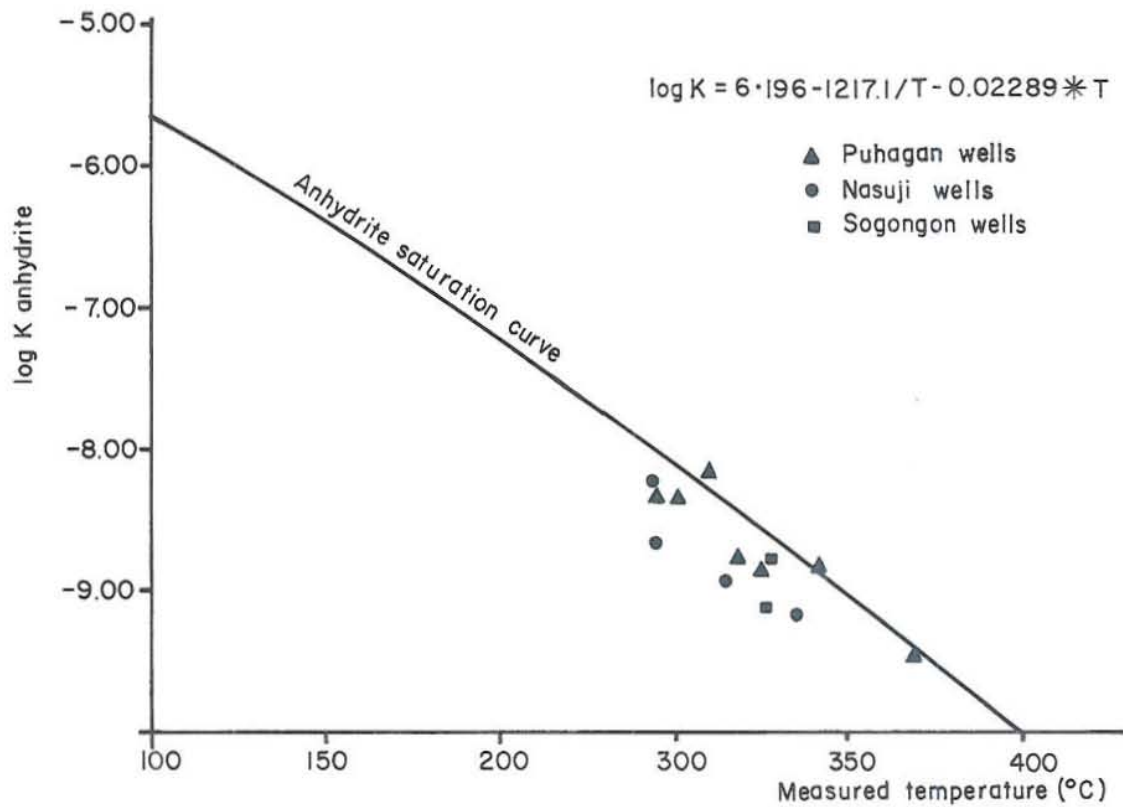
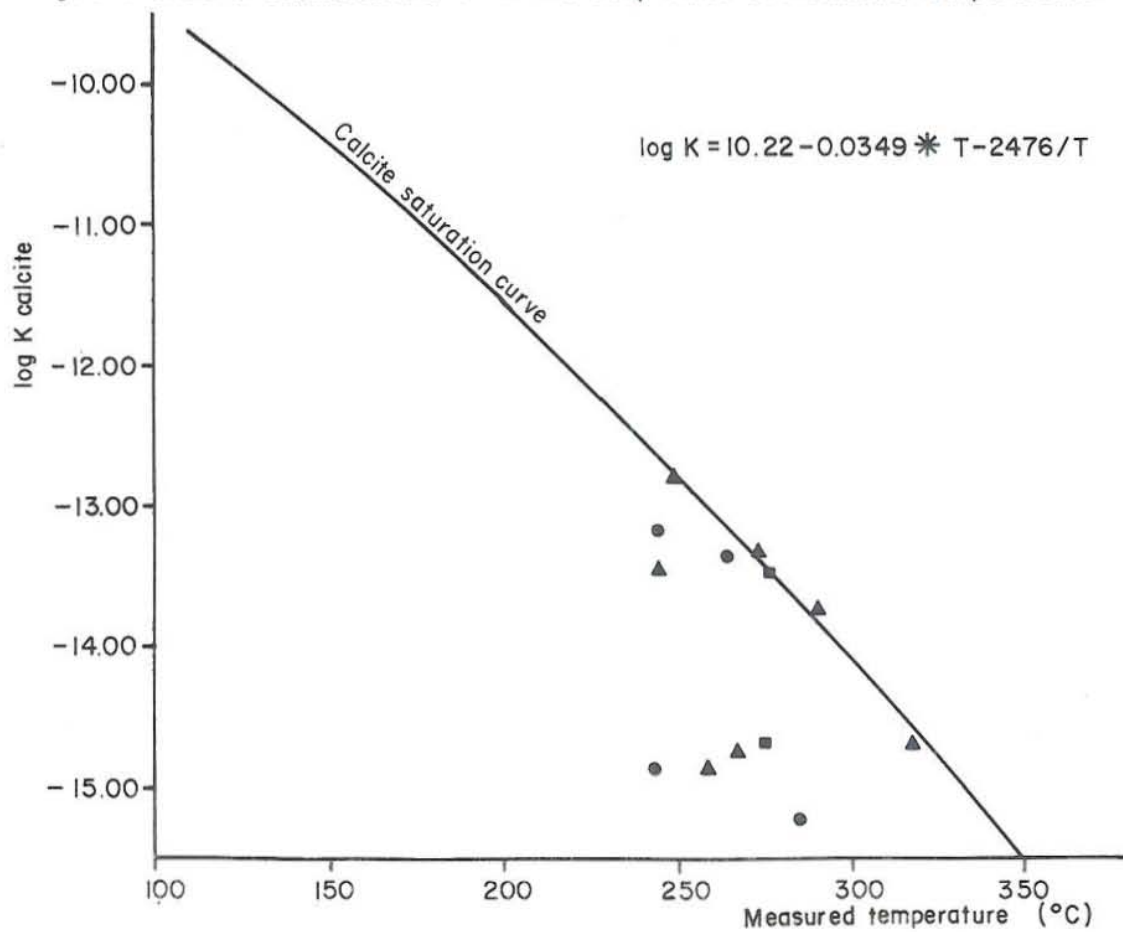
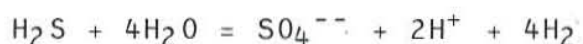


Figure 8 State of calcite saturation of the deep water at measured temperatures.



It is obvious from the figures that anhydrite and quartz saturation has been attained for most wells, only OK-6, OK-8, NJ1D and SG-2 shows slight undersaturation with respect to anhydrite. It is believed that the lower fluid salinity of these wells, caused probably by mixing and dilution, influences the control of sulfate activity by the reaction, (Arnorsson et al. 1982 a).



The equilibrium condition of anhydrite and quartz is not so sensitive to the reference temperature selected or at simulated runs for conditions of no excess enthalpies, since the calculated activities of aqueous H_4SiO_4 , Ca^{++} , SO_4^{--} do not deviate significantly from the solubility curves.

The saturation state with calcite varies from well to well, but it generally seems to be disturbed by mixing with cooler fluid, which is more pronounced within the Nasuji-Sogongon areas. There most of the wells draw production from Calcite undersaturated fluids as shown in Fig. 8. This is supported by previous plots of aquifer chloride versus calculated silica temperatures where wells SG2, SG1 and OK-6 fall on the dilution line (Clemente 1981 a; Jordan 1981 b,c).

In order to determine the deviation from the original equilibrium state of the fluid to cooling, boiling, and degassing, and to anticipate mineral deposition tendencies, a computer simulated approach was used. Figures 9 to 11 show the deviation curves for Log K values of quartz and amorphous silica, calcite, and anhydrite in response to a computer simulation of cooling, boiling, and degassing processes. An arbitrary choice of 9 boiling steps at conditions of maximum degassing and 7 cooling steps were taken to see their respective deviation from the equilibrium curve. The first four boiling steps are selected at close temperature intervals (2°C), since the

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Figure 9 State of quartz and amorphous silica saturation at boiling and cooling

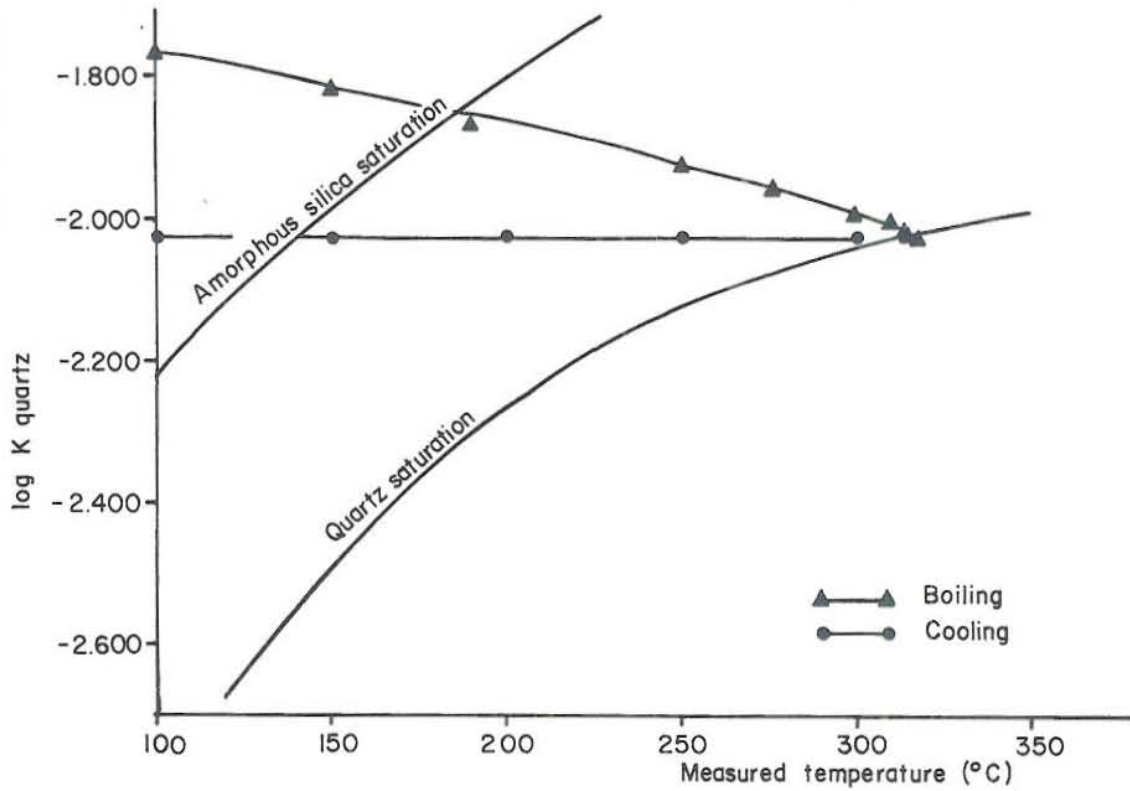
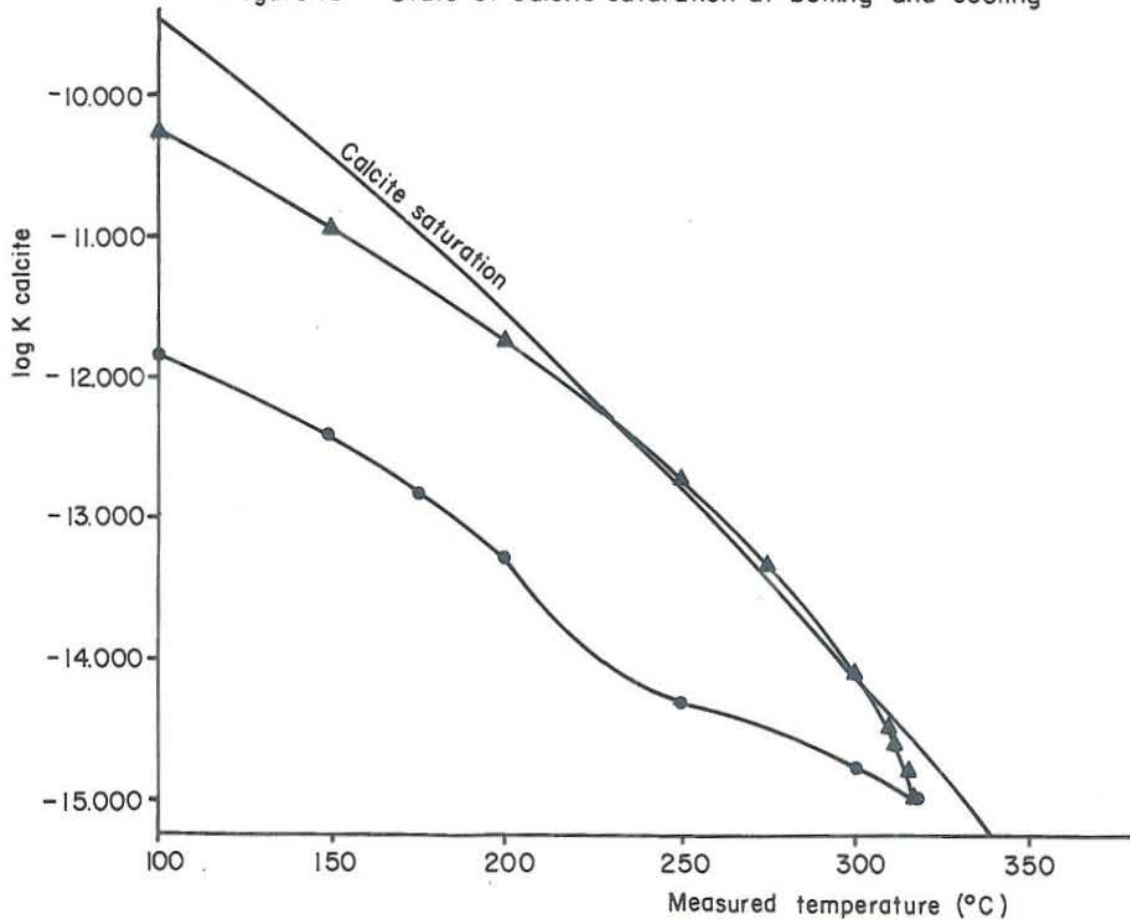
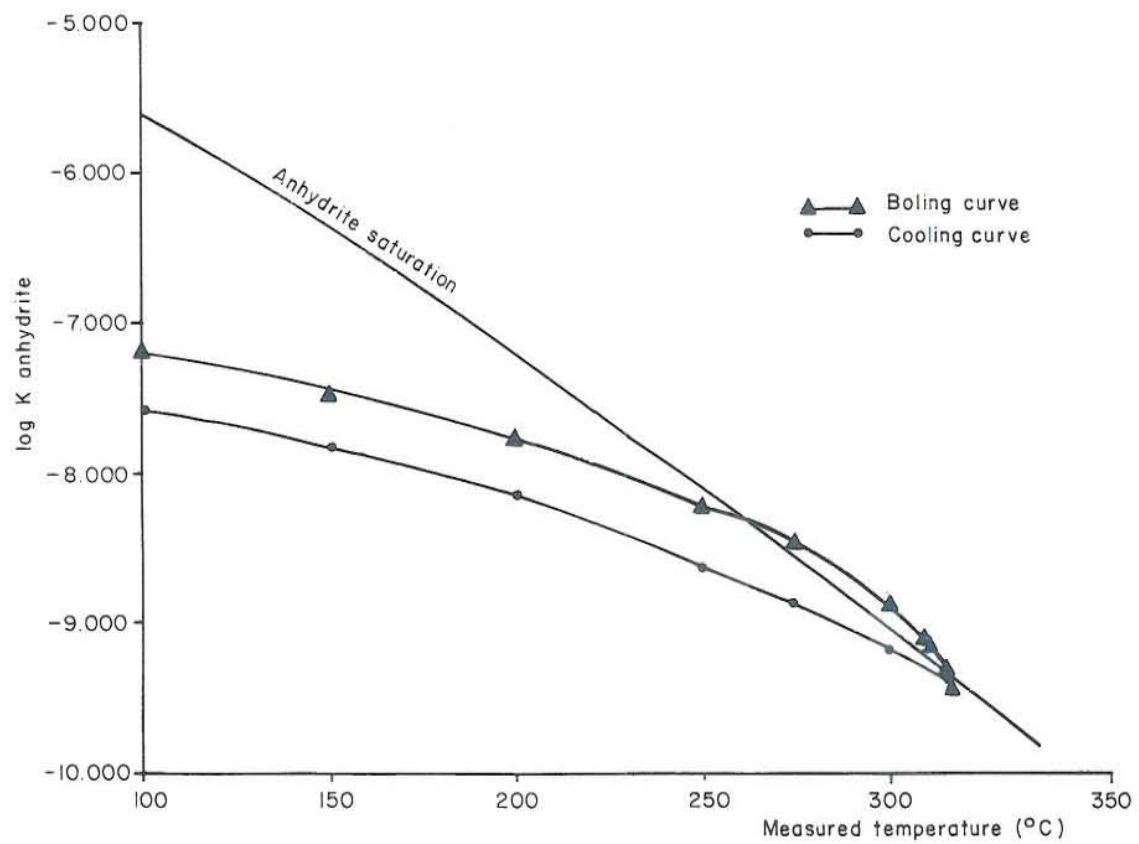


Figure 10 State of calcite saturation at boiling and cooling



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Figure 11 State of anhydrite saturation at boiling and cooling



most intense deviation from original saturation state usually occurs during the first few levels of boiling. The most representative well (OK-7) of the deep fluid was taken for this study since simulation runs from individual wells showed similar characteristic curves.

It is obvious from the figures for calcite and anhydrite that conductive cooling always yields undersaturated fluids in all the temperature range covered. The calcite and anhydrite curves indicate an approach to slight supersaturation during the first four steps of adiabatic boiling, attaining a maximum deviation from equilibrium after 43°C and 18°C adiabatic boiling for calcite and anhydrite, respectively. The rate of change decreases with progressive steam separation until undersaturation is attained. The slight approach to supersaturation with respect to calcite is not, however, regarded as an indication of potential scaling since log K values fall within the anticipated limit of experimental error.

In the case of silica, conductive cooling yields supersaturated fluids with respect to amorphous silica at about 140°C, whereas if waters cool adiabatically, amorphous silica saturation is reached at 186°C. Experience with re-injection of OK-7 waste water at 155°C for several months, however, showed no signs of serious amorphous silica scaling along the re-injection line. This would tend to indicate that other factors still control the extent of deposition apart from attaining supersaturation. Arnorsson (1981) suggests that the degree of degassing, changes in pH from steam loss, the type of flow (turbulent or laminar), and the possible separation of water and steam phases in the aquifer may control the temperature at which amorphous silica saturation is reached. These factors have not been thoroughly taken into consideration in deposition studies at S-Negros. It is anticipated that potential amorphous silica scaling would not occur until below 150°C.

3.1.3 Empirical relation between aqueous species and temperature.

As was discussed in the section 3.1.1, the attainment of an overall equilibrium between solutes and minerals enables an empirical relation to be derived between concentrations of acid molecules, cation to proton ratios, with temperatures in geothermal waters in Iceland. The selection of such components was arbitrary, i.e. some cation to proton ratios or acid molecule concentrations show better correlation than others. A good fit between $a\text{Ca}^{++}/a\text{H}^+$ and $a\text{Mg}^{++}/a\text{H}^+$ and temperature was for example noted, but no correlation found for $\text{Ca}^{++}/\text{Mg}^{++}$ and temperatures.

This is partly attributed to the strong temperature dependence of the calculated hydrogen ion activity, which is caused by the variation with temperature of the dissociation constants of weak acids which control water pH. Table 4 summarizes the temperature functions obtained for the various acid molecules and cation to proton ratios.

An attempt was made to apply such empirically derived relation to the S- Negros field. Figures 12 and 13 show plots of log aqueous acid molecules and log cation to proton ratios. The reference temperatures used are the measured temperatures at the main production zone. The best fit is noted for $\log \text{H}_4\text{SiO}_4$ and $\log \text{H}_2\text{CO}_3$. This indicates that the H_4SiO_4 and H_2CO_3 species govern the total silica and carbonate content of the fluids. The empirical relation for H_2S and H_2SO_4 , though scattered, show a closer fit for the Puhagan than the Nasuji and Sogongon wells, which suggests that the Puhagan area is within a better proximity of the reservoir. $\log \text{H}_2\text{S}$ and $\log \text{H}_2\text{SO}_4$ show a larger deviation from equilibria and lower cation to proton ratios for most of the Nasuji and Sogongon wells. It is believed that the scatter of hydrogen sulfide could be due to the interim loss of this gas from the solution if much conductive cooling and oxidation occurs during the fluid flow from the reservoir to the surface.

Table 4

Equations for temperature dependence of log acid molecules
and log cation to proton ratios(after Arnorsson et al.,1982 b)

| Species/ion ratio | Temperature function | LogK deviation mean /standard | |
|---------------------|--|----------------------------------|------|
| Log H_4SiO_4 | $-0.588-0.00441T-1515.21/T+1.3470\log T$ | 0.06 | 0.05 |
| Log H_2CO_3 | $-1.794-0.00510T-4469.63/T+4.1414\log T$ | 0.30 | 0.26 |
| Log H_2S | $-1.678-0.00355T-5071.05/T+3.8889\log T$ | 0.36 | 0.33 |
| Log H_2SO_4 | $-6.436-0.03906T-13335.68/T+14.7958\log T$ | 0.57 | 0.48 |
| Log aNa^+/aH^+ | $2.694+0.02023T+4243.47/T-6.2069\log T$ | 0.14 | 0.12 |
| Log aK^+/aH^+ | $2.505+0.01971T+3325.71/T-5.7814\log T$ | 0.12 | 0.09 |
| Log aCa^{++}/aH^+ | $1.733+0.01117T+3890.51/T-3.9977\log T$ | 0.17 | 0.12 |
| Log aMg^{++}/aH^+ | $1.816+0.01078T+3727.48/T-4.1640\log T$ | 0.34 | 0.27 |

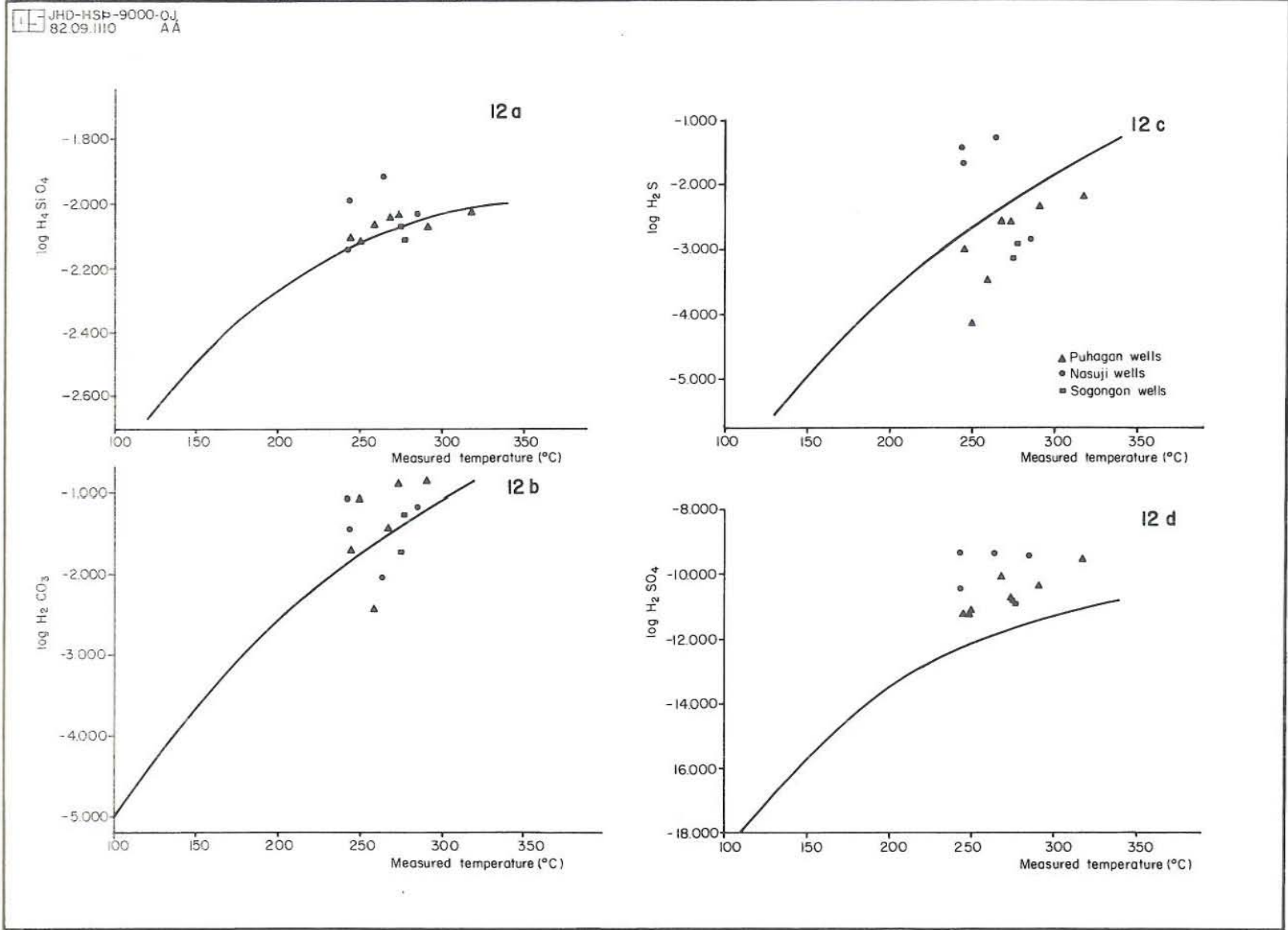


Figure 12 (a) $\log H_4SiO_4$ vs. measured temperatures
(b) $\log H_2CO_3$ vs. measured temperatures

(c) $\log H_2S$ vs. measured temperatures
(d) $\log H_2SO_4$ vs. measured temperatures

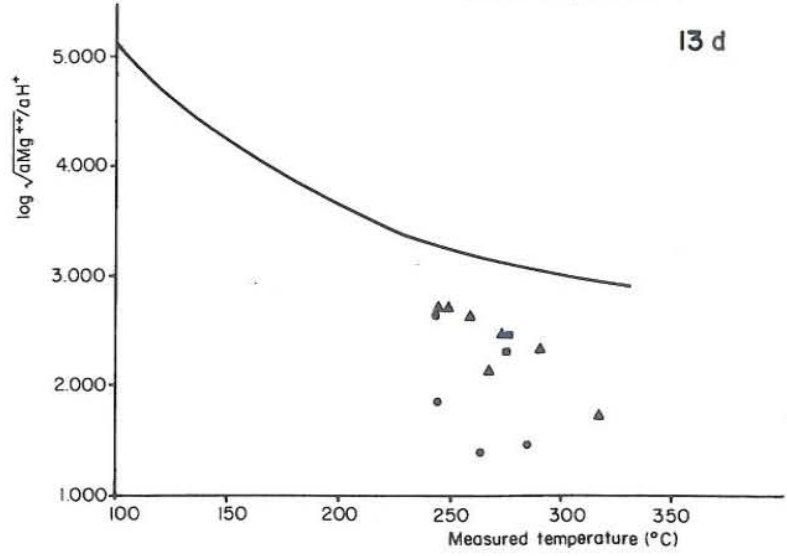
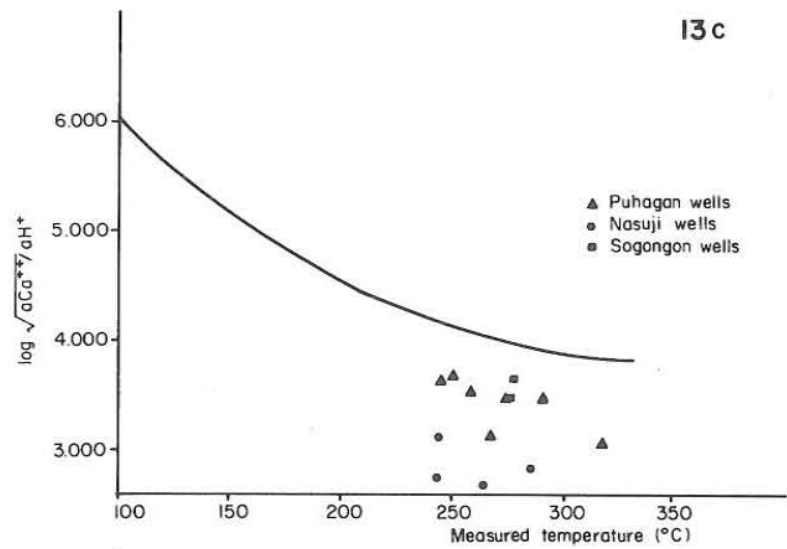
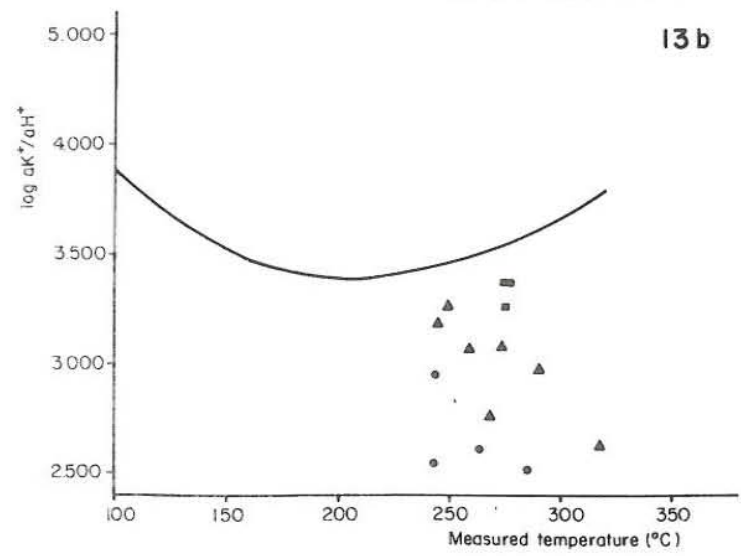
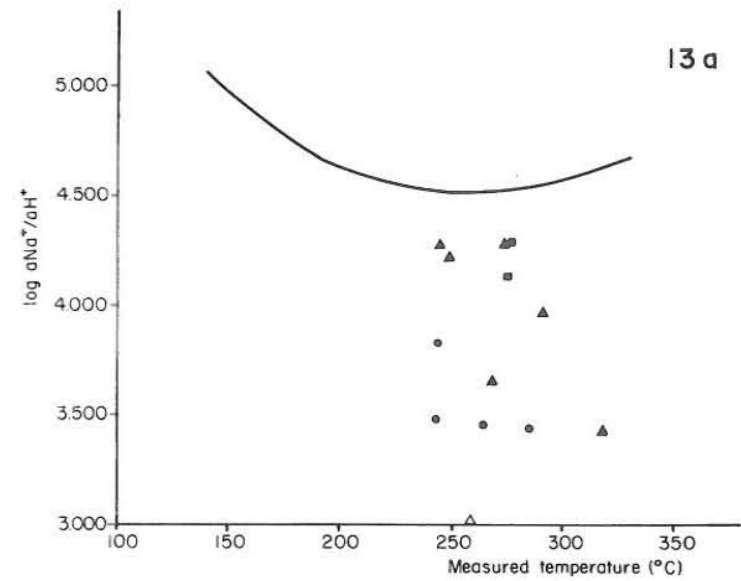


Figure 13 (a) Log aNa^+/aH^+ vs. measured temperatures (c) Log $\sqrt{aCa^{++}/aH^+}$ vs. measured temperatures
 (b) Log aK^+/aH^+ vs. measured temperatures (d) Log $\sqrt{aMg^{++}/aH^+}$ vs. measured temperatures

The rather low cation/proton ratios could readily be attributed to the sensitivity of the hydrogen ion activities to boiling and degassing mechanisms which then have a large effect on the cation to proton ratios. Computer runs at Na-K reference temperatures and at conditions of no excess enthalpy do not in any way give a better fit with the calibration curves. This may indicate that significant mineral re-equilibration due to the presence of excess steam is not responsible for the low cation to proton ratio. It is, however, evident that no equilibrium has been attained between the fluid and the respective species and ratios. Two reasons are suggested:

- (1) Figures 13a-13d indicates that to attain a better fit with the equilibrium curve would require more leaching of cations from the rock. Apparently, leaching rates are not sufficiently high.
- (2) Dilution and mixing with cooler fluid could readily decrease cation concentrations and affect the H^+ concentrations. A mixing model is attempted in section 3.2.4..

3.2 Geothermometry

Table 5 shows the equations, temperature range, and source of the various chemical geothermometers that will be considered in this study. Table 6 shows chemically calculated and measured temperatures of the deep fluids of the production wells in S-Negros, and Table 7 shows the deviations of the calculated temperatures from measured values.

3.2.1 Quartz geothermometry

Since the pH of the separated waters are all below 8.5, the total silica as analyzed can be taken to represent the H_4SiO_4 fraction. This approximation, however, is not valid when the water pH approaches 9, since a significant

Table 5

Table showing equations of various chemical geothermometers used to compare with the measured temp. of main production zone at S-Negros (legend for Table 6)

| Geothermometer | Range (°C) | Source |
|---|------------|--|
| (1.) $T(°C) = \frac{1309}{5.19 - \log SiO_2} - 273.15$ | 0 - 250 | Fournier(1977), SiO_2 in mg/Kg of the deep water; assumes no steam loss |
| (2.) $T(°C) = \frac{1164}{4.90 - \log SiO_2} - 273.15$ | 180 - 300 | Arnorsson et al.(1982 c); SiO_2 in mg/Kg of the deep water; assumes no steam loss |
| (3.) $T(°C) = 771 + (336 + 0.017P) * \log mSi(OH)_4 * 2H_2O + 35.8(\log mSi(OH)_4 * 2H_2O)^2$ | 180 - 340 | Ragnarsdottir and Walther(1982); SiO_2 in moles per liter at 1atm; P - fluid pressure in bars. |
| (4.) $T(°C) = \frac{1647}{\log Na/K + B \log Ca/Na + 2.24} - 273.15$ | 4 - 340 | Fournier-Truesdell(1973); $B=1/3$ for $Ca/Na < 1$ and $T > 100°C$; Na, K, Ca in moles per liter |
| (5.) $T(°C) = \frac{1217}{\log Na/K + 1.453} - 273.15$ | 100 - 300 | Fournier(1979); Na, K in mg/Kg |
| (6.) $T(°C) = \frac{933}{\log Na/K + 0.993} - 273.15$ | 25 - 250 | Arnorsson et al.(1982 c); Na, K in mg/Kg |
| (7.) $T(°C) = \frac{1319}{1.699 - \log Na/K} - 273.15$ | 250 - 350 | Arnorsson et al.(1982 c); Na, K in mg/Kg |
| (8.) $\log Na/K = -1.752 - 2775.5/T + 558780/T^2 - 0.00964T - 4.104 \log T$ | 25 - 350 | Arnorsson et al.(1982 c); Na, K in moles/Kg, T in °K |
| (9.) $\log CO_2 = 37.43 - 73192/T - 11529000/T^2 + 0.15923T - 86.157 \log T$ | 180 - 300 | Arnorsson et al.(1982 c); CO_2 in moles/Kg steam at 1atm; T in °K |

Table 6

* Chemically calculated and measured temperatures at the main production zone in S-Negros

| Well No. | Meas. | Temperatures(°C) | | | | | | | | CO ₂ 9) |
|----------|-------|-------------------|-----|---------|-----|-----|------|-----|-----|-----------------------|
| | | Quartz | | Na-K-Ca | | | Na/K | | | |
| | | 1) | 2) | 3) | 4) | 5)a | 6)a | 7)a | 8)b | |
| OK-7 | 318 | (264) | 269 | 295 | 280 | 300 | -- | 291 | 309 | 296 |
| OK-9D | 291 | (252) | 256 | 273 | 276 | 293 | -- | 284 | 289 | 290 |
| OK-10D | 259 | (256) | 260 | 268 | 269 | 287 | -- | 277 | 277 | 270 |
| OK-12D | 250 | 243 | 245 | 253 | 265 | 281 | -- | 274 | 273 | 289 |
| PN13D | 268 | (260) | 265 | 272 | 281 | 297 | -- | 288 | 289 | 286 |
| PN15D | 274 | (262) | 266 | 276 | 265 | 272 | -- | 265 | 265 | 290 |
| PN17D | 245 | 244 | 247 | 253 | 249 | 249 | 234 | -- | 238 | 284 |
| OK-5 | 264 | (288) | 297 | 298 | 304 | 308 | -- | 298 | 299 | 278 |
| OK-6 | 285 | (261) | 266 | 279 | 279 | 284 | -- | 276 | 279 | 276 |
| OK-8 | 244 | (273) | 279 | 279 | 294 | 300 | -- | 291 | 290 | 276 |
| NJ1D | 243 | 239 | 241 | 249 | 284 | 287 | -- | 279 | 276 | 278 |
| SG-1 | 277 | 246 | 248 | 263 | 278 | 286 | -- | 278 | 280 | 271 |
| SG-2 | 275 | (254) | 258 | 269 | 292 | 298 | -- | 289 | 291 | 280 |

* For description of chemical geothermometers, see Table 5

a Na/K ratio from Na,K analyzed

b Na/K ratio from calculated Na⁺,K⁺ activities

Figures in brackets indicates extrapolated values

Table 7

Table showing mean and standard deviations of chemically calculated temperatures from measured temperatures at main production zones in S-Negros

| Well No. | Quartz | | | Na-K-Ca | Na/K | | | | CO ₂ |
|-----------|--------|------|------|---------|------|-----|------|------|-----------------|
| | 1) | 2) | 3) | 4) | 5)a | 6)a | 7)a | 8)b | 9) |
| OK-7 | -54 | -49 | -23 | -38 | -18 | - | -27 | -9 | -22 |
| OK-9D | -39 | -35 | -18 | -15 | +2 | - | -7 | -2 | -1 |
| OK-10D | -3 | +1 | +9 | +10 | +28 | - | +18 | +18 | +11 |
| OK-12D | -7 | -5 | +3 | +15 | +31 | - | +24 | +23 | +39 |
| PN13D | -8 | -3 | +4 | +13 | +29 | - | +20 | +21 | +18 |
| PN15D | -12 | -8 | +2 | -9 | -2 | - | -9 | -9 | +16 |
| PN17D | -1 | +2 | +8 | +5 | +4 | -11 | 0 | -7 | +39 |
| OK-5 | +24 | +33 | +34 | +40 | +44 | - | +34 | +35 | +14 |
| OK-6 | -24 | -19 | -6 | -6 | -1 | - | -9 | -6 | -9 |
| OK-8 | +29 | +35 | +35 | +50 | +56 | - | +47 | +46 | +32 |
| NJ1D | -4 | -2 | +6 | +41 | +44 | - | +36 | +33 | +35 |
| SG-1 | -31 | -29 | -14 | +1 | +9 | - | +1 | +3 | -6 |
| SG-2 | -21 | -17 | -6 | +15 | +23 | - | +14 | +16 | +5 |
| Mean | | | | | | | | | |
| dev. (°C) | 19.8 | 18.3 | 12.9 | 19.8 | 22.4 | - | 18.9 | 17.5 | 19.0 |
| M(%) | 7.31 | 6.69 | 4.72 | 7.21 | 8.09 | - | 6.90 | 6.46 | 7.01 |
| g(%) | 5.63 | 5.78 | 4.10 | 5.90 | 6.73 | - | 5.30 | 5.11 | 5.06 |

+ Sign refers to degrees C higher than measured temperatures

- Sign refers to degrees C lower than measured temperatures

M mean percent deviation

g standard deviation

fraction of the total silica exists as H_3SiO_4^- , which does not participate in the equilibria with quartz. The various quartz geothermometers (Fournier, 1977; Arnorsson et al., 1982 c) presented in Table 5 assumes no steam loss and are calculated from the deep water silica. The "pressure sensitive silica geothermometer" of Ragnarsdottir and Walther (1982) is calculated from silica concentrations at atmospheric pressure and from the fluid pressure at the depth of inflow. The silica concentrations at sampling pressures were recalculated to 1 atm assuming single stage adiabatic boiling.

The silica geothermometers of Fournier (1977) and Arnorsson et al. (1982 c) generally gives lower values than measured for most wells. This suggests silica being precipitated (probably within the liner and casing walls) at depth due to its relatively fast rate of re-equilibration which could be enhanced by fluid salinity, as readily shown in the high chloride wells OK-7 and OK-9D where the calculated silica geothermometers are lower than measured by 35°C to 45°C. Similarly the lower values of the geothermometers may indicate dilution with cooler fluid which was previously shown to be more pronounced in the Nasuji and Sogongon sector of the field. The quartz geothermometers, on the average, give 19°C lower than measured temperatures for wells SG-2, and SG-1.

The "pressure sensitive silica geothermometer" of Ragnarsdottir and Walther (1982) yields the least average deviations from the measured values (12.9°C). The average deviations of the quartz geothermometers of Fournier (1977) and Arnorsson et al. (1982 c) are 19.8°C and 18.3°C, respectively.

3.2.2 Cation geothermometry

Tables 6 and 7 show that the Na/K geothermometer, on the average, yields temperatures which are higher than measured. The Na-K-Ca geothermometer gives slightly lower

temperatures than the Na-K which can be attributed to the fact that the function Ca/Na is more sensitive to cooling than the Na-K ratio. The above trend could also imply that higher temperatures could still exist, probably greater than $300^{\circ}C$.

The calibration curves of Arnorsson et al. (1982 c) for the Na-K geothermometer are based on low albite-microcline data from Helgeson (1969) and calibrated empirically using analytical values for Na^{+} and K^{+} on one hand and computed activities on the other. When applied to the S-Negros data the calibration curves show a temperature difference of more than $2^{\circ}C$ for almost half of the samples, particularly in the more saline Puhagan fluids. The high salinity and high temperature of the fluid could cause the analytical ratios of Na/K to deviate significantly from the Na^{+}/K^{+} activity ratio due to the higher degree of complexing of sodium to form NaCl species. The use of the Na^{+}/K^{+} activity ratio, which cancels out the complexing process, is regarded the most reliable geothermometer to estimate subsurface temperatures in the S-Negros geothermal field. Well NJ1D, which shows exceptionally high sulfate (155 ppm) is the only Nasuji well whose Na/K temperature as calculated from analytical concentrations and computed activities differ by as much as $5^{\circ}C$. The high sulfate content of the fluid could give higher stability to the KSO_4^{-} -complex relative to $NaSO_4^{-}$, which yields high temperatures when analytical concentrations are used instead of activities, (Arnorsson et al. 1982 c).

Arnorsson et al. (1982 c) empirically calibrated the Na/K geothermometer using the computed activities of Na^{+} and K^{+} and showed this geothermometer to have the least deviation among cation geothermometers. The significant differences between the Na-K temperatures from quartz temperatures supports the concept of precipitation of quartz in saline fluids and dilution with cooler fluid for less saline waters.

3.2.3 CO₂ geothermometry

Arnorsson et al. (1982 c) proposed an empirical CO₂ geothermometer based on data from deep drillholes in Iceland, which is applicable in the temperature range 180-300°C. The lower limit is set by the effect of the salinity of the parent water on the CO₂ concentrations in the steam, while the upper limit is set by the mobility of the gas at higher temperatures.

CO₂ concentrations in steam at sampling pressure was recalculated to 1 atmosphere assuming adiabatic boiling in one stage. Tables 6 and 7 show that the calculated temperatures gave generally higher values than measured by about 19°C which provides evidence for higher temperatures elsewhere, as was also concluded for the Na/K geothermometers.

3.2.4 Chloride-enthalpy mixing models

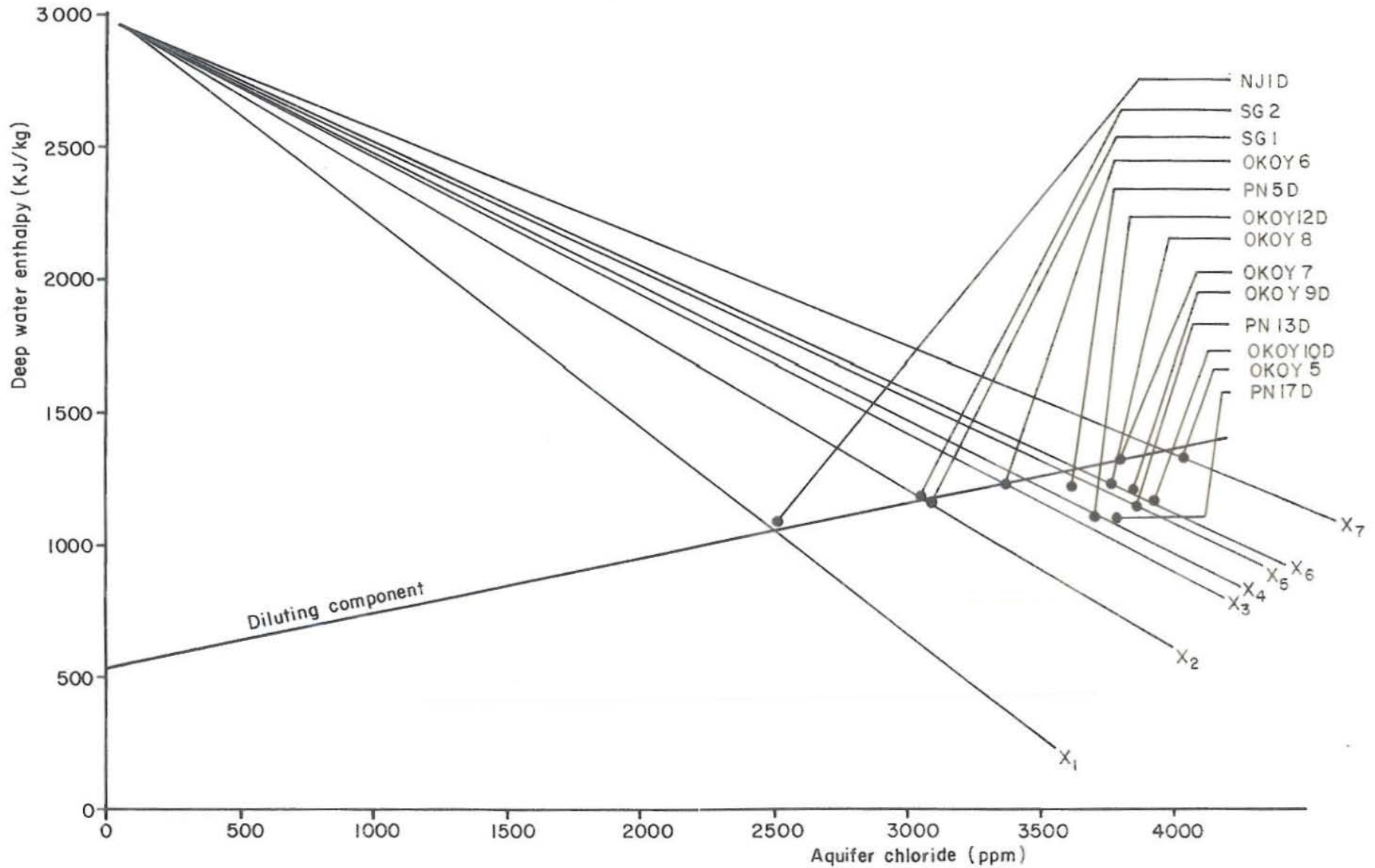
The following factors which have been described earlier might be explained by mixing: a) the discrepancies between the quartz and the Na/K geothermometers, b) non-attainment of overall equilibrium between certain species and cation to proton ratios, and c) linear relationship between silica and chloride.

The chloride-enthalpy mixing model was utilized to test the hypothesis that cold meteoric water recharge from the northwest is the diluting component of the discharges in the Sogongon and some Nasuji wells, and to estimate the temperature of the hot water component. Fig. 14 shows the chloride concentrations plotted against the deep water enthalpy. The deep water enthalpy was taken from the silica calculated temperatures of Ragnarsdottir and Walther (1982), which gave the least deviation from measured values. Deep water chloride values were calculated from analytical results at sampling pressure to the measured



JHD-HSP-9000-OJ
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Figure 14 Deep water enthalpy-aquifer chloride plot



temperatures at the main production zones, after appropriate corrections for excess enthalpies had been made.

Following the procedure described by Fournier (1977), assuming minimal heat gain or loss after mixing, extrapolation of the best fit line (dilution line) yields maximum and minimum temperatures of the hot water component to be 310-300°C. This is slightly lower than the maximum measured bottomhole temperatures of some wells (318°C OK-7, 322°C OK-9D, 317°C PN-13D, 310°C OK-5) and reveals that the hottest wells are still withdrawing from a mixed fluid, but with a significantly lower fraction of the cooling component. Fig. 14 indicates that the diluting component has a negligible chloride content (indicative of meteoric water) yet has an enthalpy of 520 KJ/Kg which corresponds to a liquid water saturation temperature of 120°C. This suggests either a) cold meteoric water picking up heat from the rock before mixing with the hydrothermal fluid, or b) Steam heated ground water rich in carbonate and sulfate. From the absolute sulfate, carbonate, and pH of the discharges of SG-1, SG-2, OK-6, NJ1D, however, cause a) is more likely.

The slightly lower enthalpy yet relatively high chloride of contents of wells OK-8, PN-13D, PN17D, OK-10D, and OK-12D are indicative of conductive cooling. Points designated as X1-X7 would represent variations in enthalpy and chloride contents of the liquid water fraction during progressive steam separation. The use of the above mixing model is believed to give more reliable temperatures than the silica-enthalpy plot since the temperature of the hot water component is close to 300°C (Fournier, 1977).

3.3 Isochemical and isothermal contours

In order to visualize the pattern of temperature distribution, to predict further areas for drilling priorities, and to determine the extent of the reservoir,

selected chemical geothermometers and species were plotted across the field and compared with the isothermal contours at 1000m depth. The isothermal contours were deduced from the estimated stable rock temperatures which were taken after a careful examination and selection of all downhole temperature logs, (Catigtig and Amistoso, 1981). Figures 15-17 show plots of iso T aNa^+/aK^+ , iso T $Si(OH)_4 \cdot 2H_2O$, and iso-chloride. Figure 18 shows the isothermal plots at 1000m depth.

The isochemical contours, more or less, shows a series of broad concentric rings towards the north and west, but exhibit a definite sharpening and distortion within the vicinity of the Puhagan wells. This sharpening trend seems to be directly associated with the Ticala fault (Harper and Clemente, 1982), which confines the flow in a narrow northeasterly direction.

The inherent low temperature and low permeability of well OK-11D results in a similar narrow contour stretch between the bottomholes of OK-5, and OK-9D, and further indicates that the flow from an upwelling fluid is initially restricted within these areas, but then broadens outwards to the west into Nasuji, and progressively into Sogongon. The present author, however, suggests that indications of a broad flow within this vicinity cannot be presented with reasonable certainty. Connection of contour lines between the widely spaced wells (OK-6 to SG-1 and 2, to OK-8) must be treated with extreme caution. It is evident from the plots that an outflow zone is located near wells SG-1, SG-2, and OK-12D. In this area differences in absolute chemistry from the rest of the wells indicate the flow regime to be slow and indulgent in continued water-rock reaction during conductive cooling and/or dilution with cooler fluid.

The difference in absolute values between T $Si(OH)_4 \cdot 2H_2O$ and T aNa^+/K^+ is significant in the Sogongon wells (32-33°C) and some Nasuji wells, (OK-6 13°C, NJ1D 35°C), indicative of

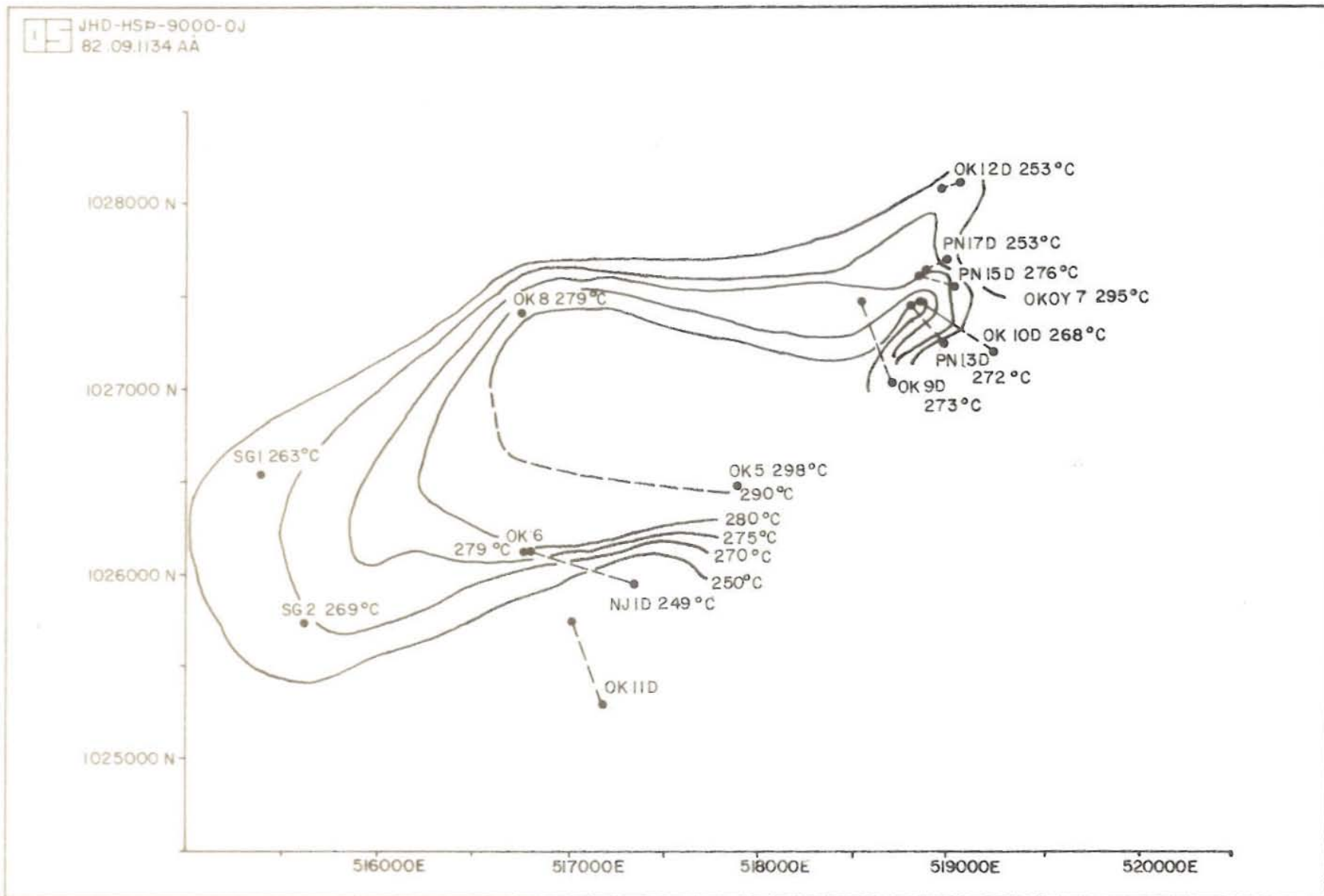


Figure 15. Iso $T_{\text{Si(OH)}_4 \cdot 2\text{H}_2\text{O}}$ map of the Palinpinon field.

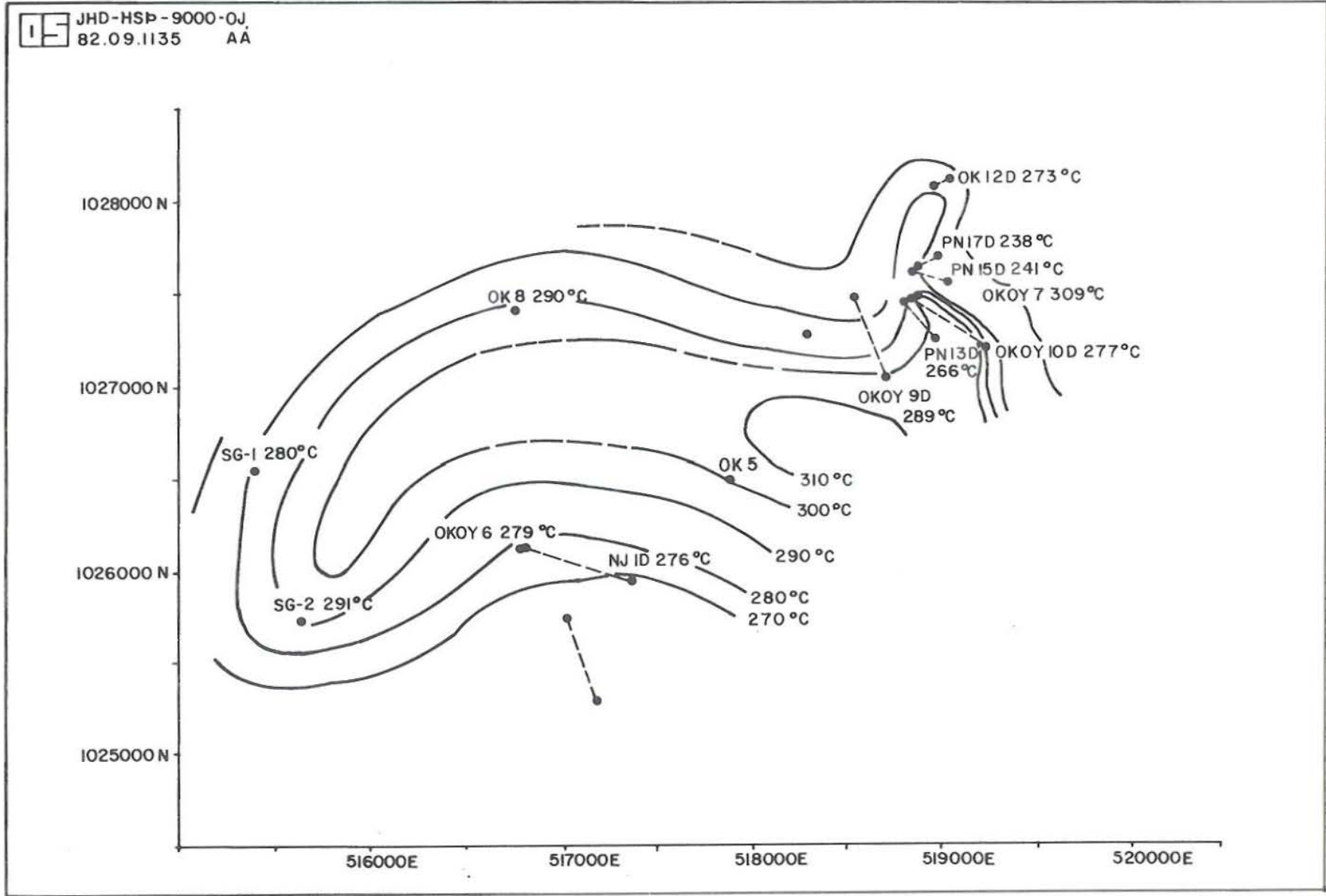


Figure 16. Iso $T_{aNa^+/-aK^+}$ map of the Palinpinon field.

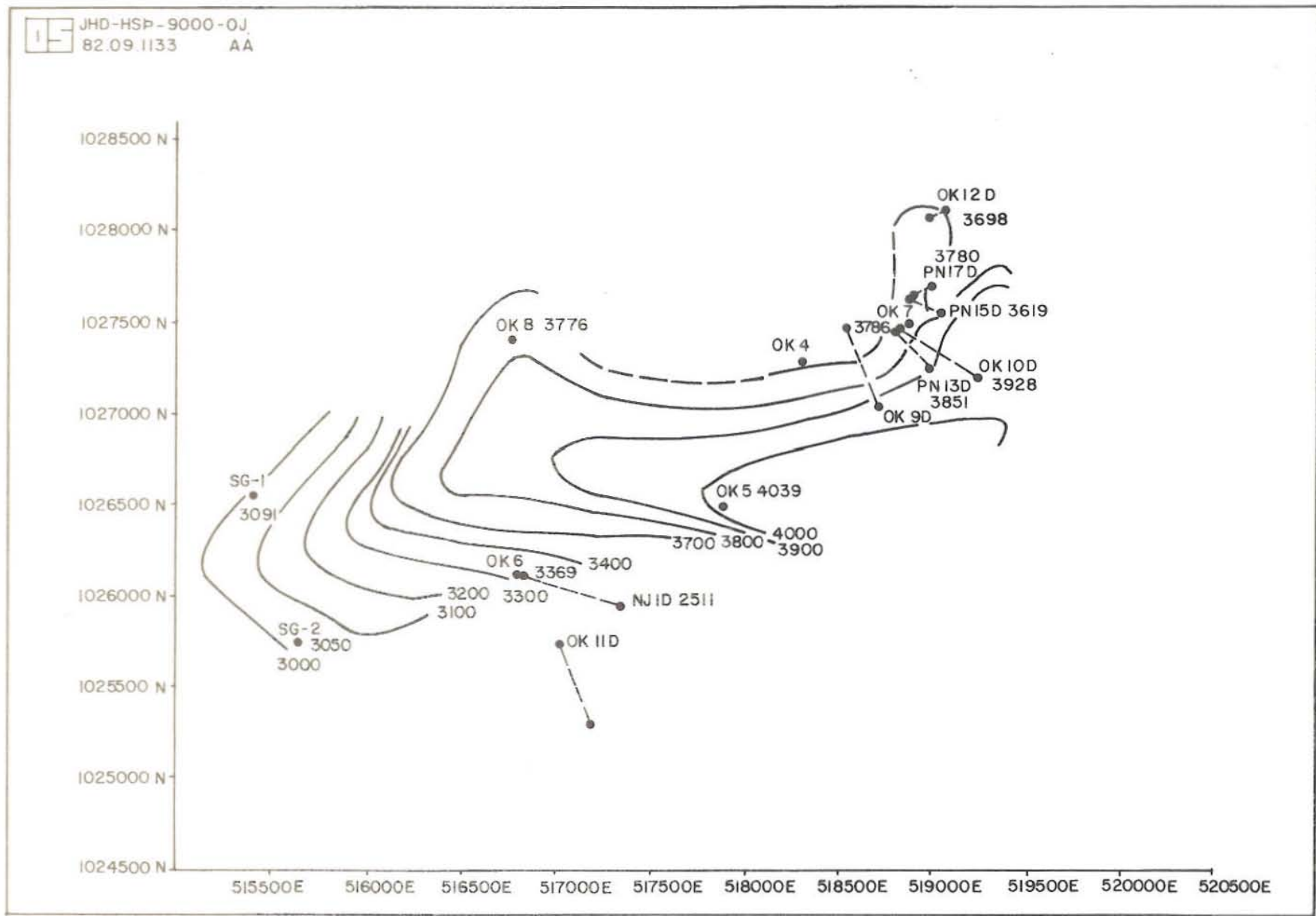


Figure 17. Iso chloride map of the Palinpinon field.

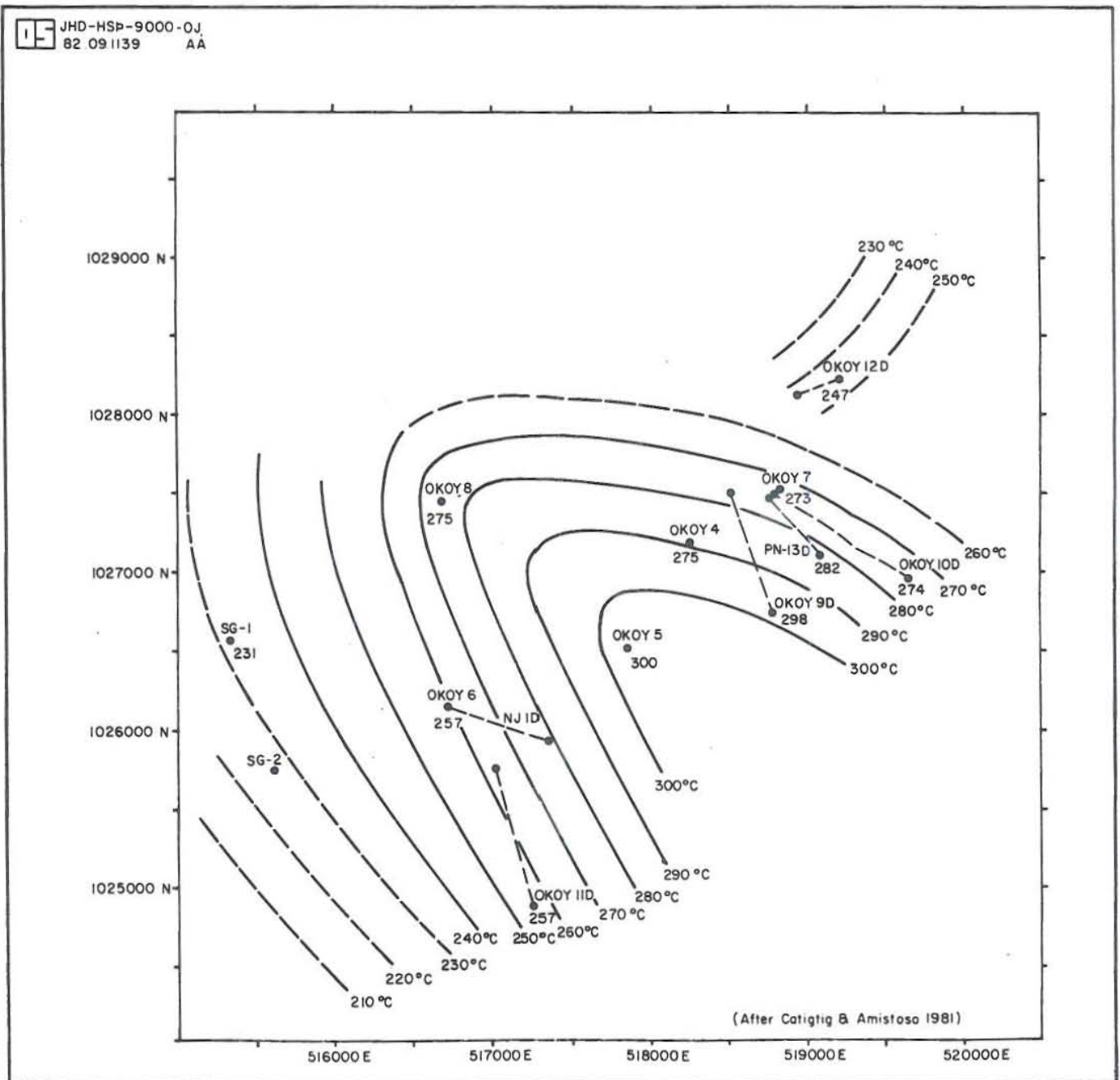


Figure 18 Isotherms at 1000m bmsl of the Palinpinon field.

mixing. This is readily supported by the iso chloride contours (Fig. 17) which show a definite dilution trend, increasing in magnitude from east to west. The absolute deep water chloride values of wells SG-1 and 2, OK-6 and NJ1D are lower (3000-3360 ppm), than the Puhagan wells (3400-3950 ppm). All contours point to the Cuernos volcano as the ultimate heat source of the field, beneath which the upwelling zone is postulated to be at temperatures of over 300°C.

3.4 CONCLUSIONS

a) The complex and iterative calculations involved in the interpretation of the chemical data from the production wells in SNGF made the use of a computer programme necessary.

b) Anhydrite and quartz are shown to be in equilibrium with the deep water for most wells. Calcite saturation, however, appears to be controlled by temperature and affected by mixing with cooler fluid. Potential calcite deposition is not likely to occur, but amorphous silica scaling could be a problem below 150°C reinjection or production separation temperature.

c) The empirically derived relation between aqueous species and temperature, Arnorsson et al. (1982 b) compares well for the silica and bicarbonate species; but the low cation to proton ratios could be due to a different rock leaching property and the effects of dilution with extraneous fluids.

d) For S-Negros fluids, the "pressure sensitive silica geothermometer" of Ragnarsdottir and Walther (1982), and the Na-K geothermometer of Arnorsson et al. (1982 c) using the computed activities of Na⁺ and K⁺, yield the least deviation from measured values of the various geothermometers tested.

e) Chloride - enthalpy mixing models suggest that cold meteoric water recharge from the northwest picks up heat from the rock before mixing with an outflowing hydrothermal fluid, which explains its high enthalpy yet almost zero chloride content. True reservoir temperatures are estimated to be at least 300 °C.

4.0 RECOMMENDATIONS

The following are recommendations for further geochemical studies based upon the results which have emerged during the evaluation of the chemical data presented in this report:

(a) A representative sample from each well should be analyzed for iron, aluminum, and flouride. Most of the minerals found in SNGF contain these elements and analysis of these elements therefore would allow more solute/mineral equilibria studies.

(b) To facilitate and speed up studies of mineral deposition and scaling, a computer programme such as that developed in Iceland should be utilized. This is highly relevant considering the accelerated discharge testing of wells and especially for the programmed reinjection wells in SNGF.

(c) The use of gas geothermometers should be given equal attention to that of the quartz and the cation geothermometers. This could much aid in predicting subsurface temperatures and in identifying flow processes in the field, especially in areas where intense steaming grounds are located, e.g. Kaipohan and Baslay de Dauin.

(d) Variations of chemistry with time, especially for wells on long term discharge tests, should be examined carefully to detect the response of the reservoir to exploitation. This may be detectable only after years of continuous discharge.

(e) A revision of certain aspects in sampling procedures and analytical methods are still necessary. For more details of this recommendation, see manual on methods of collection and analysis of geothermal fluids by the Virkir Consulting Group Ltd.

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1982-11-12 OLIVER

GEOCHEMISTRY OF WELL SONGON-1

XXXXXXXXXX811214XXXX SONGONGON-1

S.NEGROS, PHILIPPINES SONGONGON

PROGRAM WATCH1.

WATER SAMPLE (PPM)

STEAM SAMPLE

| | | | | | |
|-------------|-----------|----------------------------|-----------------|-------------------------|---------------------------|
| PH/DEG.C | 7.30/25.0 | GAS (VOL.%) | REFERENCE TEMP. | DEGREES C | 277.0 (MEASURED) |
| SI02 | 626.00 | CO2 | 0.00 | | |
| NA | 2250.00 | H2S | 0.00 | SAMPLING PRESSURE | BARS ABS. 6.9 |
| K | 457.00 | H2 | 0.00 | DISCHARGE ENTHALPY | KJOUL/KG 1,222 (MEASURED) |
| CA | 34.40 | O2 | 0.00 | DISCHARGE | KG/SEC. 35.8 |
| MG | 0.060 | CH4 | 0.00 | | |
| CO2 | 32.10 | N2 | 0.00 | MEASURED TEMPERATURE | DEGREES C 0.0 |
| SO4 | 39.80 | | | RESISTIVITY/TEMP. | OHMM/DEG.C 0.0/ 0.0 |
| H2S | 3.07 | | | EH/TEMP. | MV/DEG.C 0.000/ 0.0 |
| CL | 4148.00 | | | | |
| F | 0.00 | LITERS GAS PER KG | | | |
| DISS.SOLIDS | 0.00 | CONDENSATE/DEG.C | 5.23/25.0 | MEASURED DOWNHOLE TEMP. | FLUID INFLOW |
| AL | 0.0000 | | | DEGREES C/METERS | DEPTH (METERS) |
| B | 51.1000 | CONDENSATE (PPM) | | 0.0 | 0.0 |
| FE | 0.0000 | PH/DEG.C | 0.00/ 0.0 | 0.0 | 0.0 |
| NH3 | 0.0000 | CO2 | 0.00 | 0.0 | 0.0 |
| | | H2S | 0.00 | 0.0 | 0.0 |
| | | NA | 0.00 | 0.0 | 0.0 |
| | | | | 0.0 | 0.0 |
| | | | | 0.0 | 0.0 |
| | | | | 0.0 | 0.0 |
| | | | | 0.0 | 0.0 |
| | | CONDENSATE WITH NAOH (PPM) | | 0.0 | 0.0 |
| | | CO2 | 9848.70 | 0.0 | 0.0 |
| | | H2S | 154.40 | 0.0 | 0.0 |

IONIC STRENGTH = 0.11594

IONIC BALANCE : CATIONS (MOL.EQ.)0.11101650
ANIONS (MOL.EQ.)0.11839960
DIFFERENCE (%) -6.44

DEEP WATER (PPM)

DEEP STEAM (PPM)

GAS PRESSURES (BARS ABS.)

| | | | | | | | |
|---------|---------|-----|---------|-----------------|-----------|-------|-----------|
| SI02 | 466.63 | CO2 | 2305.96 | CO2 | 195879.34 | CO2 | 0.492E+01 |
| NA | 1677.05 | H2S | 40.53 | H2S | 1043.08 | H2S | 0.338E-01 |
| K | 340.61 | H2 | 0.00 | H2 | 0.00 | H2 | 0.000E+00 |
| CA | 25.64 | O2 | 0.00 | O2 | 0.00 | O2 | 0.000E+00 |
| MG | 0.045 | CH4 | 0.00 | CH4 | 0.00 | CH4 | 0.000E+00 |
| SO4 | 29.66 | N2 | 0.00 | N2 | 0.00 | N2 | 0.000E+00 |
| CL | 3091.47 | NH3 | 0.00 | NH3 | 0.00 | NH3 | 0.000E+00 |
| F | 0.00 | | | | | H2O | 0.613E+02 |
| DISS.S. | 0.00 | | | | | TOTAL | 0.663E+02 |
| AL | 0.0000 | | | | | | |
| B | 38.0841 | | | H2O (%) | 0.12 | | |
| FE | 0.0000 | | | BOILING PORTION | 0.25 | | |

GAS SOLUBILITY MULTIPLYING FACTOR : 1.00

ACTIVITY COEFFICIENTS IN DEEP WATER

| | | | | | | | |
|----------|-------|---------|-------|-----------|-------|----------|-------|
| H+ | 0.698 | KSO4- | 0.639 | FE++ | 0.179 | FECL+ | 0.609 |
| OH- | 0.598 | F- | 0.598 | FE+++ | 0.040 | AL+++ | 0.040 |
| H3SiO4- | 0.609 | CL- | 0.585 | FE(OH) | 0.631 | ALOH++ | 0.167 |
| H2SiO4-- | 0.167 | NA+ | 0.609 | FE(OH)3- | 0.631 | AL(OH)2+ | 0.639 |
| H2BO3- | 0.572 | K+ | 0.585 | FE(OH)4-- | 0.159 | AL(OH)4- | 0.621 |
| HCO3- | 0.609 | CA++ | 0.179 | FE(OH)2+ | 0.639 | ALSO4+ | 0.621 |
| CO3-- | 0.148 | MG++ | 0.219 | FE(OH)4- | 0.639 | ALF++ | 0.167 |
| HS- | 0.598 | CAHCO3+ | 0.651 | FESO4+ | 0.631 | ALF2+ | 0.639 |
| S-- | 0.159 | MGHCO3+ | 0.609 | FECL++ | 0.159 | ALF4- | 0.621 |
| H2SO4- | 0.621 | CAOH+ | 0.651 | FECL2+ | 0.631 | ALF5-- | 0.148 |
| SO4-- | 0.138 | MGOH+ | 0.660 | FECL4- | 0.609 | ALF6--- | 0.014 |
| NaSO4- | 0.639 | NH4+ | 0.572 | | | | |

CHEMICAL COMPONENTS IN DEEP WATER (PPM AND LOG MOLE)

| | | | | | | | | |
|-----------|---------|---------|-----------|--------|---------|-----------|------|-------|
| H+ (ACT.) | 0.00 | -5.690 | MG++ | 0.04 | -5.790 | FE(OH)3 | 0.00 | 0.000 |
| OH- | 0.13 | -5.100 | NaCl | 398.81 | -2.166 | FE(OH)4- | 0.00 | 0.000 |
| H4SiO4 | 745.95 | -2.110 | KCl | 34.14 | -3.339 | FECL+ | 0.00 | 0.000 |
| H3SiO4- | 0.31 | -5.487 | NaSO4- | 6.65 | -4.253 | FECL2 | 0.00 | 0.000 |
| H2SiO4-- | 0.00 | -10.902 | KSO4- | 8.72 | -4.190 | FECL++ | 0.00 | 0.000 |
| NAH3SiO4 | 0.19 | -5.786 | CASO4 | 2.96 | -4.662 | FECL2+ | 0.00 | 0.000 |
| H3BO3 | 217.70 | -2.453 | MGSO4 | 0.02 | -6.895 | FECL3 | 0.00 | 0.000 |
| H2BO3- | 0.12 | -5.688 | CACO3 | 0.02 | -6.627 | FECL4- | 0.00 | 0.000 |
| H2CO3 | 3215.52 | -1.285 | MGCO3 | 0.00 | -10.488 | FESO4 | 0.00 | 0.000 |
| HCO3- | 27.05 | -3.353 | CAHCO3+ | 11.58 | -3.941 | FESO4+ | 0.00 | 0.000 |
| CO3-- | 0.00 | -8.552 | MGHCO3+ | 0.00 | -7.597 | AL+++ | 0.00 | 0.000 |
| H2S | 39.92 | -2.931 | CAOH+ | 0.02 | -6.440 | ALOH++ | 0.00 | 0.000 |
| HS- | 0.59 | -4.751 | MGOH+ | 0.00 | -7.190 | AL(OH)2+ | 0.00 | 0.000 |
| S-- | 0.00 | -13.289 | NH4OH | 0.00 | 0.000 | AL(OH)3 | 0.00 | 0.000 |
| H2SO4 | 0.00 | -10.950 | NH4+ | 0.00 | 0.000 | AL(OH)4- | 0.00 | 0.000 |
| H2SO4- | 3.03 | -4.506 | FE++ | 0.00 | 0.000 | ALSO4+ | 0.00 | 0.000 |
| SO4-- | 13.01 | -3.868 | FE+++ | 0.00 | 0.000 | AL(SO4)2- | 0.00 | 0.000 |
| HF | 0.00 | 0.000 | FE(OH) | 0.00 | 0.000 | ALF++ | 0.00 | 0.000 |
| F- | 0.00 | 0.000 | FE(OH)2 | 0.00 | 0.000 | ALF2+ | 0.00 | 0.000 |
| CL- | 2833.31 | -1.097 | FE(OH)3- | 0.00 | 0.000 | ALF3 | 0.00 | 0.000 |
| NA+ | 1518.83 | -1.180 | FE(OH)4-- | 0.00 | 0.000 | ALF4- | 0.00 | 0.000 |
| K+ | 320.18 | -2.087 | FE(OH)++ | 0.00 | 0.000 | ALF5-- | 0.00 | 0.000 |
| CA++ | 20.15 | -3.299 | FE(OH)2+ | 0.00 | 0.000 | ALF6--- | 0.00 | 0.000 |

IONIC STRENGTH = 0.07866 IONIC BALANCE : CATIONS (MOL.EQ.) 0.07537772
 ANIONS (MOL.EQ.) 0.08066924
 DIFFERENCE (%) -6.78

CHEMICAL GEOTHERMOMETERS DEGREES C 1000/T DEGREES KELVIN = 1.82

QUARTZ 246.3
 CHALCEDONY 999.9
 NAK 288.3

OXIDATION POTENTIAL (VOLTS) : EH H2S= -0.486 EH CH4= 99.999 EH H2= 99.999 EH NH3= 99.999

LOG SOLUBILITY PRODUCTS OF MINERALS IN DEEP WATER

| | TEOR. | CALC. | | TEOR. | CALC. | | TEOR. | CALC. |
|-------------|---------|--------|--------------|---------|---------|-------------|---------|--------|
| ADULARIA | -14.374 | 99.999 | ALBITE LOW | -13.947 | 99.999 | ANALCIME | -11.599 | 99.999 |
| ANHYDRITE | -8.611 | -8.774 | CALCITE | -13.481 | -13.426 | CHALCEDONY | -1.891 | -2.110 |
| MG-CHLORITE | -87.175 | 99.999 | FLUORITE | -11.109 | 99.999 | GOETHITE | 3.805 | 99.999 |
| LAUMONTITE | -24.818 | 99.999 | MICROCLINE | -14.977 | 99.999 | MAGNETITE | -14.275 | 99.999 |
| CA-MONTMOR. | -72.578 | 99.999 | K-MONTMOR. | -33.814 | 99.999 | MG-MONTMOR. | -74.086 | 99.999 |
| NA-MONTMOR. | -34.099 | 99.999 | MUSCOVITE | -17.842 | 99.999 | PREHNITE | -38.313 | 99.999 |
| PYRRHOTITE | -5.596 | 99.999 | PYRITE | -21.879 | 99.999 | QUARTZ | -1.996 | -2.110 |
| WAIKAKITE | -25.119 | 99.999 | WOLLASTONITE | 7.047 | 5.225 | ZOISITE | -39.262 | 99.999 |
| EPIDOTE | -39.107 | 99.999 | NARCASITE | -5.949 | 99.999 | | | |

ORKUSTOFNUN JHD
1982-11-12 OLIVER

GEOCHEMISTRY OF WELL SOGONGON-2

XXXXXXXXXX820118XXXX SOGONGON-2

S.NEGROS,PHILIPPINES SOGONGON

PROGRAM WATCH1.

WATER SAMPLE (PPM)

STEAM SAMPLE

| | | | | | |
|----------|-----------|-------------|-----------------|----------------------|---------------------------|
| PH/DEG.C | 5.91/25.0 | GAS (VOL.%) | REFERENCE TEMP. | DEGREES C | 275.0 (MEASURED) |
| SI02 | 660.00 | CO2 | 0.00 | | |
| NA | 2139.00 | H2S | 0.00 | SAMPLING PRESSURE | BARS ABS. 9.2 |
| K | 482.00 | H2 | 0.00 | DISCHARGE ENTHALPY | MJ/OL/KG 1.280 (MEASURED) |
| CA | 21.40 | O2 | 0.00 | DISCHARGE | KG/SEC. 14.9 |
| MG | 0.060 | CH4 | 0.00 | | |
| CO2 | 25.80 | N2 | 0.00 | MEASURED TEMPERATURE | DEGREES C 0.0 |
| SO4 | 23.20 | | | RESISTIVITY/TEMP. | OHMM/DEG.C 0.0/ 0.0 |
| H2S | 3.58 | | | EH/TEMP. | MV/DEG.C 0.000/ 0.0 |
| CL | 3957.00 | | | | |

F 0.00
DISS.SOLIDS 0.00
AL 0.0000
B 48.6000
FE 0.0000
NH3 0.0000

LITERS GAS PER KG
CONDENSATE/DEG.C 7.97/25.0

CONDENSATE (PPM)
PH/DEG.C 0.00/ 0.0
CO2 0.00
H2S 0.00
NA 0.00

MEASURED DOWNHOLE TEMP.
DEGREES C/MEYERS

0.0 0.0 0.0
0.0 0.0 0.0
0.0 0.0 0.0
0.0 0.0 0.0
0.0 0.0 0.0
0.0 0.0 0.0
CONDENSATE WITH NACH (PPM)
CO2 14690.90
H2S 189.60

IONIC STRENGTH = 0.10991 IONIC BALANCE :
CATIONS (MOL.EQ.)0.10623719
ANIONS (MOL.EQ.)0.11206885
DIFFERENCE (%) -5.34

DEEP WATER (PPM)

DEEP STEAM (PPM)

GAS PRESSURES (BARS ABS.)

| | | | | | | | |
|---------|---------|-----|--------|-----------------|----------|-------|-----------|
| SI02 | 508.89 | CO2 | 786.80 | CO2 | 70179.68 | CO2 | 0.171E+01 |
| NA | 1649.14 | H2S | 24.29 | H2S | 657.35 | H2S | 0.207E-01 |
| K | 371.60 | H2 | 0.00 | H2 | 0.00 | H2 | 0.000E+00 |
| CA | 16.50 | O2 | 0.00 | O2 | 0.00 | O2 | 0.000E+00 |
| MG | 0.046 | CH4 | 0.00 | CH4 | 0.00 | CH4 | 0.000E+00 |
| SO4 | 17.89 | N2 | 0.00 | N2 | 0.00 | N2 | 0.000E+00 |
| CL | 3050.53 | NH3 | 0.00 | NH3 | 0.00 | NH3 | 0.000E+00 |
| F | 0.00 | | | | | H2O | 0.595E+02 |
| DISS.S. | 0.00 | | | | | TOTAL | 0.612E+02 |
| AL | 0.0000 | | | | | | |
| B | 37.4664 | | | H2O (%) | 4.47 | | |
| FE | 0.0000 | | | BOILING PORTION | 0.23 | | |

GAS SOLUBILITY MULTIPLYING FACTOR : 1.00

ACTIVITY COEFFICIENTS IN DEEP WATER

| | | | | | | | |
|----------|-------|---------|-------|-----------|-------|-----------|-------|
| H+ | 0.702 | KS04- | 0.643 | FE++ | 0.183 | FECL+ | 0.614 |
| OH- | 0.602 | F- | 0.602 | FE+++ | 0.041 | AL+++ | 0.041 |
| H3SI04- | 0.614 | CL- | 0.590 | FE0H+ | 0.635 | AL0H++ | 0.171 |
| H2SI04-- | 0.171 | NA+ | 0.614 | FE(OH)3- | 0.635 | AL(OH)2+ | 0.643 |
| H2B03- | 0.576 | K+ | 0.590 | FE(OH)4-- | 0.163 | AL(OH)4- | 0.625 |
| HCO3- | 0.614 | CA++ | 0.183 | FE0H++ | 0.163 | ALSO4+ | 0.625 |
| CO3-- | 0.152 | MG++ | 0.223 | FE(OH)2+ | 0.643 | AL(SO4)2- | 0.625 |
| HS- | 0.602 | CAHCO3+ | 0.654 | FE(OH)4- | 0.643 | ALF++ | 0.171 |
| S-- | 0.163 | MGHCO3+ | 0.614 | FES04+ | 0.635 | ALF2+ | 0.643 |
| H2SO4- | 0.625 | CAOH+ | 0.654 | FECL++ | 0.163 | ALF4- | 0.625 |
| SO4-- | 0.142 | MGOH+ | 0.663 | FECL2+ | 0.635 | ALF5-- | 0.152 |
| NAS04- | 0.643 | NH4+ | 0.576 | FECL4- | 0.614 | ALF6--- | 0.014 |

CHEMICAL COMPONENTS IN DEEP WATER (PPM AND LOG MOLE)

| | | | | | | | | |
|-----------|---------|---------|-----------|--------|---------|-----------|------|-------|
| H+ (ACT.) | 0.00 | -5.526 | MG++ | 0.04 | -5.753 | FE(OH)3 | 0.00 | 0.000 |
| OH- | 0.09 | -5.269 | NACL | 370.25 | -2.198 | FE(OH)4- | 0.00 | 0.000 |
| H4SI04 | 813.66 | -2.072 | KCL | 35.80 | -3.319 | FECL+ | 0.00 | 0.000 |
| H3SI04- | 0.24 | -5.600 | NAS04- | 3.83 | -4.492 | FECL2 | 0.00 | 0.000 |
| H2SI04-- | 0.00 | -11.169 | KS04- | 5.43 | -4.396 | FECL++ | 0.00 | 0.000 |
| NAH3SI04 | 0.15 | -5.902 | CAS04 | 1.31 | -5.016 | FECL2+ | 0.00 | 0.000 |
| H3B03 | 214.21 | -2.460 | MGS04 | 0.01 | -7.081 | FECL3 | 0.00 | 0.000 |
| H2B03- | 0.08 | -5.857 | CAC03 | 0.00 | -7.523 | FECL4- | 0.00 | 0.000 |
| H2CO3 | 1100.87 | -1.751 | MGC03 | 0.00 | -11.216 | FES04 | 0.00 | 0.000 |
| HCO3- | 6.66 | -3.962 | CAHCO3+ | 2.14 | -4.675 | FES04+ | 0.00 | 0.000 |
| CO3-- | 0.00 | -9.310 | MGHCO3+ | 0.00 | -8.172 | AL+++ | 0.00 | 0.000 |
| H2S | 24.02 | -3.152 | CAOH+ | 0.01 | -6.734 | AL0H++ | 0.00 | 0.000 |
| HS- | 0.25 | -5.114 | MGOH+ | 0.00 | -7.346 | AL(OH)2+ | 0.00 | 0.000 |
| S-- | 0.00 | -13.834 | NH4OH | 0.00 | 0.000 | AL(OH)3 | 0.00 | 0.000 |
| H2SO4 | 0.00 | -10.886 | NH4+ | 0.00 | 0.000 | AL(OH)4- | 0.00 | 0.000 |
| H2SO4- | 2.44 | -4.599 | FE++ | 0.00 | 0.000 | ALSO4+ | 0.00 | 0.000 |
| SO4-- | 7.59 | -4.103 | FE+++ | 0.00 | 0.000 | AL(SO4)2- | 0.00 | 0.000 |
| HF | 0.00 | 0.000 | FE0H+ | 0.00 | 0.000 | ALF++ | 0.00 | 0.000 |
| F- | 0.00 | 0.000 | FE(OH)2 | 0.00 | 0.000 | ALF2+ | 0.00 | 0.000 |
| CL- | 2808.91 | -1.101 | FE(OH)3- | 0.00 | 0.000 | ALF3 | 0.00 | 0.000 |
| NA+ | 1502.71 | -1.185 | FE(OH)4-- | 0.00 | 0.000 | ALF4- | 0.00 | 0.000 |
| K+ | 351.25 | -2.047 | FE(OH)++ | 0.00 | 0.000 | ALF5-- | 0.00 | 0.000 |
| CA++ | 15.26 | -3.419 | FE(OH)2+ | 0.00 | 0.000 | ALF6--- | 0.00 | 0.000 |

IONIC STRENGTH = 0.07779 IONIC BALANCE : CATIONS (MOL.EQ.)0.07513352
 ANIONS (MOL.EQ.)0.07951997
 DIFFERENCE (%) -5.67

CHEMICAL GEOTHERMOMETERS DEGREES C 1000/T DEGREES KELVIN = 1.82

QUARTZ 257.8
 CHALCEDONY 999.9
 NAK 304.1

OXIDATION POTENTIAL (VOLTS) : EH H2S= -0.462 EH CH4= 99.999 EH H2= 99.999 EH NH3= 99.999

LOG SOLUBILITY PRODUCTS OF MINERALS IN DEEP WATER

| | TEOR. | CALC. | | TEOR. | CALC. | | TEOR. | CALC. |
|-------------|---------|--------|--------------|---------|---------|-------------|---------|--------|
| ADULARIA | -14.373 | 99.999 | ALBITE LOW | -13.944 | 99.999 | ANALCINE | -11.591 | 99.999 |
| ANHYDRITE | -8.573 | -9.108 | CALCITE | -13.427 | -14.284 | CHALCEDONY | -1.899 | -2.072 |
| NG-CHLORITE | -86.993 | 99.999 | FLUORITE | -11.097 | 99.999 | GOETHITE | 3.689 | 99.999 |
| LAUMONTITE | -24.794 | 99.999 | MICROCLINE | -14.980 | 99.999 | MAGNETITE | -14.476 | 99.999 |
| CA-MONTMOR. | -72.578 | 99.999 | K-MONTMOR. | -33.824 | 99.999 | MG-MONTMOR. | -74.086 | 99.999 |
| NA-MONTMOR. | -34.108 | 99.999 | MUSCOVITE | -17.841 | 99.999 | PREHNITE | -38.232 | 99.999 |
| PYRRHOTITE | -6.754 | 99.999 | PYRITE | -23.333 | 99.999 | QUARTZ | -2.004 | -2.072 |
| WAIKAKITE | -25.072 | 99.999 | WOLLASTONITE | 7.078 | 4.823 | ZOISITE | -39.157 | 99.999 |
| EPIDOTE | -38.947 | 99.999 | MARCASITE | -7.344 | 99.999 | | | |

ORKUSTOFNUN JHD GEOCHEMISTRY OF WELL OKOY-5
 1982-11-12 OLIVER

XXXXXXXXXX820307XXXX OKOY-5

S.NEGROS,PHILIPPINES BALAS-BALAS

PROGRAM WATCH3. TEMPERATURE FIXING STEAM LOSS 100.0 DEGREES C

WATER SAMPLE (PPM)

STEAM SAMPLE

| | | | | | |
|-------------|-----------|----------------------------|-------------------------|------------------|-----------------------------|
| PH/DEG.C | 7.22/25.0 | GAS (VOL.%) | REFERENCE TEMP. | DEGREES C | 264.0 (ARBITRARY) |
| SI02 | 1073.00 | CO2 | | | |
| NA | 3347.00 | H2S | SAMPLING PRESSURE | BARS ABS. | |
| K | 820.00 | H2 | DISCHARGE ENTHALPY | HJOUL/KG | |
| CA | 35.10 | O2 | DISCHARGE | KG/SEC. | 29.0 |
| MG | 0.040 | CH4 | | | |
| CO2 | 68.50 | N2 | MEASURED TEMPERATURE | DEGREES C | 0.0 |
| SO4 | 61.50 | | RESISTIVITY/TEMP. | OHMM/DEG.C | 0.0/ 0.0 |
| H2S | 9.88 | | EH/TEMP. | MV/DEG.C | 0.000/ 0.0 |
| CL | 5993.00 | | | | |
| F | 0.00 | LITERS GAS PER KG | | | |
| DISS.SOLIDS | 0.00 | CONDENSATE/DEG.C | MEASURED DOWNHOLE TEMP. | DEGREES C/METERS | FLUID INFLOW DEPTH (METERS) |
| AL | 0.0000 | | | | |
| B | 73.0000 | CONDENSATE (PPM) | 0.0 | 0.0 | 0.0 |
| FE | 0.0000 | PH/DEG.C | 0.0 | 0.0 | 0.0 |
| NH3 | 0.0000 | CO2 | 0.0 | 0.0 | 0.0 |
| | | H2S | 0.0 | 0.0 | 0.0 |
| | | NA | 0.0 | 0.0 | 0.0 |
| | | | 0.0 | 0.0 | 0.0 |
| | | | 0.0 | 0.0 | 0.0 |
| | | | 0.0 | 0.0 | 0.0 |
| | | CONDENSATE WITH NAOH (PPM) | 0.0 | 0.0 | 0.0 |
| | | CO2 | 0.0 | 0.0 | 0.0 |
| | | H2S | 0.0 | 0.0 | 0.0 |

IONIC STRENGTH = 0.17111 IONIC BALANCE : CATIONS (MOL.EQ.)0.16779573
 ANIONS (MOL.EQ.)0.17154194
 DIFFERENCE (%) -2.21

| DEEP WATER (PPM) | | DEEP STEAM (PPM) | | GAS PRESSURES (BARS ABS.) | |
|------------------|---------|------------------|----------|---------------------------|-----------|
| SI02 | 723.34 | CO2 | 27568.29 | CO2 | 0.641E+02 |
| NA | 2256.11 | H2S | 1867.72 | H2S | 0.171E+01 |
| K | 552.71 | H2 | 0.00 | H2 | 0.000E+00 |
| CA | 23.66 | O2 | 0.00 | O2 | 0.000E+00 |
| MG | 0.027 | CH4 | 0.00 | CH4 | 0.000E+00 |
| SO4 | 41.45 | N2 | 0.00 | N2 | 0.000E+00 |
| CL | 4039.36 | NH3 | 0.00 | NH3 | 0.000E+00 |
| F | 0.00 | | | H2O | 0.501E+02 |
| DISS.S. | 0.00 | | | TOTAL | 0.116E+03 |
| AL | 0.0000 | | | | |
| B | 49.2025 | | | H2O (%) | 0.00 |
| FE | 0.0000 | | | BOILING PORTION | 0.00 |

GAS SOLUBILITY MULTIPLYING FACTOR : 1.00

ACTIVITY COEFFICIENTS IN DEEP WATER

| | | | | | | | |
|----------|-------|---------|-------|-----------|-------|-----------|-------|
| H+ | 0.691 | KSO4- | 0.626 | FE++ | 0.169 | FECL+ | 0.593 |
| OH- | 0.580 | F- | 0.580 | FE+++ | 0.037 | AL+++ | 0.037 |
| H3SiO4- | 0.593 | CL- | 0.566 | FE(OH) | 0.617 | ALOH++ | 0.156 |
| H2SiO4-- | 0.156 | NA+ | 0.593 | FE(OH)3- | 0.617 | AL(OH)2+ | 0.626 |
| H2BO3- | 0.551 | K+ | 0.566 | FE(OH)4-- | 0.148 | AL(OH)4- | 0.606 |
| HCO3- | 0.593 | CA++ | 0.169 | FE(OH)++ | 0.148 | ALSO4+ | 0.606 |
| CO3-- | 0.137 | MG++ | 0.211 | FE(OH)2+ | 0.626 | AL(SO4)2- | 0.606 |
| HS- | 0.580 | CAHCO3+ | 0.639 | FE(OH)4- | 0.626 | ALF++ | 0.156 |
| S-- | 0.148 | MGHCO3+ | 0.593 | FES04+ | 0.617 | ALF2+ | 0.626 |
| H2SO4- | 0.606 | CAOH+ | 0.639 | FECL++ | 0.148 | ALF4- | 0.606 |
| SO4-- | 0.126 | MGOH+ | 0.649 | FECL2+ | 0.617 | ALF5-- | 0.137 |
| NAS04- | 0.626 | NH4+ | 0.551 | FECL4- | 0.593 | ALF6--- | 0.011 |

CHEMICAL COMPONENTS IN DEEP WATER (PPM AND LOG MOLE)

| | | | | | | | | |
|-----------|----------|---------|-----------|--------|---------|-----------|------|-------|
| H+ (ACT.) | 0.02 | -4.732 | MG++ | 0.02 | -5.991 | FE(OH)3 | 0.00 | 0.000 |
| OH- | 0.01 | -6.056 | NACL | 445.01 | -2.118 | FE(OH)4- | 0.00 | 0.000 |
| H4SiO4 | 1156.94 | -1.920 | KCL | 50.00 | -3.174 | FECL+ | 0.00 | 0.000 |
| H3SiO4- | 0.07 | -6.138 | NAS04- | 6.76 | -4.246 | FECL2 | 0.00 | 0.000 |
| H2SiO4-- | 0.00 | -12.381 | KSO4- | 9.00 | -4.176 | FECL++ | 0.00 | 0.000 |
| NAH3SiO4 | 0.05 | -6.350 | CASO4 | 1.62 | -4.925 | FECL2+ | 0.00 | 0.000 |
| H3BO3 | 281.40 | -2.342 | MGS04 | 0.01 | -7.221 | FECL3 | 0.00 | 0.000 |
| H2BO3- | 0.02 | -6.480 | CACO3 | 0.00 | -7.488 | FECL4- | 0.00 | 0.000 |
| H2CO3 | 38793.44 | -0.204 | MGCO3 | 0.00 | -11.402 | FES04 | 0.00 | 0.000 |
| HCO3- | 52.90 | -3.062 | CAHCO3+ | 14.62 | -3.840 | FES04+ | 0.00 | 0.000 |
| CO3-- | 0.00 | -9.045 | MGHCO3+ | 0.00 | -7.599 | AL+++ | 0.00 | 0.000 |
| H2S | 1863.09 | -1.262 | CAOH+ | 0.00 | -7.589 | ALOH++ | 0.00 | 0.000 |
| HS- | 4.49 | -3.867 | MGOH+ | 0.00 | -8.582 | AL(OH)2+ | 0.00 | 0.000 |
| S-- | 0.00 | -13.411 | NH4OH | 0.00 | 0.000 | AL(OH)3 | 0.00 | 0.000 |
| H2SO4 | 0.00 | -9.358 | NH4+ | 0.00 | 0.000 | AL(OH)4- | 0.00 | 0.000 |
| H2SO4- | 15.73 | -3.790 | FE++ | 0.00 | 0.000 | ALSO4+ | 0.00 | 0.000 |
| SO4-- | 12.89 | -3.872 | FE+++ | 0.00 | 0.000 | AL(SO4)2- | 0.00 | 0.000 |
| HF | 0.00 | 0.000 | FE(OH) | 0.00 | 0.000 | ALF++ | 0.00 | 0.000 |
| F- | 0.00 | 0.000 | FE(OH)2 | 0.00 | 0.000 | ALF2+ | 0.00 | 0.000 |
| CL- | 3745.64 | -0.976 | FE(OH)3- | 0.00 | 0.000 | ALF3 | 0.00 | 0.000 |
| NA+ | 2079.73 | -1.044 | FE(OH)4-- | 0.00 | 0.000 | ALF4- | 0.00 | 0.000 |
| K+ | 523.88 | -1.873 | FE(OH)++ | 0.00 | 0.000 | ALF5-- | 0.00 | 0.000 |
| CA++ | 17.38 | -3.363 | FE(OH)2+ | 0.00 | 0.000 | ALF6--- | 0.00 | 0.000 |

IONIC STRENGTH = 0.10647 IONIC BALANCE : CATIONS (MOL.EQ.)0.10487515
 ANIONS (MOL.EQ.)0.10693281
 DIFFERENCE (%) -1.94

CHEMICAL GEOTHERMOMETERS DEGREES C 1000/T DEGREES KELVIN = 1.86

QUARTZ 297.6
 CHALCEDONY 999.9
 NAK 314.9

OXIDATION POTENTIAL (VOLTS) : EH H2S= -0.367 EH CH4= 99.999 EH H2= 99.999 EH NH3= 99.999

LOG SOLUBILITY PRODUCTS OF MINERALS IN DEEP WATER

| | TEOR. | CALC. | | TEOR. | CALC. | | TEOR. | CALC. |
|-------------|---------|--------|--------------|---------|---------|-------------|---------|--------|
| ADULARIA | -14.374 | 99.999 | ALBITE LOW | -13.938 | 99.999 | ANALCIME | -11.554 | 99.999 |
| ANHYDRITE | -8.366 | -8.906 | CALCITE | -13.136 | -14.042 | CHALCEDONY | -1.940 | -1.920 |
| MG-CHLORITE | -86.028 | 99.999 | FLUORITE | -11.034 | 99.999 | GOETHITE | 3.058 | 99.999 |
| LAUMONTITE | -24.679 | 99.999 | MICROCLINE | -15.008 | 99.999 | MAGNETITE | -15.567 | 99.999 |
| CA-MONTMOR. | -72.582 | 99.999 | K-MONTMOR. | -33.888 | 99.999 | MG-MONTMOR. | -74.090 | 99.999 |
| NA-MONTMOR. | -34.169 | 99.999 | MUSCOVITE | -17.837 | 99.999 | PREHNITE | -37.810 | 99.999 |
| PYRRHOTITE | -13.082 | 99.999 | PYRITE | -31.297 | 99.999 | QUARTZ | -2.047 | -1.920 |
| WATRAKITE | -24.827 | 99.999 | MOLLASTONITE | 7.253 | 3.411 | ZOISITE | -38.605 | 99.999 |
| EPIDOTE | -38.184 | 99.999 | MARCASITE | -14.975 | 99.999 | | | |

ORKUSTOFNUN JHD
1982-11-12 OLIVER

GEOCHEMISTRY OF WELL OKOY-6

XXXXXXXXXX810823XXXX OKOY-6

S.NEGROS,PHILIPPINES NASUJI

PROGRAM WATCH1.

| WATER SAMPLE (PPM) | | GAS (VOL.%) | | REFERENCE TEMP. | DEGREES C | 285.0 (MEASURED) |
|--------------------|-----------|----------------------------|-----------|-------------------------|------------|------------------|
| PH/DEG.C | 6.60/25.0 | CO2 | 0.00 | SAMPLING PRESSURE | BARS ABS. | 11.1 |
| SI02 | 725.60 | H2S | 0.00 | DISCHARGE ENTHALPY | KJOUL/KG | 1.263 (MEASURED) |
| NA | 2411.00 | H2 | 0.00 | DISCHARGE | KG/SEC. | 13.2 |
| K | 480.00 | O2 | 0.00 | MEASURED TEMPERATURE | DEGREES C | 0.0 |
| CA | 31.90 | CH4 | 0.00 | RESISTIVITY/TEMP. | OHMM/DEG.C | 0.0/ 0.0 |
| MG | 0.030 | N2 | 0.00 | EH/TEMP. | NV/DEG.C | 0.000/ 0.0 |
| CO2 | 6.60 | LITERS GAS PER KG | | MEASURED DOWNHOLE TEMP. | | |
| SO4 | 32.50 | CONDENSATE/DEG.C | 0.00/ 0.0 | DEGREES C/METERS | | FLUID INFLOW |
| H2S | 0.00 | CONDENSATE (PPM) | | 0.0 | 0.0 | DEPTH (METERS) |
| CL | 4436.00 | PH/DEG.C | 0.00/ 0.0 | 0.0 | 0.0 | 0.0 |
| F | 0.00 | CO2 | 0.00 | 0.0 | 0.0 | 0.0 |
| DISS.SOLIDS | 0.00 | H2S | 0.00 | 0.0 | 0.0 | 0.0 |
| AL | 0.0000 | NA | 0.00 | 0.0 | 0.0 | 0.0 |
| B | 52.0000 | CONDENSATE WITH NaOH (PPM) | | 0.0 | 0.0 | 0.0 |
| FE | 0.0000 | CO2 | 12738.00 | 0.0 | 0.0 | 0.0 |
| NH3 | 0.0000 | H2S | 204.40 | 0.0 | 0.0 | 0.0 |

IONIC STRENGTH = 0.12317 IONIC BALANCE : CATIONS (MOL.EQ.)0.11848436
 ANIONS (MOL.EQ.)0.12564923
 DIFFERENCE (%) -5.87

| DEEP WATER (PPM) | | DEEP STEAM (PPM) | | GAS PRESSURES (BARS ABS.) | |
|------------------|---------|------------------|---------|---------------------------|-----------|
| SI02 | 551.23 | CO2 | 2972.49 | CO2 | 0.603E+01 |
| NA | 1831.46 | H2S | 48.73 | H2S | 0.387E-01 |
| K | 364.60 | H2 | 0.00 | H2 | 0.000E+00 |
| CA | 24.23 | O2 | 0.00 | O2 | 0.000E+00 |
| MG | 0.023 | CH4 | 0.00 | CH4 | 0.000E+00 |
| SO4 | 24.69 | N2 | 0.00 | N2 | 0.000E+00 |
| CL | 3369.41 | NH3 | 0.00 | NH3 | 0.000E+00 |
| F | 0.00 | | | H2O | 0.692E+02 |
| DISS.S. | 0.00 | | | TOTAL | 0.752E+02 |
| AL | 0.0000 | | | | |
| B | 39.4969 | | | H2O (%) | 0.05 |
| FE | 0.0000 | | | BOILING PORTION | 0.24 |

GAS SOLUBILITY MULTIPLYING FACTOR : 1.00

ACTIVITY COEFFICIENTS IN DEEP WATER

| | | | | | | | |
|----------|-------|---------|-------|-----------|-------|-----------|-------|
| H+ | 0.685 | KS04- | 0.622 | FE++ | 0.162 | FECL+ | 0.591 |
| OH- | 0.578 | F- | 0.578 | FE+++ | 0.033 | AL+++ | 0.033 |
| H3SiO4- | 0.591 | CL- | 0.565 | FE0H+ | 0.613 | AL0H++ | 0.150 |
| H2SiO4-- | 0.150 | NA+ | 0.591 | FE(OH)3- | 0.613 | AL(OH)2+ | 0.622 |
| H2BO3- | 0.551 | K+ | 0.565 | FE(OH)4-- | 0.142 | AL(OH)4- | 0.602 |
| HCO3- | 0.591 | CA++ | 0.162 | FE0H++ | 0.142 | ALSO4+ | 0.602 |
| CO3-- | 0.132 | MG++ | 0.201 | FE(OH)2+ | 0.622 | AL(SO4)2- | 0.602 |
| HS- | 0.578 | CAHCO3+ | 0.634 | FE(OH)4- | 0.622 | ALF++ | 0.150 |
| S-- | 0.142 | MGHCO3+ | 0.591 | FES04+ | 0.613 | ALF2+ | 0.622 |
| HSO4- | 0.602 | CAOH+ | 0.634 | FECL++ | 0.142 | ALF4- | 0.602 |
| SO4-- | 0.122 | MGOH+ | 0.644 | FECL2+ | 0.613 | ALF5-- | 0.132 |
| NASO4- | 0.622 | NH4+ | 0.551 | FECL4- | 0.591 | ALF6--- | 0.010 |

CHEMICAL COMPONENTS IN DEEP WATER (PPM AND LOG MOLE)

| | | | | | | | | |
|-----------|---------|---------|-----------|--------|---------|-----------|------|-------|
| H+ (ACT.) | 0.02 | -4.827 | MG++ | 0.02 | -6.045 | FE(OH)3 | 0.00 | 0.000 |
| OH- | 0.02 | -5.951 | NACL | 563.27 | -2.016 | FE(OH)4- | 0.00 | 0.000 |
| H4SiO4 | 981.68 | -2.038 | KCL | 44.60 | -3.223 | FECL+ | 0.00 | 0.000 |
| H3SiO4- | 0.04 | -6.335 | NASO4- | 3.13 | -4.580 | FECL2 | 0.00 | 0.000 |
| H2SiO4-- | 0.00 | -12.653 | KS04- | 4.61 | -4.467 | FECL++ | 0.00 | 0.000 |
| NAH3SiO4 | 0.03 | -6.618 | CASO4 | 1.46 | -4.970 | FECL2+ | 0.00 | 0.000 |
| H3BO3 | 225.89 | -2.437 | MGS04 | 0.00 | -7.550 | FECL3 | 0.00 | 0.000 |
| H2BO3- | 0.02 | -6.543 | CACO3 | 0.00 | -8.267 | FECL4- | 0.00 | 0.000 |
| H2CO3 | 4184.33 | -1.171 | MGCO3 | 0.00 | -12.481 | FES04 | 0.00 | 0.000 |
| HCO3- | 3.97 | -4.187 | CAHCO3+ | 1.99 | -4.705 | FES04+ | 0.00 | 0.000 |
| CO3-- | 0.00 | -10.307 | MGHCO3+ | 0.00 | -8.676 | AL+++ | 0.00 | 0.000 |
| H2S | 48.65 | -2.845 | CAOH+ | 0.00 | -7.217 | AL0H++ | 0.00 | 0.000 |
| HS- | 0.08 | -5.618 | MGOH+ | 0.00 | -8.197 | AL(OH)2+ | 0.00 | 0.000 |
| S-- | 0.00 | -14.946 | NH4OH | 0.00 | 0.000 | AL(OH)3 | 0.00 | 0.000 |
| H2SO4 | 0.00 | -9.453 | NH4+ | 0.00 | 0.000 | AL(OH)4- | 0.00 | 0.000 |
| HSO4- | 12.14 | -3.903 | FE++ | 0.00 | 0.000 | ALSO4+ | 0.00 | 0.000 |
| SO4-- | 5.83 | -4.217 | FE+++ | 0.00 | 0.000 | AL(SO4)2- | 0.00 | 0.000 |
| HF | 0.00 | 0.000 | FE0H+ | 0.00 | 0.000 | ALF++ | 0.00 | 0.000 |
| F- | 0.00 | 0.000 | FE(OH)2 | 0.00 | 0.000 | ALF2+ | 0.00 | 0.000 |
| CL- | 3006.53 | -1.072 | FE(OH)3- | 0.00 | 0.000 | ALF3 | 0.00 | 0.000 |
| NA+ | 1609.26 | -1.155 | FE(OH)4-- | 0.00 | 0.000 | ALF4- | 0.00 | 0.000 |
| K+ | 339.88 | -2.061 | FE(OH)++ | 0.00 | 0.000 | ALF5-- | 0.00 | 0.000 |
| CA++ | 23.01 | -3.241 | FE(OH)2+ | 0.00 | 0.000 | ALF6--- | 0.00 | 0.000 |

IONIC STRENGTH = 0.08307 IONIC BALANCE : CATIONS (MOL.EQ.)0.07986062
 ANIONS (MOL.EQ.)0.08500110
 DIFFERENCE (%) -6.24

CHEMICAL GEOTHERMOMETERS DEGREES C 1000/T DEGREES KELVIN = 1.79

QUARTZ 266.4
 CHALCEDONY 999.9
 NAK 288.0

OXIDATION POTENTIAL (VOLTS) : EH H2S= -0.382 EH CH4= 99.999 EH H2= 99.999 EH NH3= 99.999

LOG SOLUBILITY PRODUCTS OF MINERALS IN DEEP WATER

| | TEOR. | CALC. | | TEOR. | CALC. | | TEOR. | CALC. |
|-------------|---------|--------|--------------|---------|---------|-------------|---------|--------|
| ADULARIA | -14.386 | 99.999 | ALBITE LOW | -13.963 | 99.999 | ANALCIME | -11.637 | 99.999 |
| ANHYDRITE | -8.762 | -9.164 | CALCITE | -13.696 | -15.220 | CHALCEDONY | -1.863 | -2.038 |
| MG-CHLORITE | -87.922 | 99.999 | FLUORITE | -11.156 | 99.999 | GOETHITE | 4.274 | 99.999 |
| LAUMONTITE | -24.921 | 99.999 | MICROCLINE | -14.972 | 99.999 | MAGNETITE | -13.468 | 99.999 |
| CA-MONTHOR. | -72.564 | 99.999 | K-MONTHOR. | -33.766 | 99.999 | MG-MONTHOR. | -74.072 | 99.999 |
| NA-MONTHOR. | -34.051 | 99.999 | MUSCOVITE | -17.845 | 99.999 | PREHNITE | -38.645 | 99.999 |
| PYRRHOTITE | -0.939 | 99.999 | PYRITE | -16.041 | 99.999 | QUARTZ | -1.965 | -2.038 |
| WAIRAKITE | -25.314 | 99.999 | WOLLASTONITE | 6.922 | 3.584 | ZOISITE | -39.690 | 99.999 |
| EPIDOTE | -39.816 | 99.999 | MARCASITE | -0.345 | 99.999 | | | |

ACTIVITY COEFFICIENTS IN DEEP WATER

| | | | | | | | |
|----------|-------|---------|-------|-----------|-------|-----------|-------|
| H+ | 0.630 | KSO4- | 0.559 | FE++ | 0.107 | FECL+ | 0.524 |
| OH- | 0.510 | F- | 0.510 | FE+++ | 0.016 | AL+++ | 0.016 |
| H3SI04- | 0.524 | CL- | 0.495 | FEOH+ | 0.549 | ALOH++ | 0.098 |
| H2SI04-- | 0.098 | NA+ | 0.524 | FE(OH)3- | 0.549 | AL(OH)2+ | 0.559 |
| H2BO3- | 0.480 | K+ | 0.495 | FE(OH)4-- | 0.091 | AL(OH)4- | 0.537 |
| HCO3- | 0.524 | CA++ | 0.107 | FEOH++ | 0.091 | ALSO4+ | 0.537 |
| CO3-- | 0.083 | MG++ | 0.141 | FE(OH)2+ | 0.559 | AL(SO4)2- | 0.537 |
| HS- | 0.510 | CAHCO3+ | 0.572 | FE(OH)4- | 0.559 | ALF++ | 0.098 |
| S-- | 0.091 | MGHCO3+ | 0.524 | FESO4+ | 0.549 | ALF2+ | 0.559 |
| HSO4- | 0.537 | CAOH+ | 0.572 | FECL++ | 0.091 | ALF4- | 0.537 |
| SO4--- | 0.075 | MGOH+ | 0.583 | FECL2+ | 0.549 | ALF5-- | 0.083 |
| NASO4- | 0.559 | NH4+ | 0.480 | FECL4- | 0.524 | ALF6--- | 0.004 |

CHEMICAL COMPONENTS IN DEEP WATER (PPM AND LOG MOLE)

| | | | | | | | | |
|-----------|----------|---------|-----------|---------|---------|-----------|------|-------|
| H+ (ACT.) | 0.01 | -4.915 | MG++ | 0.08 | -5.472 | FE(OH)3 | 0.00 | 0.000 |
| OH- | 0.02 | -5.899 | NACL | 1557.53 | -1.574 | FE(OH)4- | 0.00 | 0.000 |
| H4SI04 | 907.39 | -2.025 | KCL | 119.51 | -2.795 | FECL+ | 0.00 | 0.000 |
| H3SI04- | 0.03 | -6.510 | NASO4- | 1.22 | -4.989 | FECL2 | 0.00 | 0.000 |
| H2SI04-- | 0.00 | -12.934 | KSO4- | 5.14 | -4.420 | FECL++ | 0.00 | 0.000 |
| NAH3SI04 | 0.02 | -6.866 | CASO4 | 1.72 | -4.899 | FECL2+ | 0.00 | 0.000 |
| H3BO3 | 259.28 | -2.377 | MESO4 | 0.00 | -7.554 | FECL3 | 0.00 | 0.000 |
| H2BO3- | 0.02 | -6.442 | CACO3 | 0.00 | -7.355 | FECL4- | 0.00 | 0.000 |
| H2CO3 | 16725.57 | -0.569 | MBCO3 | 0.00 | -11.659 | FESO4 | 0.00 | 0.000 |
| HCO3- | 8.31 | -3.866 | CAHCO3+ | 14.47 | -3.844 | FESO4+ | 0.00 | 0.000 |
| CO3--- | 0.00 | -10.163 | MGHCO3+ | 0.00 | -7.729 | AL+++ | 0.00 | 0.000 |
| H2S | 240.41 | -2.152 | CAOH+ | 0.01 | -6.593 | ALOH++ | 0.00 | 0.000 |
| HS- | 0.19 | -5.247 | MGOH+ | 0.00 | -7.095 | AL(OH)2+ | 0.00 | 0.000 |
| S-- | 0.00 | -14.202 | NH4OH | 0.00 | 0.000 | AL(OH)3 | 0.00 | 0.000 |
| H2SO4 | 0.00 | -9.507 | NH4+ | 0.00 | 0.000 | AL(OH)4- | 0.00 | 0.000 |
| HSO4- | 10.10 | -3.983 | FE++ | 0.00 | 0.000 | ALSO4+ | 0.00 | 0.000 |
| SO4--- | 2.47 | -4.589 | FE+++ | 0.00 | 0.000 | AL(SO4)2- | 0.00 | 0.000 |
| HF | 0.00 | 0.000 | FEOH+ | 0.00 | 0.000 | ALF++ | 0.00 | 0.000 |
| F- | 0.00 | 0.000 | FE(OH)2 | 0.00 | 0.000 | ALF2+ | 0.00 | 0.000 |
| CL- | 2784.88 | -1.105 | FE(OH)3- | 0.00 | 0.000 | ALF3 | 0.00 | 0.000 |
| NA+ | 1455.14 | -1.199 | FE(OH)4-- | 0.00 | 0.000 | ALF4- | 0.00 | 0.000 |
| K+ | 409.50 | -1.980 | FE(OH)++ | 0.00 | 0.000 | ALF5-- | 0.00 | 0.000 |
| CA++ | 70.95 | -2.752 | FE(OH)2+ | 0.00 | 0.000 | ALF6--- | 0.00 | 0.000 |

IONIC STRENGTH = 0.07990 IONIC BALANCE : CATIONS (MOL.EQ.)0.07745818
 ANIONS (MOL.EQ.)0.07875313
 DIFFERENCE (%) -1.66

CHEMICAL GEOTHERMOMETERS DEGREES C 1000/T DEGREES KELVIN = 1.69

QUARTZ 269.5
 CHALCEDONY 999.9
 NAK 331.9

OXIDATION POTENTIAL (VOLTS) : EH H2S= -0.443 EH CH4= 99.999 EH H2= 99.999 EH NH3= 99.999

LOG SOLUBILITY PRODUCTS OF MINERALS IN DEEP WATER

| | TEOR. | CALC. | | TEOR. | CALC. | | TEOR. | CALC. |
|-------------|---------|--------|--------------|---------|---------|-------------|---------|--------|
| ADULARIA | -14.524 | 99.999 | ALBITE LOW | -14.114 | 99.999 | ANALCIME | -11.860 | 99.999 |
| ANHYDRITE | -9.396 | -9.433 | CALCITE | -14.600 | -14.964 | CHALCEDONY | -1.752 | -2.025 |
| MG-CHLORITE | -91.310 | 99.999 | FLUORITE | -11.360 | 99.999 | GOETHITE | 6.271 | 99.999 |
| LAUMONTITE | -25.485 | 99.999 | MICROCLINE | -15.050 | 99.999 | MAGNETITE | -10.043 | 99.999 |
| CA-MONTHOR. | -72.299 | 99.999 | K-MONTHOR. | -33.478 | 99.999 | MG-MONTHOR. | -73.793 | 99.999 |
| NA-MONTHOR. | -33.750 | 99.999 | MUSCOVITE | -17.840 | 99.999 | PREHNITE | -40.187 | 99.999 |
| PYRRHOTITE | 18.677 | 99.999 | PYRITE | 8.400 | 99.999 | QUARTZ | -1.849 | -2.025 |
| WAIKITE | -26.235 | 99.999 | WOLLASTONITE | 6.431 | 4.083 | ZOISITE | -41.640 | 99.999 |
| EPIDOTE | -43.932 | 99.999 | MARCASITE | 23.189 | 99.999 | | | |

ORKUSTOFNUN JHD GEOCHEMISTRY OF WELL OKDY-8
 1982-11-12 OLIVER

XXXXXXXXXX820422XXXX OKDY-8

S.NEGROS,PHILIPPINES NASUJI

PROGRAM WATCH3. TEMPERATURE FIXING STEAM LOSS 100.0 DEGREES C

WATER SAMPLE (PPM)

STEAM SAMPLE

| | | | | | |
|-------------|-----------|----------------------------|-------------------------|------------------|-----------------------------|
| PH/DEG.C | 7.85/25.0 | GAS (VOL.%) | REFERENCE TEMP. | DEGREES C | 244.0 (ARBITRARY) |
| SI02 | 864.00 | CO2 | | | |
| NA | 2771.00 | H2S | SAMPLING PRESSURE | BARS ABS. | |
| K | 636.00 | H2 | DISCHARGE ENTHALPY | MJ/OL/KG | |
| CA | 33.60 | O2 | DISCHARGE | KG/SEC. | 32.6 |
| MG | 0.050 | CH4 | | | |
| CO2 | 6.98 | N2 | MEASURED TEMPERATURE | DEGREES C | 0.0 |
| SO4 | 36.10 | | RESISTIVITY/TEMP. | OHMM/DEG.C | 0.0/ 0.0 |
| H2S | 6.48 | | EH/TEMP. | MV/DEG.C | 0.000/ 0.0 |
| CL | 5262.00 | | | | |
| F | 0.00 | LITERS GAS PER KG | | | |
| DISS.SOLIDS | 0.00 | CONDENSATE/DEG.C | MEASURED DOWNHOLE TEMP. | DEGREES C/METERS | FLUID INFLOW DEPTH (METERS) |
| AL | 0.0000 | | | | |
| B | 64.4000 | | | | |
| FE | 0.0000 | CONDENSATE (PPM) | 0.0 | 0.0 | 0.0 |
| NH3 | 0.0000 | PH/DEG.C | 0.0 | 0.0 | 0.0 |
| | | CO2 | 0.0 | 0.0 | 0.0 |
| | | H2S | 0.0 | 0.0 | 0.0 |
| | | NA | 0.0 | 0.0 | 0.0 |
| | | | 0.0 | 0.0 | 0.0 |
| | | | 0.0 | 0.0 | 0.0 |
| | | | 0.0 | 0.0 | 0.0 |
| | | CONDENSATE WITH NAOH (PPM) | 0.0 | 0.0 | 0.0 |
| | | CO2 | 0.0 | 0.0 | 0.0 |
| | | H2S | 0.0 | 0.0 | 0.0 |

IONIC STRENGTH = 0.14498 IONIC BALANCE : CATIONS (MOL.EQ.)0.13799763
 ANIONS (MOL.EQ.)0.14961219
 DIFFERENCE (%) -8.08

| DEEP WATER (PPM) | | DEEP STEAM (PPM) | | GAS PRESSURES (BARS ABS.) | |
|------------------|---------|------------------|---------|---------------------------|-----------|
| SI02 | 620.16 | CO2 | 1672.52 | CO2 | 0.440E+01 |
| NA | 1988.78 | H2S | 757.91 | H2S | 0.792E+00 |
| K | 456.44 | H2 | 0.00 | H2 | 0.000E+00 |
| CA | 24.12 | O2 | 0.00 | O2 | 0.000E+00 |
| MG | 0.036 | CH4 | 0.00 | CH4 | 0.000E+00 |
| SO4 | 25.91 | N2 | 0.00 | N2 | 0.000E+00 |
| CL | 3776.28 | NH3 | 0.00 | NH3 | 0.000E+00 |
| F | 0.00 | | | H2O | 0.359E+02 |
| DISS.S. | 0.00 | | | TOTAL | 0.411E+02 |
| AL | 0.0000 | | | | |
| B | 46.2164 | | | H2O (%) | 0.00 |
| FE | 0.0000 | | | BOILING PORTION | 0.00 |

GAS SOLUBILITY MULTIPLYING FACTOR : 1.00

ACTIVITY COEFFICIENTS IN DEEP WATER

| | | | | | | | |
|----------|-------|---------|-------|-----------|-------|-----------|-------|
| H+ | 0.712 | KS04- | 0.653 | FE++ | 0.200 | FECL+ | 0.623 |
| OH- | 0.611 | F- | 0.611 | FE+++ | 0.050 | AL+++ | 0.050 |
| H3SI04- | 0.623 | CL- | 0.598 | FE0H+ | 0.645 | AL0H++ | 0.187 |
| H2SI04-- | 0.187 | NA+ | 0.623 | FE(OH)3- | 0.645 | AL(OH)2+ | 0.653 |
| H2B03- | 0.584 | K+ | 0.598 | FE(OH)4-- | 0.177 | AL(OH)4- | 0.634 |
| HCO3- | 0.623 | CA++ | 0.200 | FE0H++ | 0.177 | ALSO4+ | 0.634 |
| CO3-- | 0.166 | MG++ | 0.243 | FE(OH)2+ | 0.653 | AL(SO4)2- | 0.634 |
| HS- | 0.611 | CAHCO3+ | 0.665 | FE(OH)4- | 0.653 | ALF++ | 0.187 |
| S-- | 0.177 | MGHCO3+ | 0.623 | FES04+ | 0.645 | ALF2+ | 0.653 |
| HSD4- | 0.634 | CAOH+ | 0.665 | FECL++ | 0.177 | ALF4- | 0.634 |
| S04-- | 0.154 | MGOH+ | 0.674 | FECL2+ | 0.645 | ALF5-- | 0.166 |
| NAS04- | 0.653 | NH4+ | 0.584 | FECL4- | 0.623 | ALF6--- | 0.018 |

CHEMICAL COMPONENTS IN DEEP WATER (PPM AND LOG MOLE)

| | | | | | | | | |
|-----------|---------|---------|-----------|--------|---------|-----------|------|-------|
| H+ (ACT.) | 0.01 | -5.119 | MG++ | 0.03 | -5.861 | FE(OH)3 | 0.00 | 0.000 |
| OH- | 0.03 | -5.733 | NACL | 256.65 | -2.357 | FE(OH)4- | 0.00 | 0.000 |
| H4SI04 | 991.70 | -1.986 | KCL | 29.03 | -3.410 | FECL+ | 0.00 | 0.000 |
| H3SI04- | 0.19 | -5.699 | NAS04- | 5.52 | -4.334 | FECL2 | 0.00 | 0.000 |
| H2SI04-- | 0.00 | -11.449 | KS04- | 5.75 | -4.371 | FECL++ | 0.00 | 0.000 |
| NAH3SI04 | 0.13 | -5.949 | CAS04 | 1.84 | -4.870 | FECL2+ | 0.00 | 0.000 |
| H3B03 | 264.29 | -2.369 | MGS04 | 0.01 | -7.063 | FECL3 | 0.00 | 0.000 |
| H2B03- | 0.05 | -6.092 | CACO3 | 0.00 | -7.669 | FECL4- | 0.00 | 0.000 |
| H2CO3 | 2342.25 | -1.423 | MGCO3 | 0.00 | -11.456 | FES04 | 0.00 | 0.000 |
| HCO3- | 12.61 | -3.685 | CAHCO3+ | 3.69 | -4.437 | FES04+ | 0.00 | 0.000 |
| CO3-- | 0.00 | -9.123 | MGHCO3+ | 0.00 | -8.146 | AL+++ | 0.00 | 0.000 |
| H2S | 750.65 | -1.657 | CAOH+ | 0.00 | -7.235 | AL0H++ | 0.00 | 0.000 |
| HS- | 7.05 | -3.672 | MGOH+ | 0.00 | -8.365 | AL(OH)2+ | 0.00 | 0.000 |
| S-- | 0.00 | -12.992 | NH4OH | 0.00 | 0.000 | AL(OH)3 | 0.00 | 0.000 |
| H2S04 | 0.00 | -10.492 | NH4+ | 0.00 | 0.000 | AL(OH)4- | 0.00 | 0.000 |
| HSD4- | 3.52 | -4.440 | FE++ | 0.00 | 0.000 | ALSO4+ | 0.00 | 0.000 |
| S04-- | 12.58 | -3.883 | FE+++ | 0.00 | 0.000 | AL(SO4)2- | 0.00 | 0.000 |
| HF | 0.00 | 0.000 | FE0H+ | 0.00 | 0.000 | ALF++ | 0.00 | 0.000 |
| F- | 0.00 | 0.000 | FE(OH)2 | 0.00 | 0.000 | ALF2+ | 0.00 | 0.000 |
| CL- | 3606.79 | -0.992 | FE(OH)3- | 0.00 | 0.000 | ALF3 | 0.00 | 0.000 |
| NA+ | 1886.72 | -1.086 | FE(OH)4-- | 0.00 | 0.000 | ALF4- | 0.00 | 0.000 |
| K+ | 439.55 | -1.949 | FE(OH)++ | 0.00 | 0.000 | ALF5-- | 0.00 | 0.000 |
| CA++ | 22.11 | -3.258 | FE(OH)2+ | 0.00 | 0.000 | ALF6--- | 0.00 | 0.000 |

IONIC STRENGTH = 0.09912 IONIC BALANCE : CATIONS (MOL.EQ.)0.09445145
 ANIONS (MOL.EQ.)0.10242935
 DIFFERENCE (%) -8.10

CHEMICAL GEOTHERMOMETERS DEGREES C 1000/T DEGREES KELVIN = 1.93

QUARTZ 279.4
 CHALCEDONY 999.9
 NAK 303.5

OXIDATION POTENTIAL (VOLTS) ; EH H2S= -0.390 EH CH4= 99.999 EH H2= 99.999 EH NH3= 99.999

LOG SOLUBILITY PRODUCTS OF MINERALS IN DEEP WATER

| | TEOR. | CALC. | | TEOR. | CALC. | | TEOR. | CALC. |
|-------------|---------|--------|--------------|---------|---------|-------------|---------|--------|
| ADULARIA | -14.424 | 99.999 | ALBITE LOW | -13.972 | 99.999 | ANALCINE | -11.525 | 99.999 |
| ANHYDRITE | -7.996 | -8.652 | CALCITE | -12.616 | -13.860 | CHALCEDONY | -2.019 | -1.986 |
| MG-CHLORITE | -84.428 | 99.999 | FLUORITE | -10.925 | 99.999 | GOETHITE | 1.944 | 99.999 |
| LAUMONTITE | -24.541 | 99.999 | MICROCLINE | -15.112 | 99.999 | MAGNETITE | -17.506 | 99.999 |
| CA-MONTMOR. | -72.617 | 99.999 | K-MONTMOR. | -34.020 | 99.999 | MG-MONTMOR. | -74.115 | 99.999 |
| NA-MONTMOR. | -34.295 | 99.999 | MUSCOVITE | -17.850 | 99.999 | PREHNITE | -37.131 | 99.999 |
| PYRRHOTITE | -24.401 | 99.999 | PYRITE | -45.633 | 99.999 | QUARTZ | -2.131 | -1.986 |
| WAIKAKITE | -24.442 | 99.999 | WOLLASTONITE | 7.585 | 4.294 | ZOISITE | -37.695 | 99.999 |
| EPIDOTE | -37.289 | 99.999 | MARCASITE | -28.672 | 99.999 | | | |

ACTIVITY COEFFICIENTS IN DEEP WATER

| | | | | | | | |
|----------|-------|---------|-------|-----------|-------|-----------|-------|
| H+ | 0.681 | KS04- | 0.620 | FE++ | 0.159 | FECL+ | 0.589 |
| OH- | 0.576 | F- | 0.576 | FE+++ | 0.032 | AL+++ | 0.032 |
| H3SI04- | 0.589 | CL- | 0.563 | FE0H+ | 0.611 | AL0H++ | 0.147 |
| H2SI04-- | 0.147 | NA+ | 0.589 | FE(OH)3- | 0.611 | AL(OH)2+ | 0.620 |
| H2B03- | 0.550 | K+ | 0.563 | FE(OH)4-- | 0.140 | AL(OH)4- | 0.600 |
| HCO3- | 0.589 | CA++ | 0.159 | FE0H++ | 0.140 | ALSO4+ | 0.600 |
| CO3-- | 0.130 | MG++ | 0.197 | FE(OH)2+ | 0.620 | AL(SO4)2- | 0.600 |
| HS- | 0.576 | CAHCO3+ | 0.631 | FE(OH)4- | 0.620 | ALF++ | 0.147 |
| S-- | 0.140 | MGHCO3+ | 0.589 | FESO4+ | 0.611 | ALF2+ | 0.620 |
| HSO4- | 0.600 | CAOH+ | 0.631 | FECL++ | 0.140 | ALF4- | 0.600 |
| SO4-- | 0.120 | MGOH+ | 0.641 | FECL2+ | 0.611 | ALF5-- | 0.130 |
| NASO4- | 0.620 | NH4+ | 0.550 | FECL4- | 0.589 | ALF6--- | 0.010 |

CHEMICAL COMPONENTS IN DEEP WATER (PPM AND LOG MOLE)

| | | | | | | | | |
|-----------|---------|---------|-----------|--------|---------|-----------|------|-------|
| H+ (ACT.) | 0.00 | -5.374 | MG++ | 0.12 | -5.310 | FE(OH)3 | 0.00 | 0.000 |
| OH- | 0.07 | -5.408 | NACL | 577.80 | -2.005 | FE(OH)4- | 0.00 | 0.000 |
| H4SI04 | 801.37 | -2.079 | KCL | 46.81 | -3.202 | FECL+ | 0.00 | 0.000 |
| H3SI04- | 0.12 | -5.882 | NASO4- | 4.07 | -4.466 | FECL2 | 0.00 | 0.000 |
| H2SI04-- | 0.00 | -11.704 | KS04- | 7.24 | -4.271 | FECL++ | 0.00 | 0.000 |
| NAH3SI04 | 0.08 | -6.177 | CASO4 | 3.62 | -4.575 | FECL2+ | 0.00 | 0.000 |
| H3B03 | 202.96 | -2.484 | MGS04 | 0.02 | -6.697 | FECL3 | 0.00 | 0.000 |
| H2B03- | 0.05 | -6.060 | CAC03 | 0.02 | -6.625 | FECL4- | 0.00 | 0.000 |
| H2CO3 | 9133.78 | -0.832 | MGC03 | 0.00 | -10.385 | FESO4 | 0.00 | 0.000 |
| HCO3- | 25.81 | -3.374 | CAHCO3+ | 24.90 | -3.608 | FESO4+ | 0.00 | 0.000 |
| CO3-- | 0.00 | -9.014 | MGHCO3+ | 0.01 | -7.098 | AL+++ | 0.00 | 0.000 |
| H2S | 153.50 | -2.346 | CAOH+ | 0.02 | -6.386 | AL0H++ | 0.00 | 0.000 |
| HS- | 0.74 | -4.650 | MGOH+ | 0.01 | -6.816 | AL(OH)2+ | 0.00 | 0.000 |
| S-- | 0.00 | -13.397 | NH4OH | 0.00 | 0.000 | AL(OH)3 | 0.00 | 0.000 |
| H2SO4 | 0.00 | -10.324 | NH4+ | 0.00 | 0.000 | AL(OH)4- | 0.00 | 0.000 |
| HSO4- | 5.38 | -4.256 | FE++ | 0.00 | 0.000 | ALSO4+ | 0.00 | 0.000 |
| SO4-- | 7.36 | -4.116 | FE+++ | 0.00 | 0.000 | AL(SO4)2- | 0.00 | 0.000 |
| HF | 0.00 | 0.000 | FE0H+ | 0.00 | 0.000 | ALF++ | 0.00 | 0.000 |
| F- | 0.00 | 0.000 | FE(OH)2 | 0.00 | 0.000 | ALF2+ | 0.00 | 0.000 |
| CL- | 2567.81 | -1.140 | FE(OH)3- | 0.00 | 0.000 | ALF3 | 0.00 | 0.000 |
| NA+ | 1533.11 | -1.176 | FE(OH)4-- | 0.00 | 0.000 | ALF4- | 0.00 | 0.000 |
| K+ | 352.20 | -2.045 | FE(OH)++ | 0.00 | 0.000 | ALF5-- | 0.00 | 0.000 |
| CA++ | 40.57 | -2.995 | FE(OH)2+ | 0.00 | 0.000 | ALF6--- | 0.00 | 0.000 |

IONIC STRENGTH = 0.07660 IONIC BALANCE : CATIONS (MOL.EQ.)0.07797492
 ANIONS (MOL.EQ.)0.07303931
 DIFFERENCE (%) 6.54

CHEMICAL GEOTHERMOMETERS DEGREES C 1000/T DEGREES KELVIN = 1.77

QUARTZ 256.2
 CHALCEDONY 999.9
 NAK 300.8

OXIDATION POTENTIAL (VOLTS) : EH H2S= -0.470 EH CH4= 99.999 EH H2= 99.999 EH NH3= 99.999

LOG SOLUBILITY PRODUCTS OF MINERALS IN DEEP WATER

| | TEOR. | CALC. | | TEOR. | CALC. | | TEOR. | CALC. |
|-------------|---------|--------|--------------|---------|---------|-------------|---------|--------|
| ADULARIA | -14.401 | 99.999 | ALBITE LOW | -13.981 | 99.999 | AMALCINE | -11.669 | 99.999 |
| ANHYDRITE | -8.876 | -8.830 | CALCITE | -13.858 | -13.694 | CHALCEDONY | -1.842 | -2.079 |
| MG-CHLORITE | -88.501 | 99.999 | FLUORITE | -11.192 | 99.999 | GOETHITE | 4.629 | 99.999 |
| LAUMONTITE | -25.008 | 99.999 | MICROCLINE | -14.974 | 99.999 | MAGNETITE | -12.857 | 99.999 |
| CA-MONTMOR. | -72.549 | 99.999 | K-MONTMOR. | -53.728 | 99.999 | MG-MONTMOR. | -74.053 | 99.999 |
| NA-MONTMOR. | -34.010 | 99.999 | MUSCOVITE | -17.848 | 99.999 | PREHNITE | -38.906 | 99.999 |
| PYRRHOTITE | 2.578 | 99.999 | PYRITE | -11.641 | 99.999 | QUARTZ | -1.943 | -2.079 |
| WAIRAKITE | -25.468 | 99.999 | WOLLASTONITE | 6.831 | 4.876 | ZOISITE | -40.023 | 99.999 |
| EPIDOTE | -40.420 | 99.999 | MARCASITE | 3.883 | 99.999 | | | |

ORKUSTOFNUN JHD
1982-11-12 OLIVER

GEOCHEMISTRY OF WELL OKDY-10D

XXXXXXXXXX810530XXXX OKDY-10D

S.NEGROS,PHILIPPINES PUHAGAN

PROGRAM WATCH1.

WATER SAMPLE (PPM)

STEAM SAMPLE

| | | | | | |
|----------|-----------|-------------|-----------------|----------------------|---------------------------|
| PH/DEG.C | 6.69/25.0 | GAS (VOL.%) | REFERENCE TEMP. | DEGREES C | 259.0 (MEASURED) |
| SI02 | 717.00 | CO2 | 0.00 | | |
| NA | 3167.00 | H2S | 0.00 | SAMPLING PRESSURE | BARS ABS. 2.6 |
| K | 634.00 | H2 | 0.00 | DISCHARGE ENTHALPY | MJ/OL/KG 1.313 (MEASURED) |
| CA | 122.00 | O2 | 0.00 | DISCHARGE | KG/SEC. 50.3 |
| MG | 0.850 | CH4 | 0.00 | | |
| CO2 | 2.60 | N2 | 0.00 | MEASURED TEMPERATURE | DEGREES C 0.0 |
| SO4 | 51.50 | | | RESISTIVITY/TEMP. | OHMM/DEG.C 0.0/ 0.0 |
| H2S | 0.00 | | | EH/TEMP. | MV/DEG.C 0.000/ 0.0 |
| CL | 5424.00 | | | | |

F 0.00
DISS.SOLIDS 0.00
AL 0.0000
B 76.1000
FE 0.0000
NH3 0.0000

LITERS GAS PER KG
CONDENSATE/DEG.C 3.76/25.0

MEASURED DOWNHOLE TEMP.
DEGREES C/METERS

FLUID INFLOW
DEPTH (METERS)

| | | | |
|----------------------------|-----------|-----|-----|
| CONDENSATE (PPM) | 0.0 | 0.0 | 0.0 |
| PH/DEG.C | 0.00/ 0.0 | 0.0 | 0.0 |
| CO2 | 0.00 | 0.0 | 0.0 |
| H2S | 0.00 | 0.0 | 0.0 |
| NA | 0.00 | 0.0 | 0.0 |
| | | 0.0 | 0.0 |
| | | 0.0 | 0.0 |
| | | 0.0 | 0.0 |
| | | 0.0 | 0.0 |
| CONDENSATE WITH NAOH (PPM) | 0.0 | 0.0 | 0.0 |
| CO2 | 7079.00 | 0.0 | 0.0 |
| H2S | 165.80 | 0.0 | 0.0 |

IONIC STRENGTH = 0.16024 IONIC BALANCE : CATIONS (MOL.EQ.)0.15971066
 ANIONS (MOL.EQ.)0.15369166
 DIFFERENCE (%) 3.84

DEEP WATER (PPM)

DEEP STEAM (PPM)

GAS PRESSURES (BARS ABS.)

| | | | | | | | |
|---------|---------|-----|--------|-----------------|----------|-------|-----------|
| SI02 | 519.39 | CO2 | 168.08 | CO2 | 21418.91 | CO2 | 0.405E+00 |
| NA | 2293.94 | H2S | 11.32 | H2S | 442.03 | H2S | 0.108E-01 |
| K | 459.20 | H2 | 0.00 | H2 | 0.00 | H2 | 0.000E+00 |
| CA | 88.37 | O2 | 0.00 | O2 | 0.00 | O2 | 0.000E+00 |
| MG | 0.616 | CH4 | 0.00 | CH4 | 0.00 | CH4 | 0.000E+00 |
| SO4 | 37.30 | N2 | 0.00 | N2 | 0.00 | N2 | 0.000E+00 |
| CL | 3928.42 | NH3 | 0.00 | NH3 | 0.00 | NH3 | 0.000E+00 |
| F | 0.00 | | | | | H2O | 0.462E+02 |
| DISS.S. | 0.00 | | | | | TOTAL | 0.466E+02 |
| AL | 0.0000 | | | | | | |
| B | 55.1162 | | | H2O (%) | 11.07 | | |
| FE | 0.0000 | | | BOILING PORTION | 0.28 | | |

GAS SOLUBILITY MULTIPLYING FACTOR : 1.00

ACTIVITY COEFFICIENTS IN DEEP WATER

| | | | | | | | |
|----------|-------|---------|-------|-----------|-------|-----------|-------|
| H+ | 0.694 | KS04- | 0.630 | FE++ | 0.174 | FECL+ | 0.597 |
| OH- | 0.584 | F- | 0.584 | FE+++ | 0.039 | AL+++ | 0.039 |
| H3SI04- | 0.597 | CL- | 0.570 | FE0H+ | 0.621 | AL0H++ | 0.161 |
| H2SI04-- | 0.161 | NA+ | 0.597 | FE(OH)3- | 0.621 | AL(OH)2+ | 0.630 |
| H2B03- | 0.555 | K+ | 0.570 | FE(OH)4-- | 0.152 | AL(OH)4- | 0.609 |
| HCO3- | 0.597 | CA++ | 0.174 | FE0H++ | 0.152 | ALSO4+ | 0.609 |
| CO3-- | 0.141 | MG++ | 0.216 | FE(OH)2+ | 0.630 | AL(SO4)2- | 0.609 |
| HS- | 0.584 | CAHCO3+ | 0.642 | FE(OH)4- | 0.630 | ALF++ | 0.161 |
| S-- | 0.152 | MGHCO3+ | 0.597 | FES04+ | 0.621 | ALF2+ | 0.630 |
| HSO4- | 0.609 | CAOH+ | 0.642 | FECL+ | 0.152 | ALF4- | 0.609 |
| SO4-- | 0.130 | MGOH+ | 0.652 | FECL2+ | 0.621 | ALF5-- | 0.141 |
| NAS04- | 0.630 | NH4+ | 0.555 | FECL4- | 0.597 | ALF6--- | 0.012 |

CHEMICAL COMPONENTS IN DEEP WATER (PPM AND LOG MOLE)

| | | | | | | | | |
|-----------|---------|---------|-----------|--------|---------|-----------|------|-------|
| H+ (ACT.) | 0.01 | -5.271 | MG++ | 0.57 | -4.628 | FE(OH)3 | 0.00 | 0.000 |
| OH- | 0.05 | -5.528 | NACL | 393.55 | -2.172 | FE(OH)4- | 0.00 | 0.000 |
| H4SI04 | 830.50 | -2.063 | KCL | 36.98 | -3.305 | FECL+ | 0.00 | 0.000 |
| H3SI04- | 0.19 | -5.709 | NAS04- | 7.67 | -4.191 | FECL2 | 0.00 | 0.000 |
| H2SI04-- | 0.00 | -11.380 | KS04- | 7.90 | -4.233 | FECL++ | 0.00 | 0.000 |
| NAH3SI04 | 0.14 | -5.913 | CAS04 | 8.61 | -4.199 | FECL2+ | 0.00 | 0.000 |
| H3B03 | 315.17 | -2.293 | MGS04 | 0.19 | -5.801 | FECL3 | 0.00 | 0.000 |
| H2B03- | 0.08 | -5.882 | CAC03 | 0.00 | -7.905 | FECL4- | 0.00 | 0.000 |
| H2CO3 | 234.63 | -2.422 | MGC03 | 0.00 | -11.119 | FES04 | 0.00 | 0.000 |
| HCO3- | 1.26 | -4.686 | CAHCO3+ | 1.60 | -4.800 | FES04+ | 0.00 | 0.000 |
| CO3-- | 0.00 | -10.084 | MGHCO3+ | 0.00 | -7.879 | AL+++ | 0.00 | 0.000 |
| H2S | 11.21 | -3.483 | CAOH+ | 0.02 | -6.397 | AL0H++ | 0.00 | 0.000 |
| HS- | 0.11 | -5.494 | MGOH+ | 0.01 | -6.760 | AL(OH)2+ | 0.00 | 0.000 |
| S-- | 0.00 | -14.534 | NH4OH | 0.00 | 0.000 | AL(OH)3 | 0.00 | 0.000 |
| H2SO4 | 0.00 | -10.469 | NH4+ | 0.00 | 0.000 | AL(OH)4- | 0.00 | 0.000 |
| HSO4- | 4.46 | -4.337 | FE++ | 0.00 | 0.000 | ALSO4+ | 0.00 | 0.000 |
| SO4-- | 14.85 | -3.811 | FE+++ | 0.00 | 0.000 | AL(SO4)2- | 0.00 | 0.000 |
| HF | 0.00 | 0.000 | FE0H+ | 0.00 | 0.000 | ALF++ | 0.00 | 0.000 |
| F- | 0.00 | 0.000 | FE(OH)2 | 0.00 | 0.000 | ALF2+ | 0.00 | 0.000 |
| CL- | 3672.10 | -0.985 | FE(OH)3- | 0.00 | 0.000 | ALF3 | 0.00 | 0.000 |
| NA+ | 2137.61 | -1.032 | FE(OH)4-- | 0.00 | 0.000 | ALF4- | 0.00 | 0.000 |
| K+ | 437.52 | -1.951 | FE(OH)++ | 0.00 | 0.000 | ALF5-- | 0.00 | 0.000 |
| CA++ | 85.18 | -2.673 | FE(OH)2+ | 0.00 | 0.000 | ALF6--- | 0.00 | 0.000 |

IONIC STRENGTH = 0.10851 IONIC BALANCE : CATIONS (MOL.EQ.)0.10848394
 ANIONS (MOL.EQ.)0.10392465
 DIFFERENCE (%) 4.29

CHEMICAL GEOTHERMOMETERS DEGREES C 1000/T DEGREES KELVIN = 1.88

QUARTZ 259.9
 CHALCEDONY 999.9
 MAK 283.1

OXIDATION POTENTIAL (VOLTS) : EH H2S= -0.403 EH CH4= 99.999 EH H2= 99.999 EH NH3= 99.999

LOG SOLUBILITY PRODUCTS OF MINERALS IN DEEP WATER

| | TEOR. | CALC. | | TEOR. | CALC. | | TEOR. | CALC. |
|-------------|---------|--------|--------------|---------|---------|-------------|---------|--------|
| ADULARIA | -14.381 | 99.999 | ALBITE LOW | -13.941 | 99.999 | ANALCIME | -11.542 | 99.999 |
| ANHYDRITE | -8.273 | -8.129 | CALCITE | -13.005 | -14.367 | CHALCEDONY | -1.959 | -2.063 |
| MG-CHLORITE | -85.609 | 99.999 | FLUORITE | -11.006 | 99.999 | GOETHITE | 2.776 | 99.999 |
| LAUMONTITE | -24.635 | 99.999 | MICROCLINE | -15.028 | 99.999 | MAGNETITE | -16.058 | 99.999 |
| CA-MONTMOR. | -72.386 | 99.999 | K-MONTMOR. | -33.916 | 99.999 | MG-MONTMOR. | -74.090 | 99.999 |
| NA-MONTMOR. | -34.197 | 99.999 | MUSCOVITE | -17.838 | 99.999 | PREHNITE | -37.629 | 99.999 |
| PYRRHOTITE | -15.934 | 99.999 | PYRITE | -34.898 | 99.999 | QUARTZ | -2.067 | -2.063 |
| WAIKAKITE | -24.723 | 99.999 | WOLLASTONITE | 7.335 | 5.046 | ZOISITE | -38.365 | 99.999 |
| EPIDOTE | -37.900 | 99.999 | MARCASITE | -18.420 | 99.999 | | | |

ORKUSTOFNUN JHD
1982-11-12 OLIVER

GEOCHEMISTRY OF WELL OKOY-12D

XXXXXXXXXX810912XXXX OKOY-12D

S.NEGROS, PHILIPPINES PUHAGAN

PROGRAM WATCH1.

WATER SAMPLE (PPM)

STEAM SAMPLE

| | | | | | |
|----------|-----------|-------------|-----------------|----------------------|---------------------------|
| PH/DEG.C | 6.70/25.0 | GAS (VOL.%) | REFERENCE TEMP. | DEGREES C | 250.0 (MEASURED) |
| SI02 | 532.00 | CO2 | 0.00 | | |
| NA | 2352.00 | H2S | 0.00 | SAMPLING PRESSURE | BARS ABS. 11.0 |
| K | 457.00 | H2 | 0.00 | DISCHARGE ENTHALPY | MJOUL/KG 1.089 (MEASURED) |
| CA | 76.00 | O2 | 0.00 | DISCHARGE | KG/SEC. 16.8 |
| MG | 0.360 | CH4 | 0.00 | | |
| CO2 | 43.00 | N2 | 0.00 | MEASURED TEMPERATURE | DEGREES C 0.0 |
| SO4 | 32.00 | | | RESISTIVITY/TEMP. | OHMM/DEG.C 0.0/ 0.0 |
| H2S | 82.60 | | | EH/TEMP. | NV/DEG.C 0.000/ 0.0 |
| CL | 4361.00 | | | | |

F 0.00
DISS.SOLIDS 0.00
AL 0.0000
B 49.2000
FE 0.0000
NH3 0.0000

LITERS GAS PER KG
CONDENSATE/DEG.C 16.90/25.0

MEASURED DOWNHOLE TEMP.
DEGREES C/METERS

FLUID INFLOW
DEPTH (METERS)

| | | | |
|----------------------------|-----------|-----|-----|
| CONDENSATE (PPM) | 0.0 | 0.0 | 0.0 |
| PH/DEG.C | 0.00/ 0.0 | 0.0 | 0.0 |
| CO2 | 0.00 | 0.0 | 0.0 |
| H2S | 0.00 | 0.0 | 0.0 |
| NA | 0.00 | 0.0 | 0.0 |
| | | 0.0 | 0.0 |
| | | 0.0 | 0.0 |
| | | 0.0 | 0.0 |
| | | 0.0 | 0.0 |
| CONDENSATE WITH NaOH (PPM) | 0.0 | 0.0 | 0.0 |
| CO2 | 31728.90 | 0.0 | 0.0 |
| H2S | 435.30 | 0.0 | 0.0 |

IONIC STRENGTH = 0.12351

IONIC BALANCE : CATIONS (MOL.EQ.)0.11755587
ANIONS (MOL.EQ.)0.12505311
DIFFERENCE (%) -6.18

DEEP WATER (PPM)

DEEP STEAM (PPM)

GAS PRESSURES (BARS ABS.)

| | | | | | | | |
|---------|---------|-----|---------|-----------------|-----------|-------|-----------|
| SI02 | 451.20 | CO2 | 3441.02 | CO2 | 532191.44 | CO2 | 0.866E+01 |
| NA | 1994.63 | H2S | 121.17 | H2S | 5742.93 | H2S | 0.121E+00 |
| K | 387.54 | H2 | 0.00 | H2 | 0.00 | H2 | 0.000E+00 |
| CA | 64.45 | O2 | 0.00 | O2 | 0.00 | O2 | 0.000E+00 |
| MG | 0.305 | CH4 | 0.00 | CH4 | 0.00 | CH4 | 0.000E+00 |
| SO4 | 27.14 | N2 | 0.00 | N2 | 0.00 | N2 | 0.000E+00 |
| CL | 3698.06 | NH3 | 0.00 | NH3 | 0.00 | NH3 | 0.000E+00 |
| F | 0.00 | | | | | H2O | 0.398E+02 |
| DISS.S. | 0.00 | | | | | TOTAL | 0.486E+02 |
| AL | 0.0000 | | | | | | |
| B | 41.7204 | | | H2O (%) | | | 0.28 |
| FE | 0.0000 | | | BOILING PORTION | | | 0.15 |

GAS SOLUBILITY MULTIPLYING FACTOR : 1.00

ACTIVITY COEFFICIENTS IN DEEP WATER

| | | | | | | | |
|----------|-------|---------|-------|-----------|-------|-----------|-------|
| H+ | 0.707 | KSO4- | 0.648 | FE++ | 0.193 | FECL+ | 0.617 |
| OH- | 0.605 | F- | 0.605 | FE+++ | 0.047 | AL+++ | 0.047 |
| H3SiO4- | 0.617 | CL- | 0.592 | FE(OH) | 0.639 | ALOH++ | 0.180 |
| H2SiO4-- | 0.180 | NA+ | 0.617 | FE(OH)3- | 0.639 | AL(OH)2+ | 0.648 |
| H2BO3- | 0.578 | K+ | 0.592 | FE(OH)4-- | 0.171 | AL(OH)4- | 0.629 |
| HCO3- | 0.617 | CA++ | 0.193 | FE(OH)++ | 0.171 | ALSO4+ | 0.629 |
| CO3-- | 0.160 | MG++ | 0.236 | FE(OH)2+ | 0.648 | AL(SO4)2- | 0.629 |
| HS- | 0.605 | CAHCO3+ | 0.659 | FE(OH)4- | 0.648 | ALF++ | 0.180 |
| S-- | 0.171 | MGHCO3+ | 0.617 | FES04+ | 0.639 | ALF2+ | 0.648 |
| HSO4- | 0.629 | CAOH+ | 0.659 | FECL++ | 0.171 | ALF4- | 0.629 |
| SO4-- | 0.149 | MGOH+ | 0.668 | FECL2+ | 0.639 | ALF5-- | 0.160 |
| NASO4- | 0.648 | NH4+ | 0.578 | FECL4- | 0.617 | ALF6--- | 0.016 |

CHEMICAL COMPONENTS IN DEEP WATER (PPM AND LOG MOLE)

| | | | | | | | | |
|-----------|---------|---------|-----------|--------|--------|-----------|------|-------|
| H+ (ACT.) | 0.00 | -5.523 | MG++ | 0.28 | -4.943 | FE(OH)3 | 0.00 | 0.000 |
| OH- | 0.08 | -5.310 | NACL | 281.43 | -2.317 | FE(OH)4- | 0.00 | 0.000 |
| H4SiO4 | 721.23 | -2.125 | KCL | 26.61 | -3.447 | FECL+ | 0.00 | 0.000 |
| H3SiO4- | 0.32 | -5.469 | NASO4- | 6.08 | -4.292 | FECL2 | 0.00 | 0.000 |
| H2SiO4-- | 0.00 | -10.850 | KSO4- | 5.62 | -4.381 | FECL++ | 0.00 | 0.000 |
| NAH3SiO4 | 0.23 | -5.717 | CASO4 | 4.47 | -4.483 | FECL2+ | 0.00 | 0.000 |
| H3BO3 | 238.52 | -2.414 | MGSO4 | 0.09 | -6.118 | FECL3 | 0.00 | 0.000 |
| H2BO3- | 0.11 | -5.744 | CAC03 | 0.06 | -6.256 | FECL4- | 0.00 | 0.000 |
| H2CO3 | 4769.28 | -1.114 | MGC03 | 0.00 | -9.493 | FES04 | 0.00 | 0.000 |
| HCO3- | 56.15 | -3.036 | CAHCO3+ | 38.37 | -3.421 | FES04+ | 0.00 | 0.000 |
| CO3-- | 0.00 | -8.122 | MGHCO3+ | 0.02 | -6.558 | AL+++ | 0.00 | 0.000 |
| H2S | 118.64 | -2.458 | CAOH+ | 0.02 | -6.446 | ALOH++ | 0.00 | 0.000 |
| HS- | 2.45 | -4.130 | MGOH+ | 0.00 | -6.950 | AL(OH)2+ | 0.00 | 0.000 |
| S-- | 0.00 | -13.002 | NH4OH | 0.00 | 0.000 | AL(OH)3 | 0.00 | 0.000 |
| H2SO4 | 0.00 | -11.162 | NH4+ | 0.00 | 0.000 | AL(OH)4- | 0.00 | 0.000 |
| HSO4- | 1.77 | -4.738 | FE++ | 0.00 | 0.000 | ALSO4+ | 0.00 | 0.000 |
| SO4-- | 13.26 | -3.860 | FE+++ | 0.00 | 0.000 | AL(SO4)2- | 0.00 | 0.000 |
| HF | 0.00 | 0.000 | FE(OH) | 0.00 | 0.000 | ALF++ | 0.00 | 0.000 |
| F- | 0.00 | 0.000 | FE(OH)2 | 0.00 | 0.000 | ALF2+ | 0.00 | 0.000 |
| CL- | 3514.69 | -1.004 | FE(OH)3- | 0.00 | 0.000 | ALF3 | 0.00 | 0.000 |
| NA+ | 1882.69 | -1.087 | FE(OH)4-- | 0.00 | 0.000 | ALF4- | 0.00 | 0.000 |
| K+ | 371.96 | -2.022 | FE(OH)++ | 0.00 | 0.000 | ALF5-- | 0.00 | 0.000 |
| CA++ | 47.89 | -2.923 | FE(OH)2+ | 0.00 | 0.000 | ALF6--- | 0.00 | 0.000 |

IONIC STRENGTH = 0.09866 IONIC BALANCE : CATIONS (MOL.EQ.)0.09419770
 ANIONS (MOL.EQ.)0.10042552
 DIFFERENCE (%) -6.40

CHEMICAL GEOTHERMOMETERS DEGREES C 1000/T DEGREES KELVIN = 1.91

| | |
|------------|-------|
| QUARTZ | 243.3 |
| CHALCEDONY | 999.9 |
| NAK | 278.7 |

OXIDATION POTENTIAL (VOLTS) : EH H2S= -0.439 EH CH4= 99.999 EH H2= 99.999 EH NH3= 99.999

LOG SOLUBILITY PRODUCTS OF MINERALS IN DEEP WATER

| | TEOR. | CALC. | | TEOR. | CALC. | | TEOR. | CALC. |
|-------------|---------|--------|--------------|---------|---------|-------------|---------|--------|
| ADULARIA | -14.402 | 99.999 | ALBITE LOW | -13.956 | 99.999 | ANALCIME | -11.528 | 99.999 |
| AMHYDRITE | -8.106 | -8.325 | CALCITE | -12.771 | -12.555 | CHALCEDONY | -1.995 | -2.125 |
| MG-CHLORITE | -84.886 | 99.999 | FLUORITE | -10.957 | 99.999 | GOETHITE | 2.273 | 99.999 |
| LAUMONTITE | -24.572 | 99.999 | MICROCLINE | -15.073 | 99.999 | MAGNETITE | -16.930 | 99.999 |
| CA-MONTMOR. | -72.600 | 99.999 | K-MONTMOR. | -33.975 | 99.999 | MG-MONTMOR. | -74.098 | 99.999 |
| NA-MONTMOR. | -34.252 | 99.999 | MUSCOVITE | -17.843 | 99.999 | PREHNITE | -37.322 | 99.999 |
| PYRRHOTITE | -21.030 | 99.999 | PYRITE | -11.351 | 99.999 | QUARTZ | -2.105 | -2.125 |
| WAIKAKITE | -24.549 | 99.999 | WOLLASTONITE | 7.484 | 5.285 | ZOISITE | -37.954 | 99.999 |
| EPIDOTE | -37.492 | 99.999 | MARCASITE | -24.586 | 99.999 | | | |

ORKUSTOFNUN JHD
1982-11-12 OLIVER

GEOCHEMISTRY OF WELL NJ-1D

XXXXXXXXXX020427XXXX NASUJI-1D

S.NEGROS,PHILIPPINES NASUJI

PROGRAM WATCH3, TEMPERATURE FIXING STEAM LOSS 100.0 DEGREES C

WATER SAMPLE (PPM)

STEAM SAMPLE

| PH/DEG.C | 5.20/25.0 | GAS (VOL.%) | REFERENCE TEMP. | DEGREES C | 243.0 (ARBITRARY) |
|-------------|-----------|----------------------------|-------------------------|------------------|--------------------------------|
| SI02 | 600.00 | CO2 | | | |
| NA | 1930.00 | H2S | SAMPLING PRESSURE | BARS ABS. | |
| K | 393.00 | H2 | DISCHARGE ENTHALPY | MJ/OL/KG | |
| CA | 16.80 | O2 | DISCHARGE | KG/SEC. | 0.0 |
| MG | 5.250 | CH4 | | | |
| CO2 | 2.54 | N2 | MEASURED TEMPERATURE | DEGREES C | 0.0 |
| SO4 | 155.00 | | RESISTIVITY/TEMP. | OHMM/DEG.C | 0.0/ 0.0 |
| H2S | 2.39 | | EH/TEMP. | MV/DEG.C | 0.000/ 0.0 |
| CL | 3489.00 | | | | |
| F | 0.00 | LITERS GAS PER KG | | | |
| DISS.SOLIDS | 0.00 | CONDENSATE/DEG.C | MEASURED DOWNHOLE TEMP. | DEGREES C/METERS | FLUID INFLOW DEPTH (METERS) |
| AL | 0.0000 | | | | |
| B | 46.9000 | CONDENSATE (PPM) | 0.0 | 0.0 | 0.0 |
| FE | 0.0000 | PH/DEG.C | 0.0 | 0.0 | 0.0 |
| NH3 | 0.0000 | CO2 | 0.0 | 0.0 | 0.0 |
| | | H2S | 0.0 | 0.0 | 0.0 |
| | | NA | 0.0 | 0.0 | 0.0 |
| | | | 0.0 | 0.0 | 0.0 |
| | | | 0.0 | 0.0 | 0.0 |
| | | | 0.0 | 0.0 | 0.0 |
| | | CONDENSATE WITH NAOH (PPM) | 0.0 | 0.0 | 0.0 |
| | | CO2 | 0.0 | 0.0 | 0.0 |
| | | H2S | 0.0 | 0.0 | 0.0 |

IONIC STRENGTH = 0.10023 IONIC BALANCE : CATIONS (MOL.EQ.)0.09498397
ANIONS (MOL.EQ.)0.10125270
DIFFERENCE (%) -6.39

PH LOWER THAN 4.00

| DEEP WATER (PPM) | DEEP STEAM (PPM) | GAS PRESSURES (BARS ABS.) |
|------------------|----------------------|---------------------------|
| SI02 431.94 | CO2 3727.97 | CO2 0.987E+01 |
| NA 1389.30 | H2S 1220.27 | H2S 0.129E+01 |
| K 282.88 | H2 0.00 | H2 0.000E+00 |
| CA 12.09 | O2 0.00 | O2 0.000E+00 |
| MG 3.779 | CH4 0.00 | CH4 0.000E+00 |
| SO4 111.57 | N2 0.00 | N2 0.000E+00 |
| CL 2511.32 | NH3 0.00 | NH3 0.000E+00 |
| F 0.00 | | H2O 0.353E+02 |
| DISS.S. 0.00 | | TOTAL 0.464E+02 |
| AL 0.0000 | | |
| B 33.7576 | H2O (%) 0.00 | |
| FE 0.0000 | BOILING PORTION 0.00 | |

GAS SOLUBILITY MULTIPLYING FACTOR : 1.00

ACTIVITY COEFFICIENTS IN DEEP WATER

| | | | | | | | |
|----------|-------|---------|-------|-----------|-------|-----------|-------|
| H+ | 0.736 | KSO4- | 0.686 | FE++ | 0.239 | FECL+ | 0.662 |
| OH- | 0.652 | F- | 0.652 | FE+++ | 0.066 | AL+++ | 0.066 |
| H3SiO4- | 0.662 | CL- | 0.641 | FEOH+ | 0.680 | ALOH++ | 0.226 |
| H2SiO4-- | 0.226 | NA+ | 0.662 | FE(OH)3- | 0.680 | AL(OH)2+ | 0.686 |
| H2BO3- | 0.630 | K+ | 0.641 | FE(OH)4-- | 0.217 | AL(OH)4- | 0.671 |
| HCO3- | 0.662 | CA++ | 0.239 | FEOH++ | 0.217 | ALSO4+ | 0.671 |
| CO3-- | 0.206 | MG++ | 0.279 | FE(OH)2+ | 0.686 | AL(SO4)2- | 0.671 |
| HS- | 0.652 | CAHCO3+ | 0.696 | FE(OH)4- | 0.686 | ALF++ | 0.226 |
| S-- | 0.217 | MGHCO3+ | 0.662 | FESO4+ | 0.680 | ALF2+ | 0.686 |
| HSO4- | 0.671 | CAOH+ | 0.696 | FECL++ | 0.217 | ALF4- | 0.671 |
| SO4-- | 0.195 | MGOH+ | 0.703 | FECL2+ | 0.680 | ALF5-- | 0.206 |
| NASO4- | 0.686 | NH4+ | 0.630 | FECL4- | 0.662 | ALF6--- | 0.029 |

PH LOWER THAN 4.00

CHEMICAL COMPONENTS IN DEEP WATER (PPM AND LOG MOLE)

| | | | | | | | | |
|-----------|---------|---------|-----------|--------|--------|-----------|------|-------|
| H+ (ACT.) | 0.01 | -4.902 | MG++ | 2.77 | -3.944 | FE(OH)3 | 0.00 | 0.000 |
| OH- | 0.02 | -5.982 | NACL | 135.56 | -2.635 | FE(OH)4- | 0.00 | 0.000 |
| H4SiO4 | 690.82 | -2.143 | KCL | 13.52 | -3.741 | FECL+ | 0.00 | 0.000 |
| H3SiO4- | 0.08 | -6.093 | NASO4- | 19.41 | -3.788 | FECL2 | 0.00 | 0.000 |
| H2SiO4-- | 0.00 | -12.110 | KSO4- | 17.70 | -3.883 | FECL++ | 0.00 | 0.000 |
| NAH3SiO4 | 0.04 | -6.444 | CASO4 | 4.71 | -4.461 | FECL2+ | 0.00 | 0.000 |
| H3BO3 | 193.06 | -2.506 | MESO4 | 4.88 | -4.392 | FECL3 | 0.00 | 0.000 |
| H2BO3- | 0.02 | -6.476 | CACO3 | 0.00 | -8.031 | FECL4- | 0.00 | 0.000 |
| H2CO3 | 5236.00 | -1.074 | MGCO3 | 0.00 | -9.555 | FESO4 | 0.00 | 0.000 |
| HCO3- | 16.52 | -3.568 | CAHCO3+ | 2.52 | -4.603 | FESO4+ | 0.00 | 0.000 |
| CO3-- | 0.00 | -9.280 | MGHCO3+ | 0.07 | -6.057 | AL+++ | 0.00 | 0.000 |
| H2S | 1213.43 | -1.448 | CAOH+ | 0.00 | -7.763 | ALOH++ | 0.00 | 0.000 |
| HS- | 6.63 | -3.698 | MGOH+ | 0.01 | -6.640 | AL(OH)2+ | 0.00 | 0.000 |
| S-- | 0.00 | -13.300 | NH4OH | 0.00 | 0.000 | AL(OH)3 | 0.00 | 0.000 |
| H2SO4 | 0.00 | -9.380 | NH4+ | 0.00 | 0.000 | AL(OH)4- | 0.00 | 0.000 |
| HSO4- | 26.51 | -3.564 | FE++ | 0.00 | 0.000 | ALSO4+ | 0.00 | 0.000 |
| SO4-- | 49.88 | -3.285 | FE+++ | 0.00 | 0.000 | AL(SO4)2- | 0.00 | 0.000 |
| HF | 0.00 | 0.000 | FEOH+ | 0.00 | 0.000 | ALF++ | 0.00 | 0.000 |
| F- | 0.00 | 0.000 | FE(OH)2 | 0.00 | 0.000 | ALF2+ | 0.00 | 0.000 |
| CL- | 2422.67 | -1.165 | FE(OH)3- | 0.00 | 0.000 | ALF3 | 0.00 | 0.000 |
| NA+ | 1332.21 | -1.237 | FE(OH)4-- | 0.00 | 0.000 | ALF4- | 0.00 | 0.000 |
| K+ | 270.67 | -2.160 | FE(OH)++ | 0.00 | 0.000 | ALF5-- | 0.00 | 0.000 |
| CA++ | 9.71 | -3.616 | FE(OH)2+ | 0.00 | 0.000 | ALF6--- | 0.00 | 0.000 |

IONIC STRENGTH = 0.06861 IONIC BALANCE : CATIONS (MOL.EQ.)0.06560814
 ANIONS (MOL.EQ.)0.06985241
 DIFFERENCE (%) -6.27

CHEMICAL GEOTHERMOMETERS DEGREES C 1000/T DEGREES KELVIN = 1.94

| | |
|------------|-------|
| QUARTZ | 239.5 |
| CHALCEDONY | 999.9 |
| NAK | 284.3 |

OXIDATION POTENTIAL (VOLTS) : EH H2S= -0.355 EH CH4= 99.999 EH H2= 99.999 EH NH3= 99.999

LOG SOLUBILITY PRODUCTS OF MINERALS IN DEEP WATER

| | TEOR. | CALC. | | TEOR. | CALC. | | TEOR. | CALC. |
|-------------|---------|--------|--------------|---------|---------|-------------|---------|--------|
| ADULARIA | -14.428 | 99.999 | ALBITE LOW | -13.975 | 99.999 | ANALCINE | -11.524 | 99.999 |
| ANHYDRITE | -7.978 | -8.233 | CALCITE | -12.591 | -14.204 | CHALCEDONY | -2.023 | -2.143 |
| MG-CHLORITE | -84.354 | 99.999 | FLUORITE | -10.920 | 99.999 | GOETHITE | 1.889 | 99.999 |
| LAUMONTITE | -24.536 | 99.999 | MICROCLINE | -15.119 | 99.999 | MAGNETITE | -17.601 | 99.999 |
| CA-MONTMOR. | -72.623 | 99.999 | K-MONTMOR. | -34.027 | 99.999 | MG-MONTMOR. | -74.117 | 99.999 |
| NA-MONTMOR. | -34.300 | 99.999 | MUSCOVITE | -17.852 | 99.999 | PREHNITE | -37.100 | 99.999 |
| PYRRHOTITE | -24.961 | 99.999 | PYRITE | -46.345 | 99.999 | QUARTZ | -2.135 | -2.143 |
| WAIRAKITE | -24.425 | 99.999 | WOLLASTONITE | 7.602 | 3.422 | ZOISITE | -37.653 | 99.999 |
| EPIDOTE | -77.210 | 99.999 | MARGARITE | -70.751 | 99.999 | | | |

ORKUSTOFNUN JHD
1982-11-12 OLIVER

GEOCHEMISTRY OF WELL PN-13D

*****820127**** PUHAGAN-13D

S, NEGROS, PHILIPPINES PUHAGAN

PROGRAM WATCH1.

WATER SAMPLE (PPM)

STEAM SAMPLE

| | | | | | |
|----------|-----------|-------------|-----------------|----------------------|---------------------------|
| PH/DEG.C | 5.95/25.0 | GAS (VOL.%) | REFERENCE TEMP. | DEGREES C | 268.0 (MEASURED) |
| SI02 | 675.00 | CO2 | 0.00 | | |
| NA | 2472.00 | H2S | 0.00 | SAMPLING PRESSURE | BARS ABS. 11.5 |
| K | 551.00 | H2 | 0.00 | DISCHARGE ENTHALPY | MJ/OL/KG 1.202 (MEASURED) |
| CA | 58.30 | O2 | 0.00 | DISCHARGE | KG/SEC. 14.1 |
| MG | 0.380 | CH4 | 0.00 | | |
| CO2 | 18.30 | N2 | 0.00 | MEASURED TEMPERATURE | DEGREES C 0.0 |
| SO4 | 24.90 | | | RESISTIVITY/TEMP. | OHMM/DEG.C 0.0/ 0.0 |
| H2S | 10.20 | | | EH/TEMP. | MV/DEG.C 0.000/ 0.0 |
| CL | 4773.00 | | | | |

F 0.00
DISS.SOLIDS 0.00
AL 0.0000
B 61.3000
FE 0.0000
NH3 0.0000

LITERS GAS PER KG
CONDENSATE/DEG.C 13.00/25.0

MEASURED DOWNHOLE TEMP.
DEGREES C/METERS

FLUID INFLOW
DEPTH (METERS)

| | | | |
|----------------------------|-----------|-----|-----|
| CONDENSATE (PPM) | 0.0 | 0.0 | 0.0 |
| PH/DEG.C | 0.00/ 0.0 | 0.0 | 0.0 |
| CO2 | 0.00 | 0.0 | 0.0 |
| H2S | 0.00 | 0.0 | 0.0 |
| NA | 0.00 | 0.0 | 0.0 |
| | | 0.0 | 0.0 |
| | | 0.0 | 0.0 |
| | | 0.0 | 0.0 |
| CONDENSATE WITH NaOH (PPM) | 0.0 | 0.0 | 0.0 |
| CO2 | 23081.90 | 0.0 | 0.0 |
| H2S | 646.30 | 0.0 | 0.0 |

IONIC STRENGTH = 0.13136 IONIC BALANCE : CATIONS (MOL.EQ.)0.12428541
ANIONS (MOL.EQ.)0.13502967
DIFFERENCE (%) -8.29

DEEP WATER (PPM)

DEEP STEAM (PPM)

GAS PRESSURES (BARS ABS.)

| | | | | | | | |
|---------|---------|-----|---------|-----------------|-----------|-------|-----------|
| SI02 | 544.67 | CO2 | 1704.92 | CO2 | 178562.48 | CO2 | 0.390E+01 |
| NA | 1994.52 | H2S | 92.14 | H2S | 2944.96 | H2S | 0.830E-01 |
| K | 444.55 | H2 | 0.00 | H2 | 0.00 | H2 | 0.000E+00 |
| CA | 47.04 | O2 | 0.00 | O2 | 0.00 | O2 | 0.000E+00 |
| MG | 0.307 | CH4 | 0.00 | CH4 | 0.00 | CH4 | 0.000E+00 |
| SO4 | 20.09 | N2 | 0.00 | N2 | 0.00 | N2 | 0.000E+00 |
| CL | 3850.75 | NH3 | 0.00 | NH3 | 0.00 | NH3 | 0.000E+00 |
| F | 0.00 | | | | | H2O | 0.533E+02 |
| DISS.S. | 0.00 | | | | | TOTAL | 0.573E+02 |
| AL | 0.0000 | | | | | | |
| B | 49.4550 | | | H2O (%) | 1.75 | | |
| FE | 0.0000 | | | BOILING PORTION | 0.19 | | |

GAS SOLUBILITY MULTIPLYING FACTOR : 1.00

ACTIVITY COEFFICIENTS IN DEEP WATER

| | | | | | | | |
|----------|-------|---------|-------|-----------|-------|-----------|-------|
| H+ | 0.692 | KS04- | 0.629 | FE++ | 0.172 | FECL+ | 0.597 |
| OH- | 0.584 | F- | 0.584 | FE+++ | 0.038 | AL+++ | 0.038 |
| H3SI04- | 0.597 | CL- | 0.571 | FE0H+ | 0.621 | AL0H++ | 0.159 |
| H2SI04-- | 0.159 | NA+ | 0.597 | FE(OH)3- | 0.621 | AL(OH)2+ | 0.629 |
| H2B03- | 0.556 | K+ | 0.571 | FE(OH)4-- | 0.150 | AL(OH)4- | 0.609 |
| HCO3- | 0.597 | CA++ | 0.172 | FE0H++ | 0.150 | ALB04+ | 0.609 |
| CO3-- | 0.140 | MG++ | 0.213 | FE(OH)2+ | 0.629 | AL(S04)2- | 0.609 |
| HS- | 0.584 | CAHCO3+ | 0.642 | FE(OH)4- | 0.629 | ALF++ | 0.159 |
| S-- | 0.150 | MGHCO3+ | 0.597 | FES04+ | 0.621 | ALF2+ | 0.629 |
| HS04- | 0.609 | CAOH+ | 0.642 | FECL++ | 0.150 | ALF4- | 0.609 |
| S04-- | 0.129 | MGOH+ | 0.651 | FECL2+ | 0.621 | ALF5-- | 0.140 |
| NAS04- | 0.629 | NH4+ | 0.556 | FECL4- | 0.597 | ALF6--- | 0.012 |

CHEMICAL COMPONENTS IN DEEP WATER (PPM AND LOG MOLE)

| | | | | | | | | |
|-----------|---------|---------|-----------|--------|---------|-----------|------|-------|
| H+ (ACT.) | 0.01 | -4.979 | MG++ | 0.29 | -4.917 | FE(OH)3 | 0.00 | 0.000 |
| OH- | 0.03 | -5.808 | NACL | 427.67 | -2.136 | FE(OH)4- | 0.00 | 0.000 |
| H4SI04 | 871.11 | -2.043 | KCL | 42.87 | -3.240 | FECL+ | 0.00 | 0.000 |
| H3SI04- | 0.08 | -6.049 | NAS04- | 3.43 | -4.541 | FECL2 | 0.00 | 0.000 |
| H2SI04-- | 0.00 | -12.083 | KS04- | 4.38 | -4.490 | FECL++ | 0.00 | 0.000 |
| NAH3SI04 | 0.06 | -6.304 | CAS04 | 2.52 | -4.733 | FECL2+ | 0.00 | 0.000 |
| H3B03 | 282.83 | -2.340 | MGS04 | 0.05 | -6.387 | FECL3 | 0.00 | 0.000 |
| H2B03- | 0.03 | -6.247 | CACO3 | 0.00 | -7.807 | FECL4- | 0.00 | 0.000 |
| H2CO3 | 2395.35 | -1.413 | MGCO3 | 0.00 | -11.083 | FES04 | 0.00 | 0.000 |
| HCO3- | 5.13 | -4.075 | CAHCO3+ | 3.97 | -4.406 | FES04+ | 0.00 | 0.000 |
| CO3-- | 0.00 | -9.863 | MGHCO3+ | 0.00 | -7.512 | AL+++ | 0.00 | 0.000 |
| H2S | 91.78 | -2.570 | CAOH+ | 0.01 | -6.893 | AL0H++ | 0.00 | 0.000 |
| HS- | 0.35 | -4.980 | MGOH+ | 0.00 | -7.191 | AL(OH)2+ | 0.00 | 0.000 |
| S-- | 0.00 | -14.260 | NH4OH | 0.00 | 0.000 | AL(OH)3 | 0.00 | 0.000 |
| H2S04 | 0.00 | -10.029 | NH4+ | 0.00 | 0.000 | AL(OH)4- | 0.00 | 0.000 |
| HS04- | 5.59 | -4.239 | FE++ | 0.00 | 0.000 | ALB04+ | 0.00 | 0.000 |
| S04-- | 6.86 | -4.146 | FE+++ | 0.00 | 0.000 | AL(S04)2- | 0.00 | 0.000 |
| HF | 0.00 | 0.000 | FE0H+ | 0.00 | 0.000 | ALF++ | 0.00 | 0.000 |
| F- | 0.00 | 0.000 | FE(OH)2 | 0.00 | 0.000 | ALF2+ | 0.00 | 0.000 |
| CL- | 3570.94 | -0.997 | FE(OH)3- | 0.00 | 0.000 | ALF3 | 0.00 | 0.000 |
| NA+ | 1825.60 | -1.100 | FE(OH)4-- | 0.00 | 0.000 | ALF4- | 0.00 | 0.000 |
| K+ | 420.80 | -1.968 | FE(OH)++ | 0.00 | 0.000 | ALF5-- | 0.00 | 0.000 |
| CA++ | 44.72 | -2.952 | FE(OH)2+ | 0.00 | 0.000 | ALF6--- | 0.00 | 0.000 |

IONIC STRENGTH = 0.09792 IONIC BALANCE : CATIONS (MOL.EQ.)0.09246597
 ANIONS (MOL.EQ.)0.10097204
 DIFFERENCE (%) -8.79

CHEMICAL GEOTHERMOMETERS DEGREES C

1000/T DEGREES KELVIN = 1.85

QUARTZ 265.1
 CHALCEDONY 999.9
 MAK 301.1

OXIDATION POTENTIAL (VOLTS) : EH H2S= -0.390 EH CH4= 99.999 EH H2= 99.999 EH NH3= 99.999

LOG SOLUBILITY PRODUCTS OF MINERALS IN DEEP WATER

| | TEOR. | CALC. | | TEOR. | CALC. | | TEOR. | CALC. |
|-------------|---------|--------|--------------|---------|---------|-------------|---------|--------|
| ADULARIA | -14.371 | 99.999 | ALBITE LOW | -13.938 | 99.999 | ANALCINE | -11.566 | 99.999 |
| ANHYDRITE | -8.441 | -8.754 | CALCITE | -13.242 | -14.436 | CHALCEDONY | -1.925 | -2.043 |
| MG-CHLORITE | -86.372 | 99.999 | FLUORITE | -11.057 | 99.999 | GOETHITE | 3.286 | 99.999 |
| LAUMONTITE | -24.717 | 99.999 | MICROCLINE | -14.996 | 99.999 | MAGNETITE | -15.172 | 99.999 |
| CA-MONTMOR. | -72.580 | 99.999 | K-MONTMOR. | -33.864 | 99.999 | MG-MONTMOR. | -74.086 | 99.999 |
| NA-MONTMOR. | -34.146 | 99.999 | MUSCOVITE | -17.838 | 99.999 | PREHNITE | -37.960 | 99.999 |
| PYRRHOTITE | -10.789 | 99.999 | PYRITE | -28.408 | 99.999 | QUARTZ | -2.031 | -2.043 |
| WAIKAKITE | -24.913 | 99.999 | MOLLASTONITE | 7.189 | 4.197 | ZOISITE | -38.801 | 99.999 |
| EPIDOTE | -38.439 | 99.999 | MARCASITE | -12.208 | 99.999 | | | |

ACTIVITY COEFFICIENTS IN DEEP WATER

| | | | | | | | |
|----------|-------|---------|-------|-----------|-------|-----------|-------|
| H+ | 0.692 | KS04- | 0.628 | FE++ | 0.168 | FECL+ | 0.596 |
| OH- | 0.583 | F- | 0.583 | FE+++ | 0.036 | AL+++ | 0.036 |
| H3S104- | 0.596 | CL- | 0.570 | FE0H+ | 0.620 | AL0H++ | 0.153 |
| H2S104-- | 0.155 | NA+ | 0.596 | FE(OH)3- | 0.620 | AL(OH)2+ | 0.628 |
| H2B03- | 0.555 | K+ | 0.570 | FE(OH)4-- | 0.147 | AL(OH)4- | 0.608 |
| HCO3- | 0.596 | CA++ | 0.168 | FE0H++ | 0.147 | ALS04+ | 0.608 |
| CO3-- | 0.137 | MG++ | 0.209 | FE(OH)2+ | 0.628 | AL(S04)2- | 0.608 |
| HS- | 0.583 | CAHCO3+ | 0.641 | FE(OH)4- | 0.628 | ALF++ | 0.155 |
| S-- | 0.147 | MGHCO3+ | 0.596 | FES04+ | 0.620 | ALF2+ | 0.628 |
| HS04- | 0.608 | CAOH+ | 0.641 | FECL++ | 0.147 | ALF4- | 0.608 |
| S04-- | 0.126 | MGOH+ | 0.650 | FECL2+ | 0.620 | ALF5-- | 0.137 |
| NAS04- | 0.628 | NH4+ | 0.555 | FECL4- | 0.596 | ALF6--- | 0.011 |

CHEMICAL COMPONENTS IN DEEP WATER (PPM AND LOG MOLE)

| | | | | | | | | |
|-----------|---------|---------|-----------|--------|--------|-----------|------|-------|
| H+ (ACT.) | 0.00 | -5.398 | MG++ | 0.18 | -5.140 | FE(OH)3 | 0.00 | 0.000 |
| OH- | 0.07 | -5.384 | NACL | 486.54 | -2.080 | FE(OH)4- | 0.00 | 0.000 |
| H4S104 | 882.49 | -2.037 | KCL | 37.28 | -3.301 | FECL+ | 0.00 | 0.000 |
| H3S104- | 0.20 | -5.672 | NAS04- | 4.10 | -4.463 | FECL2 | 0.00 | 0.000 |
| H2S104-- | 0.00 | -11.331 | KS04- | 4.51 | -4.477 | FECL++ | 0.00 | 0.000 |
| NAH3S104 | 0.14 | -5.911 | CAS04 | 2.39 | -4.755 | FECL2+ | 0.00 | 0.000 |
| H3B03 | 237.57 | -2.415 | MBS04 | 0.03 | -6.554 | FECL3 | 0.00 | 0.000 |
| H2B03- | 0.07 | -5.921 | CACO3 | 0.03 | -6.574 | FECL4- | 0.00 | 0.000 |
| H2CO3 | 8319.67 | -0.872 | MGCO3 | 0.00 | -9.999 | FES04 | 0.00 | 0.000 |
| HCO3- | 39.66 | -3.187 | CAHCO3+ | 26.06 | -3.589 | FES04+ | 0.00 | 0.000 |
| CO3-- | 0.00 | -8.617 | MGHCO3+ | 0.01 | -6.820 | AL+++ | 0.00 | 0.000 |
| H2S | 87.75 | -2.589 | CAOH+ | 0.02 | -6.539 | AL0H++ | 0.00 | 0.000 |
| HS- | 0.73 | -4.653 | MGOH+ | 0.01 | -6.899 | AL(OH)2+ | 0.00 | 0.000 |
| S-- | 0.00 | -13.476 | NH4OH | 0.00 | 0.000 | AL(OH)3 | 0.00 | 0.000 |
| H2S04 | 0.00 | -10.703 | NH4+ | 0.00 | 0.000 | AL(OH)4- | 0.00 | 0.000 |
| HS04- | 2.88 | -4.527 | FE++ | 0.00 | 0.000 | ALS04+ | 0.00 | 0.000 |
| S04-- | 7.58 | -4.103 | FE+++ | 0.00 | 0.000 | AL(S04)2- | 0.00 | 0.000 |
| HF | 0.00 | 0.000 | FE0H+ | 0.00 | 0.000 | ALF++ | 0.00 | 0.000 |
| F- | 0.00 | 0.000 | FE(OH)2 | 0.00 | 0.000 | ALF2+ | 0.00 | 0.000 |
| CL- | 3306.31 | -1.030 | FE(OH)3- | 0.00 | 0.000 | ALF3 | 0.00 | 0.000 |
| NA+ | 1848.97 | -1.095 | FE(OH)4-- | 0.00 | 0.000 | ALF4- | 0.00 | 0.000 |
| K+ | 341.44 | -2.059 | FE(OH)++ | 0.00 | 0.000 | ALF5-- | 0.00 | 0.000 |
| CA++ | 34.94 | -3.060 | FE(OH)2+ | 0.00 | 0.000 | ALF6--- | 0.00 | 0.000 |

IONIC STRENGTH = 0.09360 IONIC BALANCE : CATIONS (MOL.EQ.)0.09117395
 ANIONS (MOL.EQ.)0.09410425
 DIFFERENCE (%) -3.16

CHEMICAL GEOTHERMOMETERS DEGREES C 1000/T DEGREES KELVIN = 1.83

QUARTZ 266.5
 CHALCEDONY 999.9
 NAK 268.6

OXIDATION POTENTIAL (VOLTS) : EH H2S= -0.452 EH CH4= 99.999 EH H2= 99.999 EH NH3= 99.999

LOG SOLUBILITY PRODUCTS OF MINERALS IN DEEP WATER

| | TEOR. | CALC. | | TEOR. | CALC. | | TEOR. | CALC. |
|-------------|---------|--------|--------------|---------|---------|-------------|---------|--------|
| ADULARIA | -14.372 | 99.999 | ALBITE LOW | -13.943 | 99.999 | ANALCIME | -11.587 | 99.999 |
| AMHYDRITE | -8.554 | -8.837 | CALCITE | -13.401 | -13.316 | CHALCEDONY | -1.902 | -2.037 |
| MG-CHLORITE | -86.903 | 99.999 | FLUORITE | -11.091 | 99.999 | GOETHITE | 3.631 | 99.999 |
| LAUMONTITE | -24.782 | 99.999 | MICROCLINE | -14.982 | 99.999 | MAGNETITE | -14.576 | 99.999 |
| CA-MONTMOR. | -72.580 | 99.999 | K-MONTMOR. | -33.830 | 99.999 | MG-MONTMOR. | -74.084 | 99.999 |
| NA-MONTMOR. | -34.115 | 99.999 | MUSCOVITE | -17.840 | 99.999 | PREHNITE | -38.193 | 99.999 |
| PYRRHOTITE | -7.333 | 99.999 | PYRITE | -24.060 | 99.999 | QUARTZ | -2.007 | -2.037 |
| WAIKAKITE | -25.049 | 99.999 | MOLLASTONITE | 7.094 | 4.924 | ZOISITE | -39.106 | 99.999 |
| EPIDOTE | -38.871 | 99.999 | MARCASITE | -8.041 | 99.999 | | | |

ORKUSTOFNUN JHD
1982-11-12 OLIVER

GEOCHEMISTRY OF WELL PN-17D

XXXXXXXXXX820131XXXX PUHAGAN-17D

S.NEGROS,PHILIPPINES PUHAGAN

PROGRAM WATCH1.

WATER SAMPLE (PPM)

STEAM SAMPLE

| | | | | | |
|----------|-----------|-------------|-----------------|----------------------|----------------------------|
| PH/DEG.C | 6.21/25.0 | GAS (VOL.%) | REFERENCE TEMP. | DEGREES C | 245.0 (MEASURED) |
| SI02 | 559.00 | CO2 | 0.00 | | |
| NA | 2965.00 | H2S | 0.00 | SAMPLING PRESSURE | BARS ABS. 6.6 |
| K | 421.00 | H2 | 0.00 | DISCHARGE ENTHALPY | MJ/0UL/KG 1.094 (MEASURED) |
| CA | 68.60 | O2 | 0.00 | DISCHARGE | KG/SEC. 33.4 |
| MG | 0.420 | CH4 | 0.00 | | |
| CO2 | 40.20 | N2 | 0.00 | MEASURED TEMPERATURE | DEGREES C 0.0 |
| SO4 | 30.60 | | | RESISTIVITY/TEMP. | OHMM/DEG.C 0.0/ 0.0 |
| H2S | 7.16 | | | EH/TEMP. | MV/DEG.C 0.000/ 0.0 |
| CL | 4617.00 | | | | |

| | | | | | |
|-------------|---------|-------------------|------------|-------------------------|----------------|
| F | 0.00 | LITERS GAS PER KG | | | |
| DISS.SOLIDS | 0.00 | CONDENSATE/DEG.C | 10.40/25.0 | MEASURED DOWNHOLE TEMP. | FLUID INFLOW |
| AL | 0.0000 | | | DEGREES C/METERS | DEPTH (METERS) |
| B | 53.3000 | | | | |

| | | | | | |
|-----|--------|----------------------------|-----------|-----|-----|
| FE | 0.0000 | CONDENSATE (PPM) | 0.0 | 0.0 | 0.0 |
| NH3 | 0.0000 | PH/DEG.C | 0.00/ 0.0 | 0.0 | 0.0 |
| | | CO2 | 0.00 | 0.0 | 0.0 |
| | | H2S | 0.00 | 0.0 | 0.0 |
| | | NA | 0.00 | 0.0 | 0.0 |
| | | | 0.0 | 0.0 | 0.0 |
| | | | 0.0 | 0.0 | 0.0 |
| | | | 0.0 | 0.0 | 0.0 |
| | | CONDENSATE WITH NAOH (PPM) | 0.0 | 0.0 | 0.0 |
| | | CO2 | 19184.00 | 0.0 | 0.0 |
| | | H2S | 327.80 | 0.0 | 0.0 |

IONIC STRENGTH = 0.13896 IONIC BALANCE :

| | |
|-------------------|------------|
| CATIONS (MOL.EQ.) | 0.14288795 |
| ANIONS (MOL.EQ.) | 0.13101509 |
| DIFFERENCE (%) | 8.67 |

| | | | | | |
|------------------|---------|------------------|--------|---------------------------|-----------|
| DEEP WATER (PPM) | | DEEP STEAM (PPM) | | GAS PRESSURES (BARS ABS.) | |
| SI02 | 457.71 | CO2 | 874.01 | CO2 | 0.226E+01 |
| NA | 2427.52 | H2S | 34.86 | H2S | 0.358E-01 |
| K | 344.67 | H2 | 0.00 | H2 | 0.000E+00 |
| CA | 56.16 | O2 | 0.00 | O2 | 0.000E+00 |
| MG | 0.344 | CH4 | 0.00 | CH4 | 0.000E+00 |
| SO4 | 25.05 | N2 | 0.00 | N2 | 0.000E+00 |
| CL | 3779.73 | NH3 | 0.00 | NH3 | 0.000E+00 |
| F | 0.00 | | | H2O | 0.365E+02 |
| DISS.S. | 0.00 | | | TOTAL | 0.388E+02 |
| AL | 0.0000 | | | | |
| B | 43.6340 | | | H2O (%) | 1.96 |
| FE | 0.0000 | | | BOILING PORTION | 0.18 |

GAS SOLUBILITY MULTIPLYING FACTOR : 1.00

ACTIVITY COEFFICIENTS IN DEEP WATER

| | | | | | | | |
|----------|-------|---------|-------|-----------|-------|-----------|-------|
| H+ | 0.706 | KS04- | 0.645 | FE++ | 0.191 | FECL+ | 0.613 |
| OH- | 0.601 | F- | 0.601 | FE+++ | 0.047 | AL+++ | 0.047 |
| H3SiO4- | 0.613 | CL- | 0.587 | FE0H+ | 0.636 | AL0H++ | 0.177 |
| H2SiO4-- | 0.177 | NA+ | 0.613 | FE(OH)3- | 0.636 | AL(OH)2+ | 0.645 |
| H2B03- | 0.573 | K+ | 0.587 | FE(OH)4-- | 0.168 | AL(OH)4- | 0.625 |
| HCO3- | 0.613 | CA++ | 0.191 | FE0H++ | 0.168 | ALSO4+ | 0.625 |
| CO3-- | 0.157 | MG++ | 0.234 | FE(OH)2+ | 0.645 | AL(SO4)2- | 0.625 |
| HS- | 0.601 | CAHCO3+ | 0.656 | FE(OH)4- | 0.645 | ALF++ | 0.177 |
| S-- | 0.168 | MGHCO3+ | 0.613 | FES04+ | 0.636 | ALF2+ | 0.645 |
| HSO4- | 0.625 | CAOH+ | 0.656 | FECL++ | 0.168 | ALF4- | 0.625 |
| SO4-- | 0.145 | MGOH+ | 0.666 | FECL2+ | 0.636 | ALF5-- | 0.157 |
| NASO4- | 0.645 | NH4+ | 0.573 | FECL4- | 0.613 | ALF6--- | 0.016 |

CHEMICAL COMPONENTS IN DEEP WATER (PPM AND LOG MOLE)

| | | | | | | | | |
|-----------|---------|---------|-----------|--------|---------|-----------|------|-------|
| H+ (ACT.) | 0.00 | -5.499 | MG++ | 0.32 | -4.880 | FE(OH)3 | 0.00 | 0.000 |
| OH- | 0.08 | -5.343 | NACL | 307.64 | -2.279 | FE(OH)4- | 0.00 | 0.000 |
| H4SiO4 | 731.58 | -2.119 | KCL | 21.44 | -3.541 | FECL+ | 0.00 | 0.000 |
| H3SiO4- | 0.34 | -5.451 | NASO4- | 6.43 | -4.267 | FECL2 | 0.00 | 0.000 |
| H2SiO4-- | 0.00 | -10.813 | KS04- | 4.17 | -4.511 | FECL++ | 0.00 | 0.000 |
| NAH3SiO4 | 0.28 | -5.626 | CASO4 | 3.90 | -4.543 | FECL2+ | 0.00 | 0.000 |
| H3B03 | 249.46 | -2.394 | HGSO4 | 0.09 | -6.117 | FECL3 | 0.00 | 0.000 |
| H2B03- | 0.11 | -5.731 | CACO3 | 0.01 | -6.858 | FECL4- | 0.00 | 0.000 |
| H2CO3 | 1209.94 | -1.710 | MGCO3 | 0.00 | -10.028 | FES04 | 0.00 | 0.000 |
| HCO3- | 15.46 | -3.596 | CAHCO3+ | 10.12 | -4.000 | FES04+ | 0.00 | 0.000 |
| CO3-- | 0.00 | -8.647 | MGHCO3+ | 0.01 | -7.086 | AL+++ | 0.00 | 0.000 |
| H2S | 34.07 | -3.000 | CAOH+ | 0.02 | -6.497 | AL0H++ | 0.00 | 0.000 |
| HS- | 0.76 | -4.638 | MGOH+ | 0.00 | -6.997 | AL(OH)2+ | 0.00 | 0.000 |
| S-- | 0.00 | -13.557 | NH4OH | 0.00 | 0.000 | AL(OH)3 | 0.00 | 0.000 |
| H2SO4 | 0.00 | -11.255 | NH4+ | 0.00 | 0.000 | AL(OH)4- | 0.00 | 0.000 |
| HSO4- | 1.46 | -4.823 | FE++ | 0.00 | 0.000 | ALSO4+ | 0.00 | 0.000 |
| SO4-- | 12.63 | -3.881 | FE+++ | 0.00 | 0.000 | AL(SO4)2- | 0.00 | 0.000 |
| HF | 0.00 | 0.000 | FE0H+ | 0.00 | 0.000 | ALF++ | 0.00 | 0.000 |
| F- | 0.00 | 0.000 | FE(OH)2 | 0.00 | 0.000 | ALF2+ | 0.00 | 0.000 |
| CL- | 3582.92 | -0.995 | FE(OH)3- | 0.00 | 0.000 | ALF3 | 0.00 | 0.000 |
| NA+ | 2305.20 | -0.999 | FE(OH)4-- | 0.00 | 0.000 | ALF4- | 0.00 | 0.000 |
| K+ | 332.22 | -2.071 | FE(OH)++ | 0.00 | 0.000 | ALF5-- | 0.00 | 0.000 |
| CA++ | 50.99 | -2.895 | FE(OH)2+ | 0.00 | 0.000 | ALF6--- | 0.00 | 0.000 |

IONIC STRENGTH = 0.10795 IONIC BALANCE : CATIONS (MOL.EQ.)0.11143742
 ANIONS (MOL.EQ.)0.10161908
 DIFFERENCE (%) 9.22

CHEMICAL GEOTHERMOMETERS DEGREES C 1000/T DEGREES KELVIN = 1.93

| | |
|------------|-------|
| QUARTZ | 244.5 |
| CHALCEDONY | 999.9 |
| NAK | 236.5 |

OXIDATION POTENTIAL (VOLTS) : EH H2S= -0.423 EH CH4= 99.999 EH H2= 99.999 EH NH3= 99.999

LOG SOLUBILITY PRODUCTS OF MINERALS IN DEEP WATER

| | TEOR. | CALC. | | TEOR. | CALC. | | TEOR. | CALC. |
|-------------|---------|--------|--------------|---------|---------|-------------|---------|--------|
| ADULARIA | -14.420 | 99.999 | ALBITE LOW | -13.969 | 99.999 | ANALCIME | -11.525 | 99.999 |
| ANHYDRITE | -8.014 | -8.334 | CALCITE | -12.642 | -13.067 | CHALCEDONY | -2.015 | -2.119 |
| MG-CHLORITE | -84.503 | 99.999 | FLUORITE | -10.931 | 99.999 | GOETHITE | 1.998 | 99.999 |
| LAUNMONTITE | -24.545 | 99.999 | MICROCLINE | -15.105 | 99.999 | MAGNETITE | -17.410 | 99.999 |
| CA-MONTHOR. | -72.613 | 99.999 | K-MONTHOR. | -34.012 | 99.999 | MG-MONTHOR. | -74.113 | 99.999 |
| NA-MONTHOR. | -34.286 | 99.999 | MUSCOVITE | -17.848 | 99.999 | PREHNITE | -37.162 | 99.999 |
| PYRRHOTITE | -23.841 | 99.999 | PYRITE | -44.921 | 99.999 | QUARTZ | -2.126 | -2.119 |
| HAIRAKITE | -24.459 | 99.999 | WOLLASTONITE | 7.568 | 5.265 | ZOISITE | -37.737 | 99.999 |
| EPIDOTE | -37.318 | 99.999 | NARCASITE | -27.993 | 99.999 | | | |